

1 **An assessment of geographical distribution of different plant functional types**  
2 **over North America simulated using the CLASS-CTEM modelling**  
3 **framework**

4

5 **Rudra K. Shrestha<sup>1</sup>, Vivek K. Arora<sup>1</sup>, Joe R. Melton<sup>2</sup>, and Laxmi Sushama<sup>3</sup>**

6 <sup>1</sup>Canadian Centre for Climate Modelling and Analysis, Environment and Climate Change Canada,  
7 University of Victoria, Victoria, BC, V8W 2Y2, Canada

8 <sup>2</sup>Climate Research Division, Environment and Climate Change Canada, Toronto, Ontario, Canada

9 <sup>3</sup>Département des sciences de la Terre et de l'atmosphère, Université du Québec à Montréal, Canada

10

11 *Correspondence to: V. K. Arora (Vivek.Arora@canada.ca)*

12

13 **Abstract**

14

15 The performance of the competition module of the CLASS-CTEM (Canadian Land Surface  
16 Scheme and Canadian Terrestrial Ecosystem Model) modelling framework is assessed at 1°  
17 spatial resolution over North America by comparing the simulated geographical distribution of  
18 **its** plant functional types (PFTs) with two observation-based estimates. The model successfully  
19 reproduces the broad geographical distribution of trees, grasses and bare ground although  
20 limitations remain. In particular, compared to the two observation-based estimates, the simulated  
21 fractional vegetation coverage is lower in the arid south-west North American region and higher  
22 in the Arctic region. The lower than observed simulated vegetation coverage in the south-west  
23 region is attributed to lack of representation of shrubs in the model and plausible errors in the  
24 observation-based data sets. The observation-based data indicates vegetation fractional coverage  
25 of more than 60% in this arid region, despite only 200-300 mm of precipitation that the region  
26 receives annually and observation-based leaf area index (LAI) **values** in the region are lower than  
27 one. The higher than observed vegetation fractional coverage in the Arctic is **likely** due to the  
28 lack of representation of moss and lichen PFTs and also likely because of inadequate  
29 representation of permafrost in the model as a result of which the C<sub>3</sub> grass PFT performs overly  
30 well in the region. The model generally reproduces the broad spatial distribution and the total  
31 area covered by the two primary tree PFTs (needleleaf evergreen and broadleaf cold deciduous

32 trees) reasonably well. The simulated fractional coverage of tree PFTs increases after 1960s in  
33 response to the CO<sub>2</sub> fertilization effect and climate warming. Differences between observed and  
34 simulated PFT coverages highlight model limitations ~~and suggest that the inclusion of shrubs,~~  
35 ~~and moss and lichen PFTs, and an adequate representation of permafrost will help improve~~  
36 ~~model performance.~~

**Deleted:** in the model

**Deleted:** and

**Deleted:** provide insight into physical and structural processes that need improvement

37

38

## 44 1 Introduction

45

46 The terrestrial ecosystem plays an important role in regulating climate and weather through land-  
47 atmosphere exchange of water and energy (Cramer et al., 2001; Garraud et al., 2015; Pielke et  
48 al., 1998; Ran et al., 2016) and in mitigating climate change by sequestering atmospheric CO<sub>2</sub>  
49 (Bonan, 2008; Timmons et al., 2016). The projected sink of atmospheric CO<sub>2</sub> is uncertain due to  
50 disagreements among the Earth system models (ESMs) (Arora et al., 2013; Friedlingstein et al.,  
51 2006) primarily due to differing responses of their terrestrial ecosystem modules to future  
52 changes in atmospheric CO<sub>2</sub>. This uncertainty arises primarily because of the differences in the  
53 strength of the CO<sub>2</sub> fertilization effect on the land carbon cycle components (Arora et al., 2013;  
54 Cramer et al., 2001; Friend et al., 2013) but also because of differences in the response of  
55 vegetation. Models differ in how the spatial distribution of vegetation, and its composition,  
56 changes in response to changing climate and increasing CO<sub>2</sub> (Cramer et al., 2001). These  
57 differences are also resolution dependent. For example, models with coarse grid resolutions  
58 cannot explicitly resolve climatic niches, which in turn potentially contributes to biases in  
59 simulated vegetation distribution (Melton and Arora, 2016; Shrestha et al., 2016).

60

61 Vegetation responds to changes in climate and atmospheric CO<sub>2</sub> concentration by changing its  
62 structural attributes including leaf area index (LAI), rooting depth, vegetation height, and canopy  
63 mass, as well as its areal extent. Structural vegetation changes generally occur over seasonal to  
64 decadal time scales (Kramer and Kozlowski, 1979), while the slower areal extent changes  
65 typically occur on decadal to centennial time scales (Ritchie and Macdonald, 1986). The  
66 dynamic behavior of vegetation affects weather and climate due to its strong control over  
67 biophysical processes. At hourly to daily timescales, vegetation affects the exchange of water  
68 and energy between the land surface and the atmosphere primarily through the control of leaf  
69 stomata. At longer ~~seasonal, annual and decadal~~ timescales, vegetation affects components of  
70 energy and water balance through its structure (LAI, rooting depth, etc.) and its areal extent and  
71 thereby land surface albedo. Conversely, dynamics of vegetation is directly influenced by  
72 climate and the competitive ability of the plants. In this way vegetation responds to climate by  
73 changing its structure and areal extent depending on the colonization ability of plants. These

Deleted: timescales from

Deleted: to

76 climate-vegetation interactions have been well documented (e.g. Gobron et al., 2010; Wang et  
77 al., 2011).

78  
79 Natural vegetation is typically characterized in dynamic global vegetation models (DGVMs)  
80 based on a limited number of PFTs (Sitch et al., 2003) because it is impossible to represent  
81 thousands of species in a model. Species characterized by similar attributes, mainly based on  
82 their form and interactions with the environment (Box, 1996), are grouped together as a single  
83 PFT. For example, tree species with similar leaf form such as fir (*Abies*), spruce (*Picea*) and pine  
84 (*Pinus*) are classified as needleleaf evergreen trees. The geographical distribution of the PFTs in  
85 DGVMs is determined by their ability to grow and increase their areal extent given certain  
86 climate and soil conditions and their competitive ability.

87  
88 One way of representing competition between PFTs in DGVMs is through the use of the Lotka-  
89 Volterra (LV) equations. While originally developed for predator-prey competition, the LV  
90 equations have been used in a number of DGVMs (Arora and Boer, 2006; Brentnall et al., 2005;  
91 Cox, 2001; Zhang et al., 2015). The use of the classical form of the LV equations for modelling  
92 competition between PFTs, however, leads to an amplified expression of dominance in that the  
93 dominant PFT ends up occupying a disproportionately large fraction of a grid cell leading to  
94 little co-existence between PFTs. Arora and Boer (2006) proposed changes to the classical  
95 implementation of the LV equations for modelling competition between PFTs to reduce this  
96 amplified expression of dominance. Their approach, which has been implemented in the CLASS-  
97 CTEM modelling framework and which allows improved co-existence of PFTs compared to the  
98 classical LV equations, has been shown to simulate vegetation distribution reasonably well at the  
99 global (Melton and Arora, 2016) as well as point (Shrestha et al., 2016) scales. Both these  
100 studies used climate averaged over  $\sim 3.75^\circ$  spatial resolution. The CLASS-CTEM framework  
101 consists of the Canadian Land Surface Scheme (CLASS) coupled to the Canadian Terrestrial  
102 Ecosystem Model (CTEM) which is a dynamic vegetation model.

103  
104 In this paper, we evaluate the competition module of the CLASS-CTEM modelling framework at  
105 the regional scale over the North American domain at  $1^\circ$  spatial resolution. This resolution is  
106 much finer than the  $3.75^\circ$  resolution used in the Melton and Arora (2016) study and therefore in

Deleted:

108 principle should allow a more realistic simulation of geographical distribution of PFTs as climate  
109 niches are resolved.

110  
111 The rest of this paper is organized as follows: Section 2 describes the CLASS-CTEM modelling  
112 framework, details of the observation-based data and the experimental setup. Results are  
113 presented in section 3 and a discussion follows in section 4. Finally, a summary and conclusions  
114 are provided in section 5.

115

## 116 | **2 Model, data and methods**

Deleted:

117

### 118 | **2.1 CLASS-CTEM model**

119

120 The results presented here are obtained by coupling version 2.0 of CTEM (Melton and Arora,  
121 2016), which dynamically simulates fractional coverage of its PFTs, to version 3.6 of CLASS  
122 (Verseghy et al., 1993). CTEM simulates terrestrial processes for seven non-crop and two crop  
123 PFTs (Table 1) and prognostically tracks carbon in three living vegetation components (leaves,  
124 stems and roots) and two dead carbon pools (litter and soil). The terrestrial ecosystem processes  
125 simulated in this study include photosynthesis, autotrophic respiration, heterotrophic respiration,  
126 dynamic leaf phenology, allocation of carbon from leaves to stem and root components, fire,  
127 land use change, and competition between PFTs which dynamically determines the fractional  
128 coverage of each PFT. The amount of carbon in the leaf, stem and root components is used to  
129 estimate structural attributes of vegetation. LAI is calculated from leaf biomass using PFT-  
130 dependent specific leaf area (SLA) which determines area of leaves that can be constructed per  
131 kg C of leaf biomass (Arora and Boer, 2005); vegetation height is calculated based on stem  
132 biomass for tree PFTs and LAI for grass PFTs; and rooting depth is calculated based on root  
133 biomass (Arora and Boer, 2003). CTEM operates at a time step of one day except for  
134 photosynthesis and leaf respiration which are calculated every 30 minutes for consistency with  
135 CLASS' energy and water balance calculations which require stomatal resistance calculated by  
136 the photosynthesis module of CTEM.

137

139 CLASS simulates the energy and water balance components at the land surface and operates at a  
 140 30 minutes time step. Liquid and frozen soil moisture and soil temperature are evaluated for  
 141 three soil layers (with maximum thicknesses of 0.1, 0.25 and 3.75 m). The actual thicknesses of  
 142 these permeable soil layers are determined by the depth to bedrock, which is specified on the  
 143 basis of the global data set of Zobler (1986). CLASS distinguishes four PFTs (needleleaf trees,  
 144 broadleaf trees, crops and grasses) which map directly to the nine PFTs represented in CTEM as  
 145 shown in Table 1. Needleleaf trees in CTEM are divided into deciduous and evergreen types,  
 146 broadleaf trees are divided into cold and drought deciduous and evergreen types, and crops and  
 147 grasses are divided into C<sub>3</sub> and C<sub>4</sub> types based on their photosynthetic pathways. In coupled  
 148 mode, CLASS uses the dynamically simulated vegetation attributes (including LAI, vegetation  
 149 height, canopy mass and rooting depth) and stomatal resistance calculated by CTEM, and CTEM  
 150 uses the soil moisture, soil temperature and net shortwave radiation calculated by CLASS. The  
 151 coupling frequency between CLASS and CTEM is one day.

152

### 153 2.1.1 Competition parameterization

154

155 Competition between PFTs in CTEM is parameterized following Arora and Boer (2006) who  
 156 presented a modified version of the LV equations. The approach is described in detail by Melton  
 157 and Arora (2016) and briefly summarized here. Consider, for simplicity, two PFTs that exist in a  
 158 grid cell with fractional coverages  $f_1$  and  $f_2$ . Let PFT 1 represent a tree PFT and PFT 2 represent  
 159 a grass PFT. The bare fraction of grid cell not covered by any vegetation is represented by  $f_B$ . As  
 160 a result,  $f_1 + f_2 + f_B = 1$ . The rate of change of fractional coverages of the two PFTs and bare  
 161 fraction, for this example, are given by,

162

$$163 \left| \frac{df_1}{dt} = c_1 f_1^\beta (1 - f_1) - m_1 f_1 \right. \quad (1)$$

164

$$165 \left| \frac{df_2}{dt} = c_2 f_2^\beta (1 - f_1 - f_2) - c_1 f_1^\beta f_2 - m_2 f_2 \right. \quad (2)$$

166

$$167 \left| \frac{df_B}{dt} = -c_1 f_1^\beta f_B - c_2 f_2^\beta f_B + m_1 f_1 + m_2 f_2 \right. \quad (3)$$

168

169 where  $c_1$ ,  $c_2$  and  $m_1$ ,  $m_2$  are the colonization and mortality rates for PFT 1 and PFT 2,  
170 respectively. Colonization and mortality rates cannot be negative. Equations (1) and (2) show  
171 that PFT 1 can invade the fraction covered by PFT 2 and the bare fraction; and that PFT 2 can  
172 only invade the bare fraction. PFT 2 is not allowed to invade the fraction covered by PFT 1  
173 because it is ranked lower than PFT 1. In CTEM, the superiority or ranking of the seven natural  
174 non-crop PFTs is based on the tree-grass distinction and their colonization rates. Trees are always  
175 considered to be superior than grasses because of their ability to shade them (Siemann and  
176 Rogers, 2003). Within the tree and grass PFTs the dominance is determined dynamically based  
177 on the colonization rate. The exponent  $\beta$  ( $0 \leq \beta \leq 1$ ), an empirical parameter, controls the  
178 behaviour of the LV equations. For  $\beta = 1$ , the equations represent the classical form of the LV  
179 equations. The equilibrium fractional coverages for PFT 1 and 2 and bare fraction for this  
180 classical form of the LV equations, denoted by  $\tilde{f}_1, \tilde{f}_2$  and are given by,

$$\tilde{f}_1 = \max\left\{\left(\frac{c_1 - m_1}{c_1}\right), 0\right\} \quad (4)$$

$$\tilde{f}_2 = \max\left\{\left(\frac{(c_2 - m_2) - \left(1 + \frac{c_2}{c_1}\right)(c_1 - m_1)}{c_2}\right), 0\right\} \quad (5)$$

$$\tilde{f}_B = \frac{(m_1 \tilde{f}_1 + m_2 \tilde{f}_2)}{(c_1 \tilde{f}_1 + c_2 \tilde{f}_2)} \quad (6)$$

188 In equations (1) and (2), if the fractional coverages of PFT 1 and PFT 2 are initially zero then the  
189 PFTs cannot expand for  $\beta = 1$ , implying that a minimum seeding fraction is always required.  
190 Furthermore, in equation (5) as long as  $(c_1 - m_1)$  is greater than  $(c_2 - m_2)$  then the equilibrium  
191 solution for  $\tilde{f}_2$  will always be zero and PFT 2 will not be able to coexist with PFT 1. These  
192 features of the classical form of the LV equations are avoided when  $\beta = 0$ , following Arora and  
193 Boer (2006). The equilibrium fractional coverages for PFT 1 and 2 and bare fraction for the case  
194 with  $\beta = 0$  are given by,

$$\tilde{f}_1 = \left(\frac{c_1}{c_1 + m_1}\right) \quad (7)$$

$$\tilde{f}_2 = \frac{c_2(1 - \tilde{f}_1)}{(c_1 + c_2 + m_2)} = \left(\frac{c_2 m_1}{(c_1 + m_1)(c_1 + c_2 + m_2)}\right) \quad (8)$$

198 
$$\tilde{f}_B = \frac{(m_1 \tilde{f}_1 + m_2 \tilde{f}_2)}{(c_1 + c_2)} \quad (9)$$

199 Unlike the classical version of the LV equations, the modified version of the equations with  $\beta = 0$   
 200 does not require a minimum seeding fraction, and PFTs are able to increase their areal extent as  
 201 long as the climate is favorable and  $c_i$  is positive. Also, as long as  $m_1 > 0$  and  $c_2 > 0$  then PFT 2  
 202 is able to coexist at equilibrium with PFT 1. Other values of  $\beta$  between 0 and 1 give the dominant  
 203 PFT varying levels of access to sub-dominant PFTs but coexistence is most possible in the case  
 204 with  $\beta = 0$ .

205  
 206 The calculations of colonization and mortality rates are described in detail in Melton and Arora  
 207 (2016). Briefly, the colonization rate depends on the net primary productivity of a PFT. The  
 208 better a PFT performs for given climatic and soil conditions; the higher is its colonization rate.  
 209 The mortality rate represents the combined effect of four different processes: intrinsic or age-  
 210 related mortality, growth or stress mortality, mortality due to disturbance, and mortality due to  
 211 adverse climate which ensures that tree PFTs do not venture outside their bioclimatic zones.

212  
 213 **2.2 Forcing data**

214  
 215 The Climate Research Unit – National Centre for Environmental Prediction (CRU-NCEP)  
 216 reanalysis dataset (Viovy, 2012), is used to drive the model. The meteorological variables  
 217 (surface temperature, pressure, precipitation, wind, specific humidity, and incident short-wave  
 218 and long-wave radiation fluxes) are available at a spatial resolution of  $0.5^\circ \times 0.5^\circ$  and at a six  
 219 hourly time interval for the period 1901-2010. These data are interpolated to  $1^\circ$  resolution  
 220 spatially, and disaggregated to half-hourly time resolution, a standard CLASS-CTEM model  
 221 integration time step. Temperature, pressure, wind, specific humidity, and long-wave radiation  
 222 are linearly interpolated in time while short-wave radiation is assumed to change with the solar  
 223 zenith angle with maximum radiation occurring at solar noon. Following Arora (1997), the six-  
 224 hourly precipitation amount ( $P$ , mm/6-hour) is used to estimate the number of wet half-hours  
 225 ( $w_h$ ) in a given six-hour period for  $P > 0$  as

226  
 227 
$$w_h = \text{integer}(\max[1, \min(12, 2.6 \log(6.93 P))]). \quad (10)$$

Deleted:

229

230 The total precipitation amount is then distributed randomly but conservatively over these wet  
231 half-hours. For instance, if seven out of 12 half hours intervals are calculated to be wet using  
232 equation (10) then seven random numbers varying between 0 and 1 are generated and the six-  
233 hourly precipitation amount is divided into seven parts in proportion to their respective random  
234 numbers

235

236 Figure 1 shows the spatial distribution of mean annual precipitation and surface temperature over  
237 the North American domain considered in this study. Mean annual precipitation values range  
238 from less than 200 mm in the arid south-west United States and the high Arctic to more than  
239 1500 mm on the Pacific coast. Mean annual temperature varies from around 24° C near the  
240 southern limit of the domain in Mexico to less than -20° C in the Arctic tundra.

241

## 242 **2.3 Observation-based data**

### 243 **2.3.1 Fractional coverage of PFTs**

244

245 Observation-based estimates of fractional coverages of PFTs are based on a modified version of  
246 the Wang et al. (2006) data set (hereafter WANG06) and the Moderate Resolution Imaging  
247 Spectroradiometer land cover product (Friedl et al., 2013) (hereafter MODIS). These data are  
248 used to evaluate the model results.

249

250 The WANG06 data set was developed for use by CTEM in simulations in which competition is  
251 turned off and prescribed fractional coverage of PFTs is used. It combines observation- and  
252 model-based data to estimate the annual change in fractional coverage of CTEM's nine PFTs  
253 from 1850 to 2000. The Global Land Cover for the year 2000 (GLC2000), which is considered  
254 as a base year for environmental assessment, divides the global land cover in 22 types is  
255 available at 1 km resolution. WANG06 (their Table 2) mapped the GLC2000 data to CTEM's  
256 nine PFTs aggregated to 0.5° resolution. The GLC2000 data were then extrapolated back to 1850  
257 by adjusting the changes in crop area based on the then available Ramankutty and Foley (1999)  
258 | [crop](#) data set. Here, we use a modified version of the WANG06 data set which is based on the

259 HYDE v.3.1 crop data set (Hurt et al., 2011) and generate an estimate of fractional coverage of  
260 CTEM PFTs for the period 1850-2012.

261

262 The MODIS data set is based on the International Geosphere-Biosphere Programme (IGBP)  
263 global vegetation data and University of Maryland's Science Data Set classification schemes at  
264 0.25° spatial resolution. The data are derived from NASA HDF-EOS MODIS/Terra land cover  
265 type. The data set is for the period 2001 to 2014 and contains 17 land cover types which we map  
266 to CTEM's nine PFTs following the logic used in Wang et al. (2006) as shown in Table 2. The  
267 fractional coverage of each of the nine CTEM PFT is first obtained at 0.25 degree resolution for  
268 each year using the mapping scheme described in Table 2. These fractional coverages are then  
269 re-gridded to the 1° spatial resolution for individual years. Finally, the data are averaged over the  
270 period 2001-2014 to evaluate model results. MODIS data are known to exhibit substantial  
271 interannual variability. Broxton et al. (2014), for instance, report that globally 40% of land pixels  
272 show land cover change one or more times during 2001–2010 period. This does not necessarily  
273 indicate changes in land cover but rather these differences are due to low accuracy in  
274 categorizing the remotely sensed vegetation into one of the 17 MODIS land cover types, as  
275 Broxton et al. (2014) note. This low accuracy is itself attributed to the fact that many landscapes  
276 include mixtures of vegetation classes. Our re-gridding of fractional coverages to 1° spatial  
277 resolution and averaging over the 2001-2014 time period to obtain climatology of land cover  
278 alleviates some of the uncertainty since the effect of inaccurately classified land cover categories  
279 is reduced due to both spatial and temporal averaging.

280

281 The separation of the broadleaf deciduous PFT into its drought and cold deciduous components  
282 is performed via the approach used by WANG06. They assumed that below 24 °N deciduousness  
283 is caused by soil moisture limitation and hence all broadleaf deciduous trees below this latitude  
284 are drought deciduous, and above 34 °N deciduousness is caused by low temperatures and so all  
285 broadleaf deciduous trees above this latitude are cold deciduous. Between 24 °N and 34 °N,  
286 following WANG06 we assume a linear transition from drought deciduous to cold deciduous  
287 trees. Finally, the separation of grasses into their C<sub>3</sub> and C<sub>4</sub> components is based on the  
288 geographical distributions of the C<sub>3</sub> and C<sub>4</sub> fractions in the WANG06 data set.

289 **2.3.2 Gross primary productivity and LAI**

290

291 Observation-based estimates of gross primary productivity (GPP) are based on Beer et al. (2010).  
292 These data are based on the ecosystem level GPP obtained using eddy covariance measurements  
293 from more than 250 stations across the globe. Beer et al. (2010) extrapolated GPP values based  
294 on these eddy covariance flux data ~~to the global scale using diagnostic models for the period~~  
295 1982 – 2008, and the average over this time period is used to evaluate the model results. LAI  
296 data used for validation are the same as those used by Anav et al. (2013) and are based on Zhu et  
297 al. (2013) who use normalized difference vegetation index (NDVI) data from the Advanced Very  
298 High Resolution Radiometer (AVHRR) satellite to calculate average LAI for the period 1981 –  
299 2010.

Deleted: of GPP

300

301 **2.4 Experimental setup**

302

303 **2.4.1 Equilibrium pre-industrial simulation**

304

305 The equilibrium pre-industrial simulation was initialized from zero biomass and zero fractional  
306 coverage for all non-crop PFTs. The fractions of C<sub>3</sub> and C<sub>4</sub> crop PFTs in each grid cell are  
307 specified corresponding to year 1850 based on the HYDE 3.1 dataset. The model was then run  
308 for 600 years driven by 1901-1925 CRU-NCEP climate data cycled repeatedly. These data do  
309 not show any warming trend (Wen et al., 2011) as opposed to the later part of the 20<sup>th</sup> century.  
310 Atmospheric CO<sub>2</sub> concentration was set to 285 ppm corresponding to the pre-industrial 1850  
311 level. This pre-industrial equilibrium simulation yields initial conditions including fractional  
312 coverages of PFTs and carbon in all the live and dead pools for the transient 1850-2010  
313 simulation. The 600 years simulation is sufficient for fractional vegetation cover and carbon  
314 pools to reach equilibrium.

315 **2.4.2 Transient historical simulation**

316

317 The transient historical simulation is performed for the period 1851-2010 and its carbon pools  
318 and fractional coverage of non-crop PFTs are initialized from the equilibrium pre-industrial  
319 simulation as mentioned above. The years 1851 to 1900 of this historical simulation are driven

321 with CRU-NCEP climate data corresponding to the period 1901-1925, cycled twice. For the  
322 period 1901-2010 the climate data corresponding to each year are used. Time varying  
323 concentrations of atmospheric CO<sub>2</sub> are supplied for the period 1851-2010 based on the values  
324 used in the fifth Coupled Modelling Intercomparison Project (CMIP5,  
325 <http://tntcat.iiasa.ac.at/RcpDb/>) which are extended past 2005 to 2010 based on data from the  
326 National Oceanic and Atmospheric Administration  
327 ([ftp://aftp.cmdl.noaa.gov/products/trends/co2/co2\\_annmean\\_gl.txt](ftp://aftp.cmdl.noaa.gov/products/trends/co2/co2_annmean_gl.txt)). The annual time-varying  
328 fractional coverages of C<sub>3</sub> and C<sub>4</sub> crop PFTs in each grid cell are based on the HYDE 3.1 dataset.  
329 The crop fractions in a grid cell are not available for colonization and neither are they subject to  
330 disturbance by fire. Competition between PFTs occurs over the remaining non-crop fraction of a  
331 grid cell. As total crop fraction in a grid cell changes over time (based on the HYDE 3.1 dataset)  
332 the fractional area available for competition also changes.

333

334 The simulated results are evaluated against their observation-based counterparts using averaged  
335 values over the last 30 years of the simulation corresponding to the period 1981-2010. This is the  
336 same and/or very close to the time period for modified WANG06 land cover data set (1981-  
337 2010), Beer et al. (2010) GPP (1982-2008), and Zhu et al. (2013) LAI (1981-2010). The only  
338 exception is the MODIS-based land cover data which are available for the 2001-2014 period.

### 339 **3 Results**

340

#### 341 **3.1 Continental scale values of PFT coverage**

342

343 Figures 2a compares the simulated vegetation areas summed over our North American domain  
344 with the WANG06 and MODIS observation-based estimates. In the absence of another measure  
345 of uncertainty, we use the range between these two observation-based estimates and assess if  
346 simulated areal coverage of a given land cover type lies within or outside this range. The  
347 simulated total vegetated area over North America ( $14.8 \times 10^6$  km<sup>2</sup>) is very similar to the  
348 modified WANG06 ( $14.4 \times 10^6$  km<sup>2</sup>) and MODIS derived ( $14.2 \times 10^6$  km<sup>2</sup>). At the most basic  
349 tree-grass-bare ground level, the simulated areas are closer to the MODIS-based estimates, than  
350 to the estimate based on the modified WANG06 data. The simulated area covered by tree PFTs

351  $(7.8 \times 10^6 \text{ km}^2)$  is 6% lower than the MODIS derived estimate  $(8.2 \times 10^6 \text{ km}^2)$  and 21% lower  
352 than WANG06  $(9.7 \times 10^6 \text{ km}^2)$ . The simulated grass coverage  $(4.7 \times 10^6 \text{ km}^2)$  is 35% higher  
353 than the MODIS derived estimate  $(3.5 \times 10^6 \text{ km}^2)$ . Both simulated and MODIS-based estimates  
354 of area covered by grass PFTs are, however, substantially higher than the WANG06  $(2.4 \times 10^6$   
355  $\text{km}^2)$  estimate. Averaged over the North American region, the simulated partitioning of land area  
356 (excluding cropland area) covered by trees, grasses and bare ground (45%, 27%, 28%) is much  
357 closer to the MODIS based data (48%, 20% and 32%) than to the modified WANG06 based data  
358 (56%, 14%, 30%).

359

360 Figure 2b shows a comparison of simulated areas of individual PFTs with observation-based  
361 estimates. This is a more stringent test of the performance of the competition module of CTEM.  
362 The observation-based estimates of areas of all individual PFTs are available for the modified  
363 WANG06 dataset. The MODIS based estimates were derived based on the mapping of MODIS'  
364 17 land cover types to CTEM PFTs as shown in Table 2, which itself is mostly based on  
365 WANG06. In Figure 2b, the observation-based estimates show that needleleaf evergreen (NDL  
366 EVG) and broadleaf cold deciduous (BDL DCD CLD) are the dominant tree PFTs across North  
367 America ~~and the model is able to reproduce this aspect~~. The simulated total area of the NDL  
368 EVG tree PFT  $(3.9 \times 10^6 \text{ km}^2)$  is 28% less than WANG06  $(5.3 \times 10^6 \text{ km}^2)$  and 15% less than the  
369 MODIS based estimate  $(4.7 \times 10^6 \text{ km}^2)$ . The simulated total area of BDL DCD CLD tree PFT  $(3$   
370  $\times 10^6 \text{ km}^2)$  is 13% less than WANG06  $(3.4 \times 10^6 \text{ km}^2)$  and 3% greater than MODIS based  $(2.9 \times$   
371  $10^6 \text{ km}^2)$  estimate. Overall, the model is able to capture the areas covered by individual PFTs  
372 reasonably well. However, differences remain between observations-based and simulated  
373 estimates especially the larger simulated area for  $C_3$  grasses than both observation-based  
374 estimates. Reasons for these differences include limitations in the model but also the manner in  
375 which remotely-sensed vegetation is categorized into broad-scale vegetation types and then  
376 mapped onto CTEM's nine PFTs, as discussed later.

377

378 In both Figures 2a and 2b although simulated areal coverages at the basic tree-grass-bare ground  
379 level and for individual PFTs (except for  $C_3$  grasses) are comparable to observation-based

Deleted: which is also shown by the model

382 estimates they are outside the range defined by difference of the WANG06 and MODIS based  
383 estimates.

384  
385 Figure 2c shows the time series of simulated areas summed over the domain covered by tree and  
386 grass PFTs, the total vegetated area and the remaining bare ground. The specified area covered  
387 by crop PFTs, based on the HYDE 3.1 data set, is also shown and first increases over the  
388 historical period and then stabilizes and in fact somewhat decreases in association with cropland  
389 abandonment over the north-eastern United States. The increase in the crop area results in a  
390 decrease in the area covered by tree and grass PFTs up until the time when the crop area  
391 stabilizes around 1970. In the model, this causes land use change emissions associated with  
392 deforestation. After this time, as vegetation productivity responds to increasing atmospheric CO<sub>2</sub>  
393 concentration, the area covered by tree PFTs increases somewhat and colonizes available bare  
394 areas and those covered by grass PFTs. This leads to a small reduction in the area covered by  
395 grass PFTs as well as bare ground and the associated increase in the total vegetated area.

396

## 397 **3.2 Geographical distribution of PFTs**

398

### 399 **3.2.1 Total vegetated and bare ground fractions**

400

401 Figures 3 and 4 compare the geographical distribution of simulated total vegetated and bare  
402 fractions across North America with the two observation-based estimates derived from the  
403 modified WANG06 and MODIS data sets. The two observation-based estimates are also  
404 compared amongst themselves. The metrics used are averaged root mean square difference  
405 (RMSD) and spatial correlations ( $R^2$ ).

406

407 The observation-based geographical distribution of vegetated fraction in Figure 3 (middle  
408 column) shows densely vegetated land over the eastern part of the continent and less vegetation  
409 coverage over colder regions in the North and drier regions in the south-central and south-west  
410 United States. These broad scale patterns are consistent with the precipitation and temperature  
411 climatologies of the region (Figure 1). The model reasonably reproduces the observed vegetation  
412 distribution (left panel) with some obvious limitations. Simulated vegetation cover is

413 underestimated across the arid south-west United States, Great Plains and part of the Canadian  
414 Prairies (right panel) due to lower simulated fractional coverage of tree and grass PFTs over  
415 these regions (shown in the next section). The model overestimates vegetation coverage in  
416 Northern Canada because of higher simulated grass cover in the Arctic as discussed below in  
417 more detail. The spatial correlation and RMSD when comparing simulated vegetated fraction to  
418 both observation-based estimates are 0.79 and around 18%, respectively. The spatial correlation  
419 and RMSD between the two observation-based estimates themselves are 0.86 and around 14%,  
420 respectively.

Deleted: which are

421  
422 The simulated and observation-based bare ground fractions across North America are compared  
423 in Figure 4. The observation-based estimates show that bare ground fraction is higher in Arctic  
424 Canada and Alaska where, of course, cold temperatures limit vegetation growth and in the south-  
425 west United States, Great Plains and the Prairies where low rainfall limits vegetation growth  
426 (Figure 1). The biases in simulated bare ground fraction mirror those in the simulated vegetated  
427 fraction but in an opposite manner. The model underestimates bare ground fraction across Arctic  
428 Canada due to higher simulated grass cover as discussed in the next section. The model  
429 overestimates the bare ground fraction generally across the arid and semi-arid south-west United  
430 States, Great Plains and the Prairies. The spatial correlations and RMSDs when comparing  
431 simulated bare ground fraction to both observation-based estimates, and when comparing the two  
432 observation-based data sets amongst themselves, are the same as those for the total vegetation  
433 fraction in Figure 3.

Deleted:

Deleted: around 0.46 and around 18%, respectively. The spatial correlation and RMSD between the two observation-based estimates themselves are 0.68 and around 14%, respectively.

### 435 3.2.2 Tree and grass cover

436  
437 Figure 5 compares the simulated tree cover with the two observation-based estimates. The model  
438 reasonably reproduces the broad scale patterns including the Canadian boreal forest and the  
439 temperate forests across the southeastern United States. However, the model simulates lower tree  
440 cover across the western part of the continent compared to both observation-based estimates  
441 particularly over the southwestern United States which is characterized by arid climate (Figure  
442 1). The observation-based estimates do not particularly well agree over this region either. The  
443 MODIS derived estimate suggests around 25% tree cover in the southwestern United States

451 while the WANG06 derived estimate suggests a tree cover of around 60% over a large area in  
452 the region. The spatial correlation and RMSD when comparing simulated tree cover to both  
453 observation-based estimates are around 0.68 and around 17%, respectively. The spatial  
454 correlation and RMSD between the two observation-based estimates themselves are 0.75 and  
455 around 15%, respectively. Possible reasons for differences between simulated and observation-  
456 based estimates are discussed in detail in the discussion section and include the fact that the  
457 CLASS-CTEM framework does not currently represent shrubs and there are limitations in the  
458 observation-based data sets themselves. Shrubs are more prevalent in arid and semi-arid regions  
459 where they are better suited to grow compared to both trees and grasses.

460  
461 Figure 6 compares the geographical distribution of the simulated grass cover with the two  
462 observation-based estimates. The broad geographical distribution of simulated grass cover  
463 compares well with the two observation-based estimates with the notable exception of the Arctic  
464 region including Alaska and northern Canada, where the model overestimates grass cover. This  
465 overestimation of grass cover in the Arctic region is also the reason for the overestimation of  
466 total vegetation fraction and the underestimation of bare fraction that was seen earlier in Figures  
467 3 and 4 respectively.

468  
469 As shown in Figure 6, the spatial correlation and RMSD when comparing simulated grass cover  
470 to both observation-based estimates lie between 0.33 and 0.38 and between around 15-17%,  
471 respectively. The spatial correlation and RMSD between the two observation-based estimates  
472 themselves are 0.54 and around 9%, respectively. The two observation-based estimates disagree  
473 most markedly over the western half of the United States where the MODIS derived estimates of  
474 grass cover are higher.

475  
476 **3.2.3 Needleleaf evergreen and broadleaf cold deciduous trees**

477  
478 Figures 7a and 7b compare the geographical distribution of NDL EVG and BDL DCD CLD  
479 trees, respectively, with their observation-based estimates. These two are the primary tree PFTs  
480 which exist in the North American domain considered here.

481

482 In Figure 7a, the overall simulated coverage of NDL EVG trees is lower than both observation-  
483 based estimates as was also seen in Figure 2b. The simulated values are primarily lower in  
484 western Canada and over a large area in the western United States according to estimates based  
485 on the modified WANG06 data set. This is also the case along the wide swath of the Canadian  
486 boreal forest. The model overestimates the coverage of NDL EVG trees in the eastern United  
487 States. The spatial correlation and RMSD when comparing simulated coverage of NDL EVG  
488 trees to both observation-based estimates lie between 0.36 and 0.40 and between around 16-17%,  
489 respectively. The spatial correlation and RMSD between the two observation-based estimates  
490 themselves are 0.52 and around 16%, respectively.

491  
492 The geographical distribution of BDL DCD CLD trees is compared with its observation-based  
493 estimates in Figure 7b. Although the simulated domain summed area of BDL DCD CLD trees ( $3$   
494  $\times 10^6$  km<sup>2</sup>) is comparable to estimates based on the modified WANG06 ( $3.4 \times 10^6$  km<sup>2</sup>) and  
495 MODIS ( $2.9 \times 10^6$  km<sup>2</sup>) data sets, there are two primary limitations in its simulated geographical  
496 distribution. First, the simulated values are generally overestimated in Canadian boreal forests  
497 and underestimated in the eastern United States. Second, the model simulates near zero coverage  
498 in the arid south-western United States. The spatial correlation and RMSD when comparing  
499 simulated coverage of BDL DCD CLD trees to both observation-based estimates are around 0.3  
500 and around 12%, respectively. The spatial correlation and RMSD between the two observation-  
501 based estimates themselves are 0.60 and around 8%, respectively.

502

### 503 3.2.4 C<sub>3</sub> and C<sub>4</sub> grasses

504

505 Figures 8a and 8b compare the simulated geographical distribution of C<sub>3</sub> and C<sub>4</sub> grasses with  
506 observation-based estimates.

507

508 In Figure 8a, the most obvious limitation of the model is its excessive simulated grass coverage  
509 in Alaska and in Arctic Canada. Other than this, the model reproduces the broad geographical  
510 distribution of C<sub>3</sub> grasses including the Great Plains of United States and the Canadian Prairies,  
511 where a large extent of grasslands is observed. The overestimated grass coverage at high  
512 latitudes leads to a total simulated C<sub>3</sub> grass area ( $4.4 \times 10^6$  km<sup>2</sup>) that is higher than estimates

513 based on the modified WANG06 ( $1.9 \times 10^6 \text{ km}^2$ ) and MODIS ( $2.8 \times 10^6 \text{ km}^2$ ) data sets. The  
514 spatial correlation and RMSD when comparing simulated coverage of  $C_3$  grasses to both  
515 observation-based estimates lie between 0.34-0.38 and between around 15-17%, respectively.  
516 The spatial correlation and RMSD between the two observation-based estimates themselves are  
517 0.54 and around 12%, respectively.

518  
519 Figure 8b shows the distribution of  $C_4$  grasses which mostly occur in the tropics and do not  
520 occupy large areas in North America (as was also seen in Figure 2b). The modelled geographical  
521 distribution of  $C_4$  grasses is larger than observation-based estimates but the absolute fractions  
522 remain small so that the simulated area covered over the whole domain ( $0.35 \times 10^6 \text{ km}^2$ ) is  
523 actually smaller than estimates based on the modified WANG06 ( $0.45 \times 10^6 \text{ km}^2$ ) and MODIS  
524 ( $0.7 \times 10^6 \text{ km}^2$ ) data sets. The spatial correlation and RMSD when comparing simulated  
525 coverage of  $C_4$  grasses to both observation-based estimates lie between 0.12-0.16 and between  
526 around 3-5%, respectively. The spatial correlation and RMSD between the two observation-  
527 based estimates themselves are 0.62 and around 5%, respectively.

528  
529 We do not compare the spatial distribution of broadleaf evergreen (BDL EVG) and broadleaf  
530 drought deciduous (BDL DCD DRY) trees with the two observation-based estimates for three  
531 reasons: 1) the geographical distribution of these PFTs is limited to a small total area in our  
532 domain, 2) the geographical distribution of the BDL EVG tree PFT based on observations cannot  
533 be directly compared to simulated values because, when mapping land cover types to CTEM  
534 PFTs in WANG06, evergreen shrubs (which exist much farther north than 30 °N) are assigned to  
535 the BDL EVG tree PFT, and 3) the geographical distribution of the BDL DCD DRY tree PFT in  
536 the observation-based data sets is based on the arbitrary latitudinal thresholds of 24 °N and 34 °N  
537 as mentioned earlier.

### 538 539 3.3 LAI and GPP

540  
541 Figure 9 compares the geographical distribution of simulated LAI and GPP with observation-  
542 based estimates for the present day. In Figure 9a, the simulated geographical distribution of LAI  
543 compares well with the observation-based estimates. The spatial correlation and RMSD between

Deleted: ¶  
Broadleaf evergreen and drought deciduous trees¶

The least prevalent PFTs in the North American domain considered here are broadleaf evergreen (BDL EVG) and broadleaf drought deciduous (BDL DCD DRY) trees. As they are represented in the model these are primarily tropical PFTs and hence generally do not exist above around 30 °N (see Figure 9), according to the bioclimatic limits used in the model for tree PFTs (Melton and Arora, 2016). In our simulations, these PFTs therefore exist near the southern edge of the United States. We do not evaluate spatial correlation and RMSD for these PFTs compared to the

Deleted: the

Deleted: 10

Deleted: 10a

565 simulated and observation-based estimates are 0.74 and 0.81  $\text{m}^2/\text{m}^2$ , respectively. The domain  
566 averaged simulated LAI of 2.5  $\text{m}^2/\text{m}^2$  is higher than the observation-based estimate of 2.1  $\text{m}^2/\text{m}^2$ .  
567 The model captures the broad geographical patterns with higher LAI over the boreal forest  
568 region in Canada and also in the eastern United States similar to observations. However, some  
569 differences remain particularly over the drier southwest United States where the model simulates  
570 bare ground with negligible LAI but observations suggest a small LAI of around 1  $\text{m}^2/\text{m}^2$ . In  
571 contrast, the model slightly overestimates LAI over northern and Arctic Canada where it  
572 simulates a higher fractional coverage of  $\text{C}_3$  grasses, as seen earlier.

573

574 Consistent with the geographical distribution of LAI, the simulated GPP is overestimated in the  
575 eastern United States and the Canadian boreal forest (Figure 9b). The broad geographical  
576 distribution of GPP, similar to LAI, is consistent with the observation-based estimates. The  
577 spatial correlation and RMSD between simulated and observation-based estimates are 0.78 and  
578 225  $\text{gC}/\text{m}^2\cdot\text{year}$ , respectively. The domain averaged simulated GPP of 737  $\text{gC}/\text{m}^2\cdot\text{year}$  is higher  
579 than the observation-based estimate of 628  $\text{gC}/\text{m}^2\cdot\text{year}$ . As with LAI, the simulated GPP is lower  
580 than observations over the drier southwest region of the United States where the model simulates  
581 more bare ground than observation-based estimates, and the model overestimates GPP over the  
582 northern and Arctic Canada.

Deleted: 10b

583

584 Figure 10 shows the time series of annual domain averaged GPP, LAI, net primary productivity  
585 (NPP) and domain summed net biome productivity (NBP). The NBP term is essentially the net  
586 atmosphere-land  $\text{CO}_2$  flux which is the result of all terrestrial ecosystem processes including  
587 photosynthesis, autotrophic and heterotrophic respiration, fire and land use change. NBP values  
588 of zero indicate that the system is in equilibrium such that carbon gained by photosynthesis is  
589 equal to carbon lost by respiration and other processes. Simulated GPP, LAI and NPP all show  
590 an increase over the 20<sup>th</sup> century due to the increase in atmospheric  $\text{CO}_2$  concentration and the  
591 associated change in climate. The increase in  $\text{CO}_2$  drives the increase in GPP and subsequently in  
592 NPP and LAI through the  $\text{CO}_2$  fertilization effect. The net result of this gradually increasing NPP  
593 is that the terrestrial ecosystems become a sink of carbon and this is seen in the resulting positive  
594 values of NBP. The simulated sink over the North American domain for the periods 1990-2000  
595 and 2000-2010 is around 0.4 and 0.5  $\text{Pg C}/\text{year}$ , respectively. Crevoisier et al. (2010) compare

Deleted: 1

598 the carbon sink over the North American region from five studies (their Table 1) for time periods  
599 in the 1990s and 2000s. These reported sinks vary from  $0.81 \pm 0.72$  to  $1.26 \pm 0.23$  Pg C/year for the  
600 period 1992-1996, 0.58 Pg C/yr for the period 2001-2006 and Crevoisier et al. (2010) themselves  
601 estimate a value of  $0.51 \pm 0.41$  Pg C/yr for the period 2004-2006. The sinks simulated by  
602 CLASS-CTEM over the 1990s and 2000s are broadly consistent with these estimates.

603

### 604 **3.4 Added value of finer spatial resolution**

605

606 Figure 11 assesses the added value of running the model and performing competition between  
607 PFTs at the 1° spatial resolution used in this study compared to the 3.75° resolution used in  
608 Melton and Arora (2016) study which evaluated the performance of CLASS-CTEM's  
609 competition module at the global scale. For Figure 11, the Melton and Arora (2016) results were  
610 extracted for the North American domain used in this study and observation-based estimates of  
611 fractional coverage of tree, grass and total vegetation from the modified WANG06 land cover  
612 product were re-gridded to the 3.75° resolution. The resulting spatial correlations and RMSDs  
613 between the simulated and the WANG06 estimates for fractional coverage of tree, grass and total  
614 vegetation, at the two spatial resolutions, are summarized in Figure 11. When compared to the  
615 modified WANG06 data the RMSDs are somewhat lower (Figure 11a), and spatial correlations  
616 (Figure 11b) are slightly higher for model's implementation at 3.75° resolution, compared to  
617 model's implementation at 1° resolution. This indicates that the model's performance is slightly  
618 better at the coarser 3.75° resolution. Recall that competition between PFTs occurs over the non-  
619 crop fraction of each grid cell. For this reason, we do not perform this analysis for MODIS based  
620 land cover product because the crop areas that are specified in the model are exactly same as  
621 those in the modified WANG06 land cover product making comparison of simulated and  
622 observation-based fractional coverages of PFTs more consistent for the modified WANG06 land  
623 cover product.

624

## 625 **4 Discussion**

626

627 Competition between PFTs, that determines their fractional coverage, is one of the several  
628 processes that the CLASS-CTEM modelling framework simulates. Other than competition

Deleted: ¶

630 between PFTs, terrestrial ecosystem processes of photosynthesis, autotrophic and heterotrophic  
631 respiration, allocation of carbon from leaves to stem and root components, dynamic leaf  
632 phenology, fire, and land use change are also modelled. These aspects of the model have been  
633 evaluated at point (Arora, 2003; Arora and Boer, 2005; Melton et al., 2015), regional (Garnaud et  
634 al., 2015; Peng et al., 2014; Arora et al., 2016) and global (Arora and Boer, 2010; Melton and  
635 Arora, 2014; Melton and Arora, 2016) scales. A typical model evaluation exercise at the global  
636 scale compares model simulated geographical and latitudinal distribution of GPP, vegetation  
637 biomass, and soil carbon with their respective observation-based estimates such as those from  
638 Beer et al. (2010), Ruesch and Holly (2008) and Harmonized World Soil Database  
639 (FAO/IIASA/ISRIC/ISS-CAS/JRC, 2012). Model evaluation exercises help in identifying model  
640 limitations but also yield opportunities to improve model performance by tuning model  
641 parameters. CLASS-CTEM model also participated in the 2016 TRENDY intercomparison of  
642 terrestrial ecosystem models whose results contributed to the global carbon project (Le Quéré et  
643 al., 2016). The competition module of the CLASS-CTEM modelling framework has been  
644 previously evaluated at point scales (Arora and Boer, 2006; Shrestha et al., 2016). In addition to  
645 assessing fractional coverage at which PFTs equilibrate, these point scale evaluations also assess  
646 the time the PFTs take to reach their equilibrium fractional coverages against empirical data and  
647 if the succession patterns are realistically simulated (e.g. grasses should colonize a given area  
648 before trees invade the area covered by grasses). This manuscript focusses on evaluation of the  
649 competition module of the CLASS-CTEM modelling framework at a regional scale.

Deleted: ¶

651 Dynamically simulated fractional coverages of PFTs adds another degree of freedom to a model  
652 compared to the case where the fractional coverages of its PFTs are specified. This is a more  
653 stringent test of a model's performance. Errors in the simulated geographical distribution of  
654 PFTs will, of course, lead to corresponding errors in the geographical distribution of primary  
655 terrestrial ecosystem carbon pools and fluxes. Yet, the CLASS-CTEM model is broadly able to  
656 reproduce the geographical distributions of GPP and LAI. Limitations, of course, remain. In  
657 particular, the simulated LAI and GPP are high in Alaska and in northern and Arctic Canada, and  
658 these variables are lower than their observation-based estimates in arid regions of the western  
659 United States. The simulated fractional vegetation coverage reflects these patterns.

Deleted: Allowing a terrestrial ecosystem model to

Deleted: its

Deleted: the

666 It is difficult to conclusively determine whether these model limitations are due to the limitations  
667 in the biogeochemistry parameterizations of the model for its existing PFTs or the simple  
668 structural limitation that the model does not represent shrub, moss and lichen PFTs. Shrubs are  
669 adapted to grow in arid and semi-arid regions, whether in cold or hot climates (where neither  
670 grasses nor trees are able to grow) and their representation in the model would likely help to  
671 increase the fractional vegetation cover in arid regions including those in the western United  
672 States. At high latitudes grass growth is inhibited by mosses and lichens which flourish in cold  
673 and damp conditions. A representation of moss and lichen PFTs and improved representation of  
674 permafrost in the model would likely help to decrease simulated grass coverage in Arctic  
675 regions. In the current version of the CLASS-CTEM model bioclimatic limits are used only for  
676 tree PFTs to ensure that these PFTs do not venture outside their pre-determined bioclimatic  
677 zones. In the model, bioclimatic limits are not used for grasses and their geographical  
678 distribution is entirely the result of plant physiological processes and their competitive  
679 interactions with the tree PFTs and amongst themselves. Since, in the Arctic region, grasses do  
680 not face competition from tree PFTs, and moss and lichen PFTs are not represented in the model,  
681 they are free to increase their expanse – climate permitting, of course. Another possible reason  
682 for higher than observed grass coverage in the Arctic region is that in the current implementation  
683 of CLASS only three permeable soil layers with maximum thicknesses of 0.1, 0.25 and 3.75 m  
684 are represented and a boundary condition of zero heat flux is assumed across the bottommost  
685 layer. This simple representation does not allow to model permafrost realistically. Permafrost is  
686 more realistically modelled with multiple permeable and impermeable (extending into the bed  
687 rock) layers that go sufficiently deep (> 30 m at least) to capture the slow evolution of soil  
688 temperatures in response to climate warming (Teufel et al., 2017). The current set up of three  
689 layers that go only 4.1 m deep produces soil temperatures that are warmer than in the set up  
690 when permeable and impermeable layers are sufficiently deep and produces permafrost extent  
691 that is lower than observation-based estimates (Koven et al., 2013). It is likely that warmly  
692 biased soil temperatures in the current set up contribute to promote grass growth and allow it to  
693 cover a larger area in the Arctic region than would be the case when permafrost is more  
694 realistically modelled.

695

696 The lower than observed fractional vegetation cover in the arid and semi-arid regions of the  
697 western United States, however, may not solely be due to model limitations alone. Here, we  
698 argue that the manner in which remotely sensed land cover types are mapped to CTEM PFTs,  
699 and the errors in calculating bare ground fraction in remotely sensed products also contribute to  
700 mismatch between modelled and observation-based values of fractional vegetation cover. We  
701 illustrate this by comparing the functional relationship between LAI and total vegetation cover.  
702 Figure 12a shows this relationship for model simulated values. As expected, as LAI increases so  
703 does the total vegetation cover. The relationship between these two variables is fairly tight in the  
704 model and the green line is an exponential fit. The red dots in the figure correspond to grid cells  
705 that lie in the region identified in the inset in Figure 12d and broadly correspond to the western  
706 half of the United States. Figures 12b and 12c show the same relationship but between the  
707 observation-based estimate of LAI from Zhu et al. (2013) (as mentioned in Section 2.3.2) and the  
708 total vegetation cover based on the WANG06 and MODIS derived land cover data sets,  
709 respectively. The blue and magenta lines in Figures 12b and 12c are the corresponding  
710 exponential fits. When compared with Figure 12a, Figures 12b and 12c show much more scatter  
711 around the fitted curves, and the overall relationship appears to break down for the red dots  
712 corresponding to the grid cells in the western United States. A careful look at the red dots in  
713 Figures 12b and 12c shows that the observation-based vegetation cover in the Western United  
714 States for a large fraction of grid cells is around 60% regardless of the observation-based LAI  
715 which ranges between 0.1 and 1.5  $\text{m}^2/\text{m}^2$ . Clearly, it is physically unrealistic to achieve fractional  
716 vegetation coverage of 60% below LAI values of 0.6  $\text{m}^2/\text{m}^2$  (the  $\text{m}^2/\text{m}^2$  unit implies  $\text{m}^2$  of leaf  
717 area per  $\text{m}^2$  of ground area) and this indicates that the fractional vegetation cover in this region is  
718 likely overestimated in both observation-based data sets.

719  
720 There are at least two ways in which errors in total vegetation cover can occur. The first relates  
721 to the method by which the fractional vegetation cover is calculated for the land cover types in  
722 the original remotely sensed land cover products: that is, for the 22 land cover types in the  
723 GLC2000 data set upon which the WANG06 data are based and the 17 land cover types in the  
724 MODIS data set. An example of such an error for arid regions is illustrated by Lawley et al.  
725 (2014) who suggest that the MODIS soil fractional cover product, at least in its present form, is  
726 unsuited to monitoring sparsely vegetated arid landscapes and generally unable to separate soil

727 from vegetation in situations where normalized difference vegetation index (NDVI) is low. The  
728 second way in which errors are introduced is through the mapping of the remotely sensed land  
729 cover types to the CTEM PFTs following Table 2 of WANG06 for the GLC2000 land cover  
730 types, and following Table 2 in this manuscript for the MODIS land cover types. This mapping  
731 is based on available information in the literature but is also based on expert judgement which  
732 introduces subjectiveness. For instance, it is debatable what fraction of the “open shrublands”  
733 MODIS land cover type, which exists over much of the arid southwestern United States, is in  
734 fact bare ground. In Table 2, we have allocated a fraction of 0.4 of “open shrublands” to bare  
735 ground following WANG06. Had WANG06 allocated a higher value than this to bare ground,  
736 our simulated values would have compared better with the observation-based values of bare  
737 ground fraction over arid regions. Nevertheless this would not have changed the relationship, or  
738 rather the lack thereof, between the observation-based estimates of LAI and the total vegetation  
739 cover in the western half of the United States seen in Figures 12b and 12c.

Deleted: value

Deleted: to this fraction

740  
741 Both model and observation-based results are also affected by a common limitation associated  
742 with peatlands which exists in the Hudson Bay lowlands region. Both the GLC2000 data set,  
743 upon which the modified WANG06 land cover product is based, and the MODIS land cover data  
744 do not represent peatland vegetation. In these data sets the peatland vegetation is classified either  
745 as grasses, shrubs or trees. The model also does not represent peatlands and as a result the model  
746 grows trees and grasses in regions where peatlands exists. Work is under way to incorporate a  
747 peatland model developed for CLASS-CTEM (Wu et al., 2016) into our modelling framework.

748  
749 The simulated areas covered by the primary two tree PFTs (NDL EVG and BDL DCD COLD)  
750 have their weaknesses but large differences also exist between the two observation-based  
751 estimates especially for the NDL EVG PFT. Modelling competition between two tree PFTs is  
752 much more difficult than between trees and grasses. In the latter case trees are always considered  
753 superior to grasses, but in the case of competition between two tree PFTs the superiority is based  
754 on parameterized colonization rates which depend on simulated NPP. Based on comparisons  
755 with observation-based estimates, the main limitation in model results here is that the model  
756 overestimates the coverage of NDL EVG trees, and underestimates the coverage of BDL DCD  
757 COLD trees in the eastern United States, while the opposite is true in western Canada. The

Deleted: ¶

761 model, of course, does not represent individual species, while in the real world competition  
762 occurs at the species level that is modulated by soils and nutrient availability. An example that  
763 illustrates this limitation of the model is the Jack pine tree species which occupies ecological  
764 niche of nutrient poor soils in Boreal Canada (e.g. see Ste-Marie et al., 2007). The coupling of  
765 carbon and nutrient cycles is currently not represented in CLASS-CTEM and optimizing model  
766 parameters for hundreds of species is currently extremely difficult given limited available data at  
767 the species level. Most likely before the model is applied at the species level, as a first step, the  
768 number of PFTs represented in the model should be increased. An example of how additional  
769 PFTs in the CLASS-CTEM framework can lead to improved model performance is illustrated by  
770 Peng et al. (2014). This application of the model shows how sub-dividing the NDL EVG PFT  
771 into coastal and interior types for the province of British Columbia in Canada leads to  
772 improvement in simulated LAI and GPP. A recent attempt to explicitly represent physiological  
773 process in a model to simulate competition between needleleaf and broadleaf cold deciduous  
774 trees at a regional scale is illustrated in (Fisher et al., 2015) who incorporated the concepts from  
775 the Ecosystem Demographics (ED) model into the community land model – dynamic vegetation  
776 model (CLM-DGVM). Their results provide some interesting insights; however, validation of  
777 this approach at the global scale over a wide range of PFTs remains challenging.

Deleted: .

Deleted: One

Deleted: ,

778  
779 Finally, one of the objectives of this study was to evaluate if resolving climate niches by  
780 performing CLASS-CTEM simulation at a finer resolution of 1° in this study allowed improved  
781 simulation of geographical distribution of PFTs than in the Melton and Arora (2016) study that  
782 evaluated the competition module of the CLASS-CTEM model at 3.75° spatial resolution at the  
783 global scale. Figure 11 addresses this objective and shows that while the spatial correlations and  
784 RMSDs between the simulated and the modified WANG06 land cover product for fractional  
785 coverage of tree, grass and total vegetation are fairly similar for the model outputs at 1° and  
786 3.75° resolutions, these metrics are somewhat better for model's application at the coarser 3.75°  
787 resolution. One possible reason for the slightly worse model performance at the finer resolution  
788 is that while climate niches are resolved better at the finer resolution the model does not have the  
789 additional differentiation in PFTs (the number of model PFTs is still nine) that is required to gain  
790 benefit from the resolved climate niches. In addition, comparing Melton and Arora (2016) results  
791 over North America with ones obtained here we note that the primary model limitations remain

Deleted: s

Deleted: 2

Deleted: through 4 of Melton and Arora (2016) compare simulated geographical distributions of PFTs with WANG06 data

Deleted: . C

Deleted: their

802 unchanged in the application of the model at both spatial resolutions. These include lower  
803 simulated fractional vegetation coverage in the arid south-west North American region and  
804 higher in the Arctic region (due to higher grass coverage). In addition, in both applications of the  
805 model the differences in simulated geographical distribution of NDL EVG and BDL DCD CLD  
806 PFTs, compared to the WANG06 land cover data, are also similar. Model differences, compared  
807 to the WANG06 data, therefore remain more or less similar in the application of the model at  
808 both spatial resolutions. These results are, however, based on offline applications of the CLASS-  
809 CTEM model where it is driven by reanalysis data. In a fully-coupled simulation where CLASS-  
810 CTEM is coupled to an atmospheric model it is possible that model performance at low spatial  
811 resolution is different from its performance at high spatial resolution.

812

813 The comparison between observation-based and simulated fractional coverages is the most  
814 robust at the basic tree-grass-bare ground level. The subjectiveness introduced in the process of  
815 mapping remotely sensed land cover types to the PFTs represented in a model, as mentioned  
816 above, makes the comparison of simulated and observation-based fractional coverages for  
817 individual PFTs less robust. Nevertheless, comparisons with observations allow useful insights  
818 into model limitations as we have seen here.

819

820

## 821 **5 Summary and conclusions**

822

823 This study evaluates the CLASS-CTEM simulated fractional coverages of PFTs, when driven  
824 with observed meteorological forcing, against the observation-based estimates from MODIS and  
825 the modified WANG06 data sets over the North American region. In the past, performance of the  
826 competition module of the CLASS-CTEM modelling framework has been assessed at global  
827 scale, at a coarse spatial resolution of 3.75° (Melton and Arora, 2016), as well as at point scale,  
828 for a range of locations across the globe (Shrestha et al., 2016). Our objective here was to assess  
829 the performance of the CLASS-CTEM competition module at a higher spatial resolution of 1°  
830 over North America. To achieve this objective we compared simulated present day geographical  
831 distributions of fractional coverages of PFTs, but also LAI and GPP with their observation-based  
832 estimates.

Deleted: climate

834

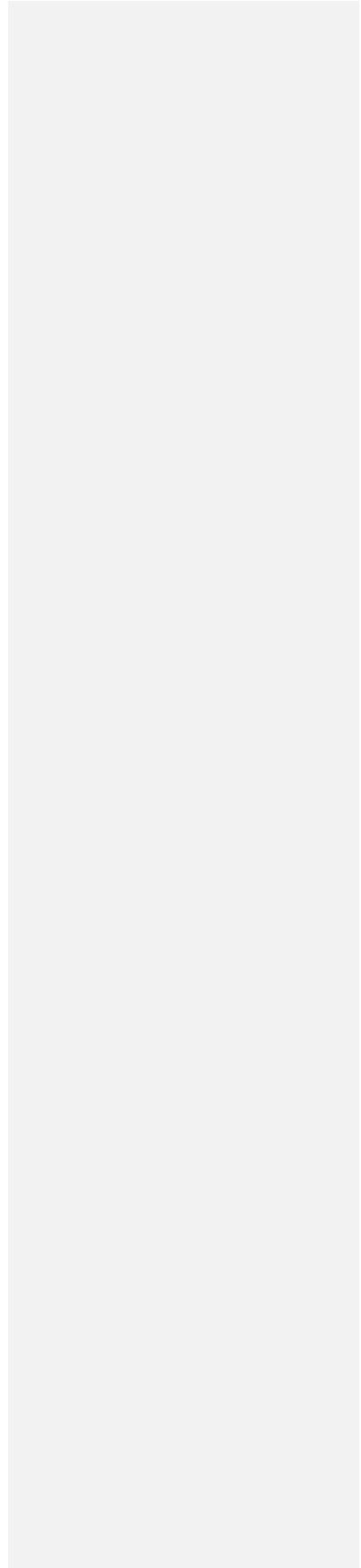
835 The CLASS-CTEM modelling framework is generally able to reproduce the dominant features  
836 of the geographic distribution of PFT coverage, and LAI and GPP over the North American  
837 region. After 1960, the model simulates increasing GPP and LAI in response to changing climate  
838 as well as increased atmospheric CO<sub>2</sub> concentrations and the resulting sink for the 1990s and  
839 2000s is broadly consistent with other estimates.

840

841 The simulated geographical distribution of PFTs, when compared to observation-based  
842 estimates, show two primary limitations which are excessive grass cover in the Arctic region and  
843 low vegetation cover in the arid western United States, although for the latter the observation-  
844 based estimates themselves may have their own weaknesses. There are three main factors in the  
845 CLASS-CTEM modelling framework that may have contributed to these differences: 1) the  
846 absence of a shrub PFT, which we believe is the reason for low simulated vegetation coverage in  
847 the arid to semi-arid western United States, 2) the absence of moss and lichen PFTs that may  
848 inhibit the establishment of grasses, and 3) probably a lack of sensitivity of C<sub>3</sub> grasses to high  
849 latitude climate and an inadequate representation of permafrost. Future model developments will  
850 focus on these aspects with a view to improving model performance.

851

852



854 **References:**

- 855 Anav, A., Friedlingstein, P., Kidston, M., Bopp, L., Ciais, P., Cox, P., Jones, C., Jung, M.,  
856 Myneni, R., and Zhu, Z.: Evaluating the land and ocean components of the global carbon  
857 cycle in the CMIP5 earth system models, *J. Climate*, 26, 6801-6843, doi:10.1175/JCLI-  
858 D-12-00417.1, 2013.
- 859
- 860 Arora, V. K.: Land surface modelling in general circulation models: a hydrological perspective,  
861 PhD Thesis, University of Melbourne, Australia, 1997.
- 862
- 863 [Arora, V.K.: Simulating energy and carbon fluxes using coupled land surface and terrestrial  
864 ecosystem models, \*Agricultural and Forest Meteorology\*, 118\(1-2\), 21-47, 2003.](#)
- 865
- 866 Arora, V. K. and Boer, G. J.: A representation of variable root distribution in dynamic vegetation  
867 models, *Earth Interactions*, 7, 1-19, 2003.
- 868
- 869 Arora, V. K. and Boer, G. J.: A parameterization of leaf phenology for the terrestrial ecosystem  
870 component of climate models, *Glob. Change Biol.*, 11, 39-59, 2005.
- 871
- 872 Arora, V. K. and Boer, G. J.: Simulating Competition and Coexistence between Plant Functional  
873 Types in a Dynamic Vegetation Model, *Earth Interactions*, 10, 1-29, 2006.
- 874
- 875 [Arora, V.K. and G.J. Boer: Uncertainties in the 20th century carbon budget associated with land  
876 use change, \*Global Change Biology\*, 16\(12\), 3327-3348, 2010.](#)
- 877
- 878 Arora, V. K., Boer, G. J., and Friedlingstein, P. E. A.: Carbon-concentration and carbon-climate  
879 feedbacks in CMIP5 earth system models, *J. Climate*, 26, 5289-5314, 2013.
- 880
- 881 [Arora, V. K., Y. Peng, W. A. Kurz, J. C. Fyfe, B. Hawkins, and A. T. Werner: Potential near-  
882 future carbon uptake overcomes losses from a large insect outbreak in British Columbia,  
883 Canada, \*Geophys. Res. Lett.\*, 43, doi:10.1002/2015GL067532, 2016.](#)
- 884
- 885 Beer, C., Reichstein, M., Tomelleri, E., Ciais, P., Jung, M., Carvalhais, N., Rodenbeck, C.,  
886 Arain, M. A., Baldocchi, D., Bonan, G. B., Bondeau, A., Cescatti, A., Lasslop, G.,  
887 Lindroth, A., Lomas, M., Luyssaert, S., Margolis, H., Oleson, K. W., Rouspard, O.,  
888 Veenendaal, E., Vivoy, N., Williams, C., Woodward, F. I., and Papale, D.: Terrestrial  
889 gross carbon dioxide uptake: global distribution and covariation with climate, *Science*,  
890 329, 834-838, doi:10.1126/science.1184984, 2010.
- 891
- 892 Bonan, G. B.: Forests and climate change: Forcings, feedbacks, and the climate benefits of  
893 forests, *Science*, 320, 1444-1449, doi:10.1126/science.1155121, 2008.
- 894
- 895 Box, E. O.: Plant functional types and climate at the global scale, *J. Veg. Sci.*, 7, 309-320,  
896 doi:10.2307/3236274, 1996.
- 897

898 Brentnall, S. J., Beerling, D. J., Osborne, C. P., Harland, M., Francis, J. E., Valdes, P. J., and  
899 Wittig, V. E.: Climatic and ecological determinants of leaf lifespan in polar forests of the  
900 high CO<sub>2</sub> Cretaceous 'greenhouse' world, *Glob. Change Biol.*, 11, 2177-2195, 2005.  
901

902 Broxton, P., X. Zeng, D. Sulla-Menashe, and P. Troch, 2014: A Global Land Cover Climatology  
903 Using MODIS Data. *J. Appl. Meteor. Climatol.*, 53, 1593–1605, doi: 10.1175/JAMC-D-  
904 13-0270.1.  
905

906 Cox, P. 2001: Description of the "TRIFFID" Dynamic Global Vegetation Model. Tech. Note 24.  
907

908 Cramer, W., Bondeau, A., Woodward, F. I., Prentice, I. C., Betts, R. A., Brovkin, V., Cox, P. M.,  
909 Fisher, V., Falloon, P. D., Foley, J., Friend, A. D., Kucharik, C., Lomas, M. R.,  
910 Ramankutty, N., Sitch, S., Smith, B., White, A., and Molling-Young, C.: Global response  
911 of terrestrial ecosystem structure and function to CO<sub>2</sub> and climate change: results from  
912 six dynamic global vegetation models, *Global Change Biology*, 7, 357-373, 2001.  
913

914 Crevoisier, C., Sweeney, C., Gloor, M., Sarmiento, J. L., and Tans, P. P.: Regional US carbon  
915 sinks from the three-dimensional atmospheric CO<sub>2</sub> sampling, *PANAS*, 107, 18348-  
916 18353, DOI: 10.1073/pnas.0900062107, 2010.  
917

918 Dai, Y., Zeng, X., Dickinson, R. E., and Coauthors: Common Land Model (CLM), Technical  
919 documentation and user's guide. [Available online at  
920 <http://climate.eas.gatech.edu/dai/clmdoc.pdf>], 2001.  
921

922 [FAO/IIASA/ISRIC/ISS-CAS/JRC: Harmonized World Soil Database \(version 1.2\), available at:  
923 \[http://www.fao.org/soils-portal/soil-survey/soil-maps-and-databases/harmonized-world-  
924 soil-database-v12/en/\]\(http://www.fao.org/soils-portal/soil-survey/soil-maps-and-databases/harmonized-world-soil-database-v12/en/\) \(last access: 20 January 2016\), 2012.](http://www.fao.org/soils-portal/soil-survey/soil-maps-and-databases/harmonized-world-soil-database-v12/en/)  
925

926 Fisher, R. A., Muszala, S., Versteinstein, M., Lawrence, P., Xu, C., McDowell, N. G., Knox, R.  
927 G., Koven, C., Holm, J., Rogers, B. M., Spessa, A., Lawrence, D., and Bonan, G.: Taking  
928 off the training wheels: the properties of a dynamic vegetation model without climate  
929 envelopes, *CLM4.5(ED)*, *Geosci. Model Dev.*, 8, 3593-3619, doi:10.5194/gmd-8-3593-  
930 2015, 2015.  
931

932 Friedl, M., Strahler, A., Schaaf, C., Hodges, J. C. F., and Salomon, J., 2013.: Binary MODIS  
933 MOD12C1 0.25 Degree Land Cover Climate Modeler Grid. Available at  
934 [<http://duckwater.bu.edu/lc/>] from the Department of Geography, Boston University,  
935 Boston, Massachusetts, USA., 2013.  
936

937 Friedlingstein, P., Cox, P., Betts, R., Bopp, L., Von Bloh, W., Brovkin, V., Cadule, P., Doney,  
938 S., Eby, M., Fung, I., Bala, G., John, J., Jones, C., Joos, F., Kato, T., Kawamiya, M.,  
939 Knorr, W., Lindsay, K., Matthews, H. D., Raddatz, T., Rayner, P., Reick, C., Roeckner,  
940 E., Schnitzler, K.-G., Schnur, R., Strassmann, K., Weaver, A. J., Yoshikawa, C., and  
941 Zeng, N.: Climate–carbon cycle feedback analysis: Results from the C4MIP model  
942 intercomparison, *J. Climate*, 19, 3337-3353, 2006.  
943

- 944 Friend, A. D., Chard, Lucht, W., Rademacher, T., Keribin, R., Betts, R., Cadule, P., Ciais, P.,  
 945 Clark, D. B., Dankers, R., Falloon, P. D., Ito, A., Kahana, R., Kleidon, A., Lomas, M. R.,  
 946 Nishina, K., Ostberg, S., Pavlick, R., Peylin, P., Schaphoff, S., Vuichard, N.,  
 947 Warszawski, L., Wiltshire, A., and Woodward, F. I.: Carbon residence time dominates  
 948 uncertainty in terrestrial vegetation response to future climate and atmospheric CO<sub>2</sub>,  
 949 Proc. Natl. Acad. Sci. USA, 111, 3280-3285, 2013.
- 950
- 951 Garnaud, C., Sushama, L., and Verseghy, D.: Impact of interactive vegetation phenology on the  
 952 Canadian RCM simulated climate over North America, Climate Dynamics, 45, 1471-  
 953 1492, doi:10.1007/s00382-014-2397-9, 2015.
- 954
- 955 Gobron, N., Belward, A., and Knorr, W.: Monitoring biosphere vegetation 1998-2009, Geophys.  
 956 Res. Lett., 37, L15402, 2010.
- 957
- 958 Hurtt, G. C., Chini, L. P., Frolking, S., Betts, R. A., Feddema, J., Fischer, G., Fisk, J. P.,  
 959 Hibbard, K., Houghton, R. A., Janetos, A., Jones, C. D., Kindermann, G., Kinoshita, T.,  
 960 Klein Goldewijk, K., Riahi, K., Shevliakova, E., Smith, S., Stehfest, E., Thomson, A.,  
 961 Thornton, P., Van Vuuren, D. P., and Wang, Y. P.: Harmonization of land-use scenarios  
 962 for the period 1500-2100: 600 years of global gridded annual land-use transitions, wood  
 963 harvest, and resulting secondary lands, Climate Change, 109, 117-161, 2011.
- 964
- 965 Kramer, P. J. and Kozlowski, T. T., 1979: Physiology of woody plants. Academic press, 1979.
- 966
- 967 Koven, C. D., Riley, W. J. and Stern, A.: Analysis of Permafrost Thermal Dynamics and  
 968 Response to Climate Change in the CMIP5 Earth System Models, J. Clim., 26(6), 1877-  
 969 1900, 2013.
- 970
- 971 Lawley, E. F., Lewis, M. M., and Ostendorf, B.: Evaluating MODIS soil fractional cover for arid  
 972 regions, using albedo from high-spatial resolution satellite imagery, Int. J. Remote Sens.,  
 973 35, 2028-2046, 2014.
- 974
- 975 [Le Quéré, C., Andrew, R. M., Canadell, J. G., Sitch, S., Korsbakken, J. I., Peters, G. P.,](#)  
 976 [Manning, A. C., Boden, T. A., Tans, P. P., Houghton, R. A., Keeling, R. F., Alin, S.,](#)  
 977 [Andrews, O. D., Anthoni, P., Barbero, L., Bopp, L., Chevallier, F., Chini, L. P., Ciais, P.,](#)  
 978 [Currie, K., Delire, C., Doney, S. C., Friedlingstein, P., Gkritzalis, T., Harris, I., Hauck, J.,](#)  
 979 [Haverd, V., Hoppema, M., Klein Goldewijk, K., Jain, A. K., Kato, E., Körtzinger, A.,](#)  
 980 [Landschützer, P., Lefèvre, N., Lenton, A., Lienert, S., Lombardozzi, D., Melton, J. R.,](#)  
 981 [Metzl, N., Millero, F., Monteiro, P. M. S., Munro, D. R., Nabel, J. E. M. S., Nakaoka, S.-](#)  
 982 [I., O'Brien, K., Olsen, A., Omar, A. M., Ono, T., Pierrot, D., Poulter, B., Rödenbeck, C.,](#)  
 983 [Salisbury, J., Schuster, U., Schwinger, J., Séférian, R., Skjelvan, I., Stocker, B. D.,](#)  
 984 [Sutton, A. J., Takahashi, T., Tian, H., Tilbrook, B., van der Laan-Luijkx, I. T., van der](#)  
 985 [Werf, G. R., Viovy, N., Walker, A. P., Wiltshire, A. J., and Zaehle, S.: Global Carbon](#)  
 986 [Budget 2016, Earth Syst. Sci. Data, 8, 605-649, doi:10.5194/essd-8-605-2016, 2016.](#)
- 987
- 988 Melton, J. R. and Arora, V. K.: Competition between plant functional types in the Canadian  
 989 Terrestrial Ecosystem Model (CTEM) v. 2.0, Geosci. Model Dev., 9, 323-361, 2016.

990  
991 [Melton, J. R. and Arora, V. K. \(2014\) Sub-grid scale representation of vegetation in global land](#)  
992 [surface schemes: Implications for estimation of the terrestrial carbon sink,](#)  
993 [Biogeosciences, 11\(4\), 1021-1036.](#)  
994  
995 [Melton, J. R., Shrestha, R. K., and Arora, V. K.: The influence of soils on heterotrophic](#)  
996 [respiration exerts a strong control on net ecosystem productivity in seasonally dry](#)  
997 [Amazonian forests, Biogeosciences, 12, 1151-1168, doi:10.5194/bg-12-1151-2015, 2015.](#)  
998  
999 Peng, Y., Arora, V. K., Kurz, W. A., Hember, R. A., Hawkins, B. J., Fyfe, J. C., and Werner, A.  
1000 T.: Climate and atmospheric drivers of historical terrestrial carbon uptake in the province  
1001 of British Columbia, Canada, Biogeosciences, 11, 635-649, doi:10.5194/bg-11-635-2014,  
1002 2014.  
1003  
1004 Pielke, R. A., Avissar, R., Raupach, M., Dolman, A. J., Zeng, X., and Denning, S.: Interactions  
1005 between the atmosphere and terrestrial ecosystem: influence on weather and climate,  
1006 Global Change Biology, 4, 461-475, 1998.  
1007  
1008 Ramankutty, N. and Foley, J. A.: Estimating historical changes in global land cover: Croplands  
1009 from 1700 to 1992, Global Biogeochem. Cycles, 13, 997-1027, 1999.  
1010  
1011 Ran, L., Pleim, J., Gilliam, R., Binkowski, F. S., Hogrefe, C., and Band, L.: Improved  
1012 meteorology from an updated WRF/CMAQ modeling system with MODIS vegetation  
1013 and albedo, J. Geophys. Res. Atmos., 121, 2393-2415, doi:10.1002/2015JD024406, 2016.  
1014  
1015 [Ruesch, A. and Holly, K.: New IPCC Tier-1 Global Biomass Carbon Map For the Year 2000,](#)  
1016 [available at: \[ftp://cdiac.ornl.gov/pub/global\\\_carbon/\]\(ftp://cdiac.ornl.gov/pub/global\_carbon/\) \(last access: 5 July 2013\), Carbon](#)  
1017 [Dioxide Information Analysis Center <http://cdiac.ornl.gov> \(last access: 5 July 2013\), Oak](#)  
1018 [Ridge National Laboratory, Oak Ridge, Tennessee, 2008.](#)  
1019  
1020 Ritchie, T. C. and Macdonald, G. M.: The patterns of post-glacial spread of White Spruce, J.  
1021 Biogeogr., 13, 527-540, 1986.  
1022  
1023 Shrestha, R. K., Arora, V. K., and Melton, J. R.: The sensitivity of simulated competition  
1024 between different plant functional types to sub-grid-scale representation of vegetation in  
1025 a land surface model, J. Geophys. Res. Biogeosci., 121, doi:10.1002/2015JG003234,  
1026 2016.  
1027  
1028 Siemann, E. and Rogers, W. E.: Changes in light and nitrogen availability under pioneer trees  
1029 may indirectly facilitate tree invasion of grasslands, J. Ecology, 91, 923-931, 2003.  
1030  
1031 Sitch, S., Smith, B., Prentice, I. C., Arneeth, A., Bondeau, A., Cramer, W., Kaplan, J. O., Lucht,  
1032 W., Sykes, M. T., Thonicke, K., and Venevsky, S.: Evaluation of ecosystem dynamics,  
1033 plant geography and terrestrial carbon cycling in the LPJ dynamic global vegetation  
1034 model, Glob. Change Biol., 9, 161-185, 2003.

1035 [Ste-Marie, C., Paré D. and Gagnon, D. : The contrasting effects of Aspen and Jack Pine on soil](#)  
1036 [nutritional properties depend on parent material, \*Ecosystems\*, 10\(8\), 1299-1310, 2007.](#)  
1037

1038 Teufel, B., Sushama, L., Arora, V., and Versegny, D.: Impact of dynamic vegetation phenology  
1039 on the simulated pan-Arctic land surface state, submitted to *Climate Dynamics*, 2017.  
1040

1041 Timmons, D., Buchholz, T., and Veeneman, C. H.: Forest biomass energy: assessing  
1042 atmospheric carbon impacts by discounting future carbon flows, *GCB Bioenergy*, 8, 631-  
1043 643, 10.1111/gcbb.12276, 2016.  
1044

1045 Versegny, D. L., Mcfarlane, N. A., and Lazare, M.: CLASS - A Canadian land surface scheme  
1046 for GCMs, II. Vegetation model and coupled runs, *Int. J. Climatol.*, 13, 347-370, 1993.  
1047

1048 Viovy, N.: CRU-NCEP reanalysis data version 4. Available at  
1049 [http://dods.extra.cea.fr/store/p529viouv/cruncep/V4\_1901\_2012/], (Accessed May 2015),  
1050 2012.  
1051

1052 Wang, A., Price, D. T., and Arora, V. K.: Estimating changes in global vegetation cover (1850-  
1053 2100) for use in climate models, *Global Biogeochem. Cycles*, 20, GB3028, 2006.  
1054

1055 Wang, G., Sun, S., and Mei, R.: Vegetation dynamics contributes to the multi-decadal variability  
1056 of precipitation in the Amazon region, *Geophys. Res. Lett.*, 38, L19703,  
1057 doi:10.1029/2011GL049017, 2011.  
1058

1059 Wen, X., Tang, G., Wang, S., and Huang, J.: Comparison of global mean temperature series,  
1060 *Advances in Climate Change Research*, 2, 187-192, 2011.  
1061

1062 [Wu, Y., Versegny, D. L. and Melton, J. R.: Integrating peatlands into the coupled Canadian Land](#)  
1063 [Surface Scheme \(CLASS\) v3.6 and the Canadian Terrestrial Ecosystem Model \(CTEM\)](#)  
1064 [v2.0, \*Geosci Model Dev\*, 9\(8\), 2639–2663, doi:10.5194/gmd-9-2639-2016, 2016.](#)  
1065

1066 Zhang, Z., Xue, Y., Macdonald, G., Cox, P. M., and Collatz, G. J.: Investigation of North  
1067 America vegetation variability under recent climate: A study using the  
1068 SSiB4/TRIFFID biophysical/dynamic vegetation model, *J. Geophys. Res. Atmos.*, 120,  
1069 1300-1321, 2015.  
1070

1071 Zhu, Z., Bi, J., Pan, Y., Ganguly, S., Anav, A., Xu, L., Samanta, A., Piao, S., Nemani, R. R., and  
1072 Myneni, R. B.: Global data sets of vegetation leaf area index (LAI)3g and fraction of  
1073 photosynthetically active radiation (FPAR)3g derived from global inventory modeling  
1074 and mapping studies (GIMMS) normalized difference vegetation index (NDVI3g) for the  
1075 period 1981- to 2011, *Remote Sens.*, 5, 927-948, doi:10.3390/rs5020927, 2013.  
1076

1077 Zobler, L.: A world soil file for global climate modelling, Tec. Rep. NASA TM-87802, 14-32,  
1078 1986.  
1079  
1080

Formatted: English (U.S.)

Formatted: English (U.S.)

Formatted: English (U.S.)

1081  
 1082  
 1083 Table 1: Plant functional types (PFTs) represented in CTEM and their relation to CLASS PFTs.  
 1084  
 1085

---

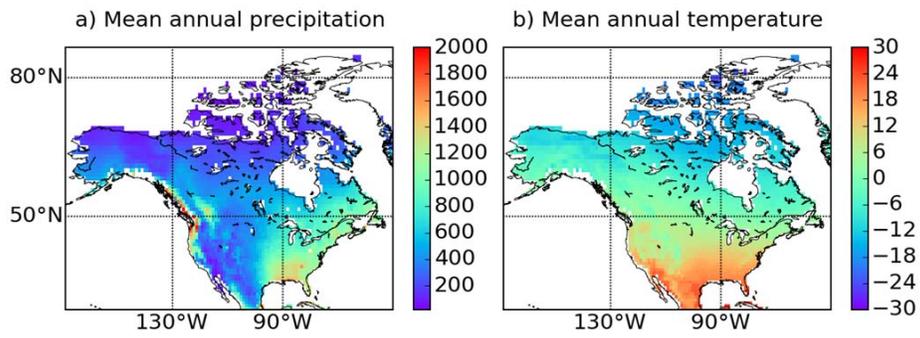
<b>CLASS PFTs</b>	<b>CTEM PFTs</b>	<b>CTEM PFT Symbol</b>
Needleleaf trees	Needleleaf Evergreen trees	NDL-EVG
	Needleleaf Deciduous trees	NDL-DCD
Broadleaf trees	Broadleaf Evergreen trees	BDL-EVG
	Broadleaf Cold Deciduous trees	BDL-DCD-CLD
	Broadleaf Drought/Dry Deciduous trees	BDL-DCD-DRY
Crops	C <sub>3</sub> Crops	CROP-C3
	C <sub>4</sub> Crops	CROP-C4
Grasses	C <sub>3</sub> Grasses	GRASS-C3
	C <sub>4</sub> Grasses	GRASS-C4

---

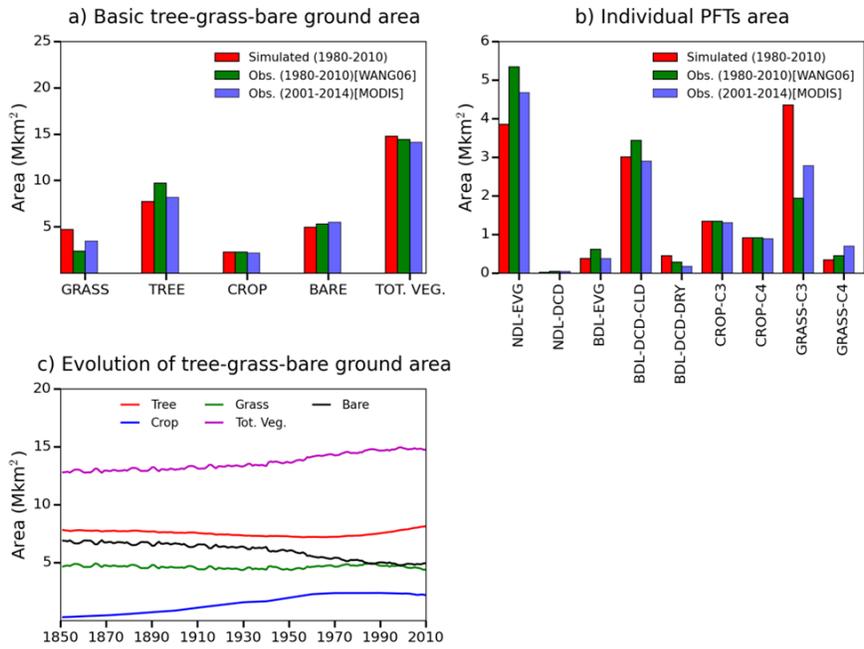
Table 2: Reclassification of the 17 MODIS land cover classes into the nine CTEM PFTs

SN	Items	Tree				Crop	Grass	Bare	Reference
		NDL EVG	NDL DCD	BDL EVG	BDL DCD				
1	Woody Savanna			0.1	0.4		0.25	0.25	Dai et al. (2001)
2	Water bodies							1	
3	Urban built up areas	0.05			0.05		0.1	0.8	Dai et al. (2001)
4	Savanna			0.05	0.3		0.4	0.25	Wang et al. (2006)
5	Permanent Wetlands						0.25	0.75	Dai et al. (2001)
6	Permanent snow and ice							1	Wang et al. (2006)
7	Open Shrublands	0.1			0.15		0.35	0.4	Wang et al. (2006)
8	Needleleaf evergreen	1							Wang et al. (2006)
9	Needleleaf deciduous		0.8				0.1	0.1	Wang et al. (2006)
10	Mixed forest	0.45			0.45		0.1		Wang et al. (2006)
11	Grasslands						0.65	0.35	Wang et al. (2006)
12	Croplands					0.9		0.1	Wang et al. (2006)
13	Cropland natural veg. mosaic			0.2		0.5	0.2	0.1	Wang et al. (2006)
14	Closed shrublands	0.2	0.2		0.4		0.2		Wang et al. (2006)
15	Broadleaf evergreen			1					Wang et al. (2006)
16	Broadleaf deciduous				1				Wang et al. (2006)
17	Bare ground							1	Wang et al. (2006)

1  
2  
3



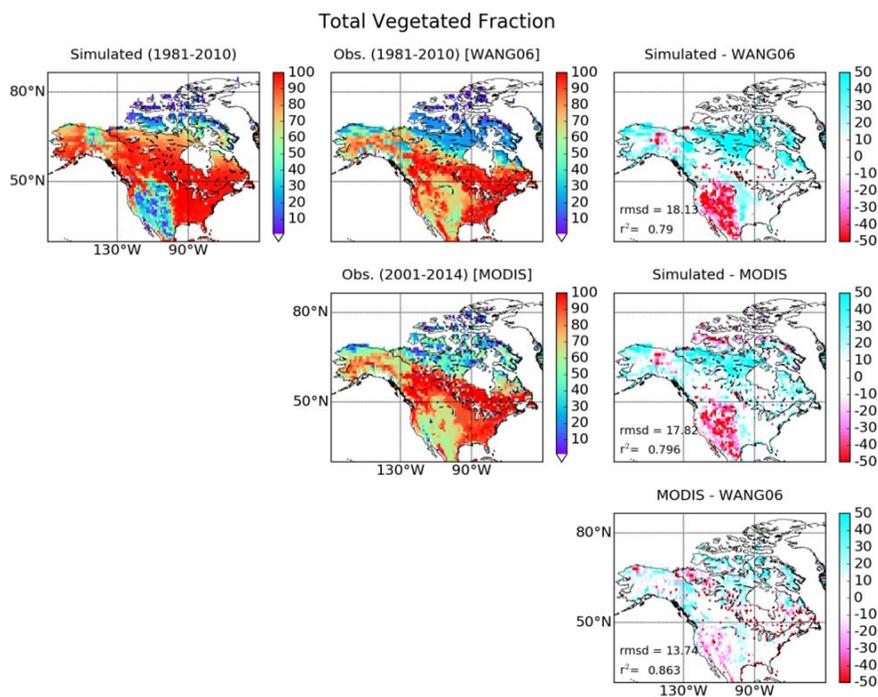
4  
5 Figure 1. Spatial distribution of mean annual a) precipitation (mm), and b) temperature (°C)  
6 across North America. Grid cells with permanent ice/glaciers have been masked out.  
7



8

9

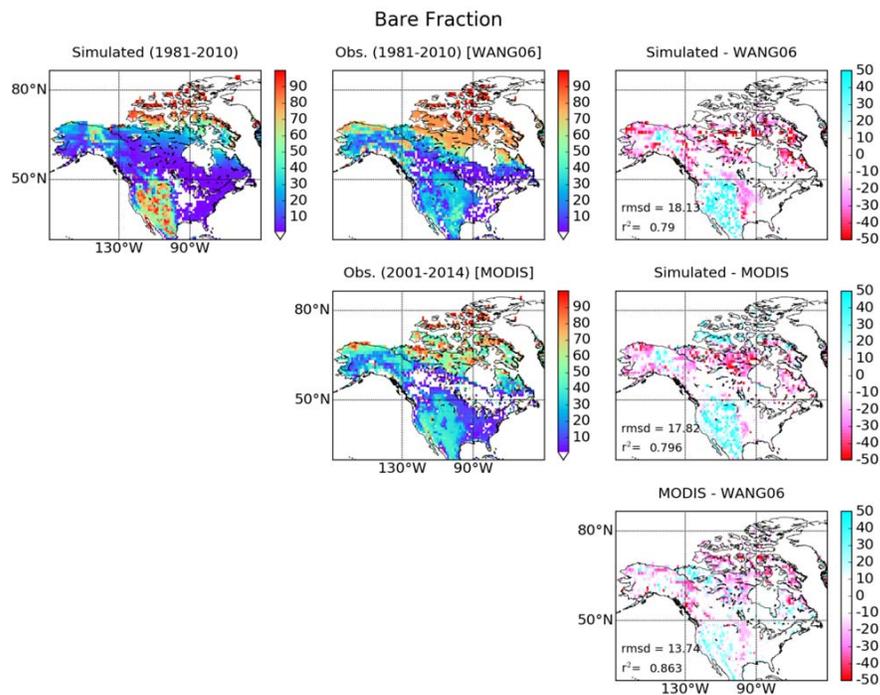
10 Figure 2. Comparison of observation-based and simulated vegetation areas summed over the  
 11 North American domain a) grass, treed, crop, bare ground and total vegetated area, b) individual  
 12 PFT areas, and c) evolution of simulated vegetation areas summed over the domain.  
 13



15

16 Figure 3. Spatial distribution of total vegetated coverage across North America. Simulated,  
 17 observation-based, and differences are presented in the left, middle and right columns,  
 18 respectively. The differences column includes model biases with respect to WANG06 (top panel)  
 19 and MODIS (middle panel), and the difference between the two observation-based estimates  
 20 (bottom panel). Root mean square difference (rmsd) and coefficient of determination ( $r^2$ ) are also  
 21 shown in each case.

22

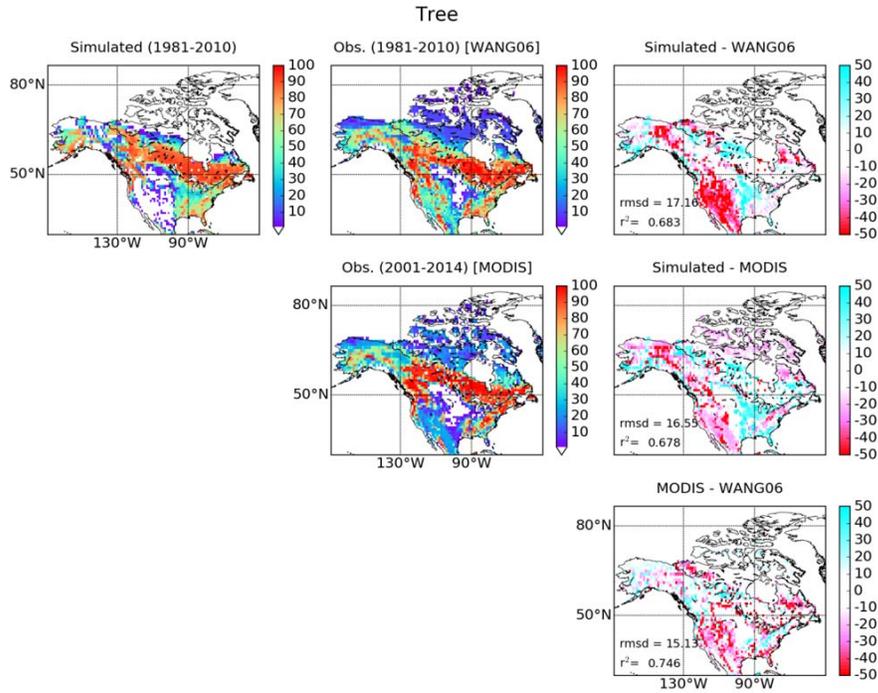


24

25 Figure 4. Spatial distribution of bare ground coverage across North America. Simulated,  
 26 observation-based, and differences are presented in the left, middle and right columns,  
 27 respectively. The differences column includes model biases with respect to WANG06 (top panel)  
 28 and MODIS (middle panel), and the difference between the two observation-based estimates  
 29 (bottom panel). Root mean square difference (rmsd) and coefficient of determination ( $r^2$ ) are also  
 30 shown in each case.

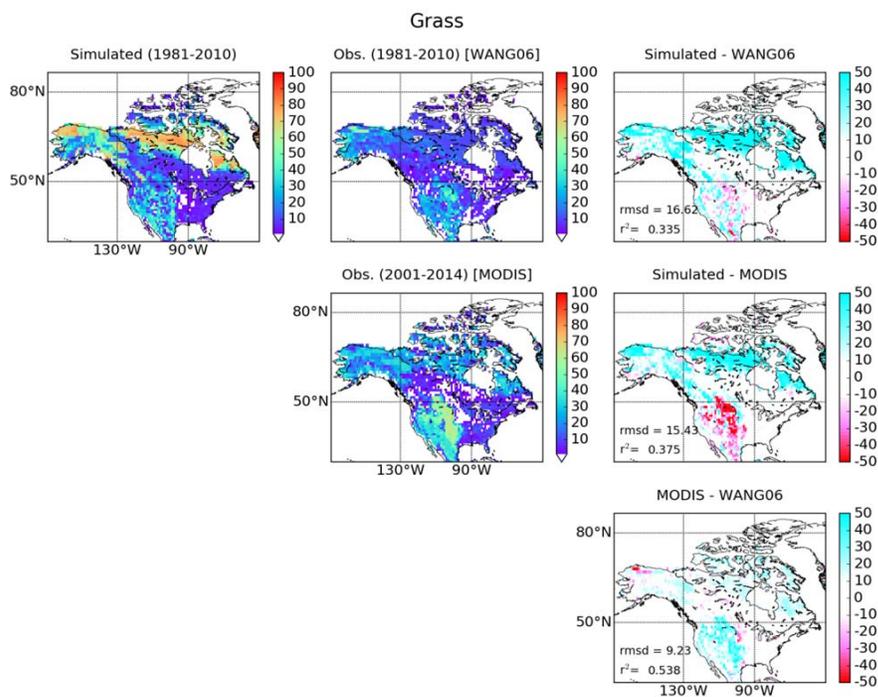
31

32



33  
 34 Figure 5. Spatial distribution of tree coverage across North America. Simulated, observation-  
 35 based, and differences are presented in the left, middle and right columns, respectively. The  
 36 differences column includes model biases with respect to WANG06 (top panel) and MODIS  
 37 (middle panel), and the difference between the two observation-based estimates (bottom panel).  
 38 Root mean square difference (rmsd) and coefficient of determination ( $r^2$ ) are also shown in each  
 39 case.

40



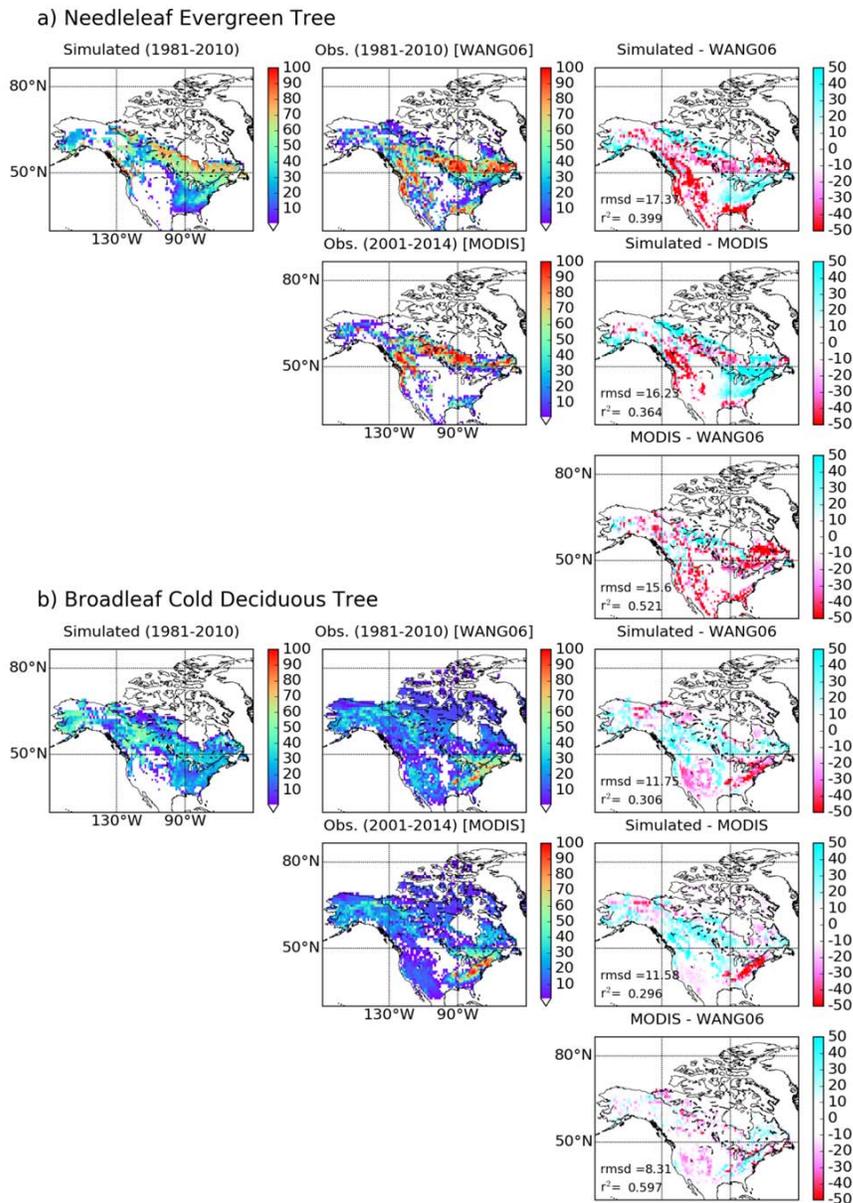
42

43 Figure 6. Spatial distribution of grass coverage across North America. Simulated, observation-  
 44 based, and differences are presented in the left, middle and right columns, respectively. The  
 45 differences column includes model biases with respect to WANG06 (top panel) and MODIS  
 46 (middle panel), and the difference between the two observation-based estimates (bottom panel).  
 47 Root mean square difference (rmsd) and coefficient of determination ( $r^2$ ) are also shown in each  
 48 case.

49

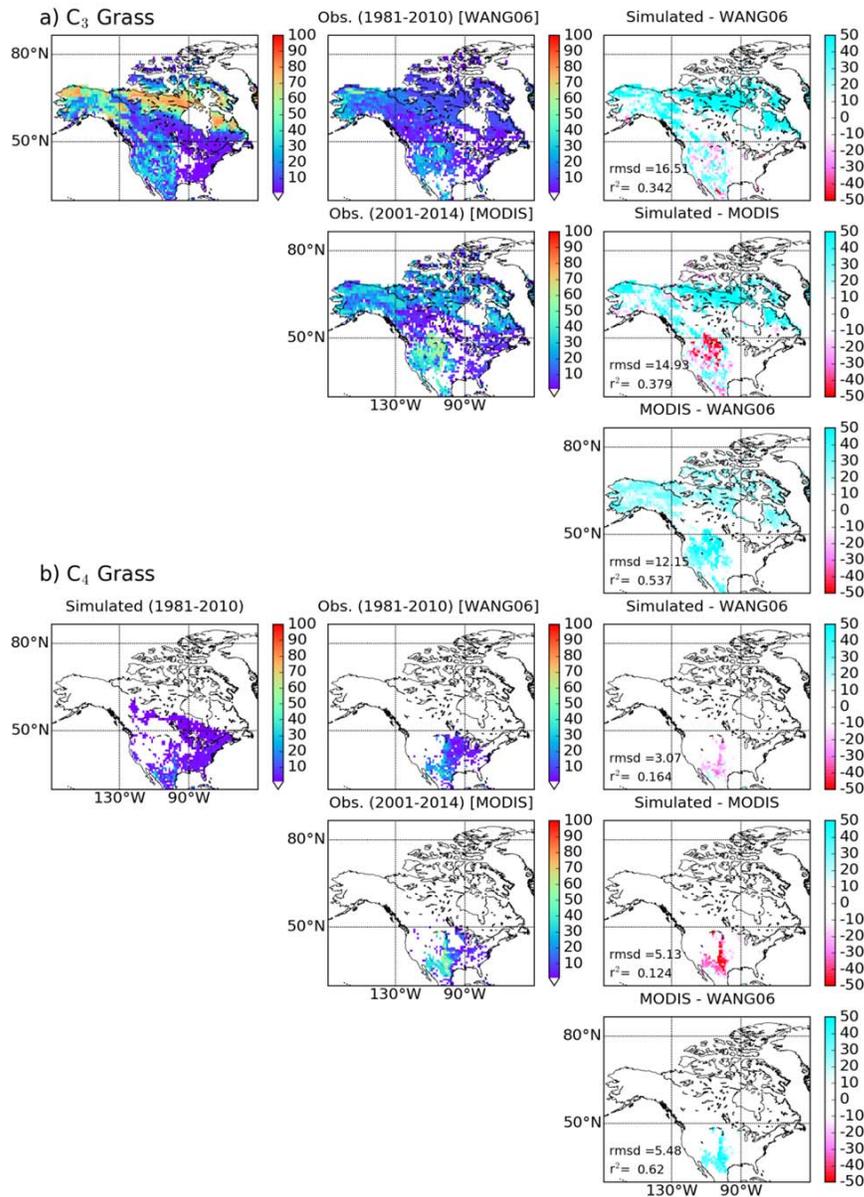
50

51



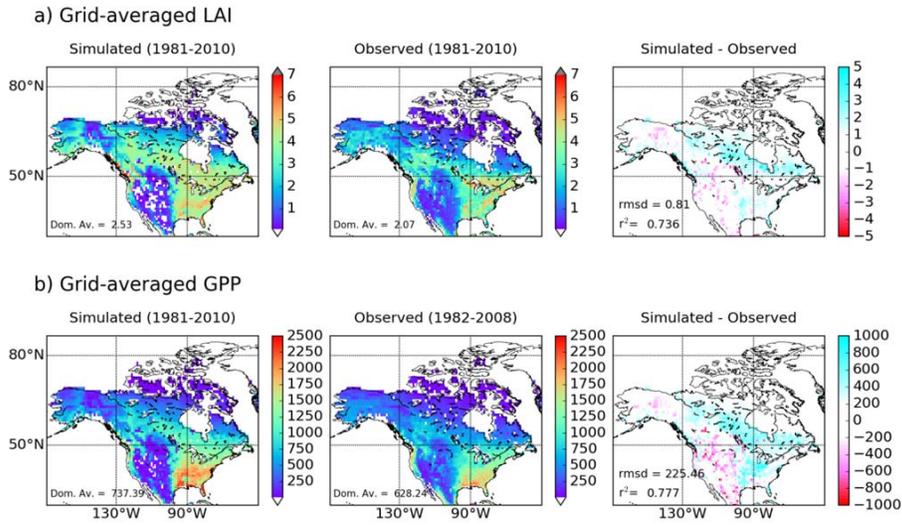
52

53 Figure 7. Spatial distribution of a) needleleaf evergreen tree, and b) broadleaf cold deciduous tree  
 54 across North America. Simulated, observation-based, and differences are presented in the left,  
 55 middle and right columns, respectively. The differences column includes model biases with  
 56 respect to WANG06 (top panel) and MODIS (middle panel), and the difference between the  
 57 observation-based estimates (bottom panel). Root mean square difference (rmsd) and coefficient  
 58 of determination ( $r^2$ ) are also shown in each case.



59  
 60 Figure 8. Spatial distribution of a) C<sub>3</sub> grasses, and b) C<sub>4</sub> grasses across North America.  
 61 Simulated, observation-based, and differences are presented in the left, middle and right  
 62 columns, respectively. The differences column includes model biases with respect to WANG06  
 63 (top panel) and MODIS (middle panel), and the difference between the the observation-based  
 64 estimates (bottom panel). Root mean square difference (rmsd) and coefficient of determination  
 65 ( $r^2$ ) are also shown in each case.

66  
67  
68  
69



70  
71  
72  
73  
74  
75  
76

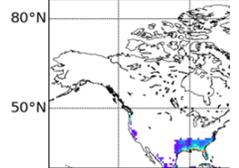
Figure 9. Spatial distribution of a) grid averaged maximum LAI ( $m^2 m^{-2}$ ), and b) grid averaged GPP ( $g C m^2 y^{-1}$ ) across North America. Simulated, observation-based, and differences between them are presented in the left, middle and right columns, respectively. Root mean square difference (rmsd) and coefficient of determination ( $r^2$ ) are also shown in each case.

Deleted: ¶

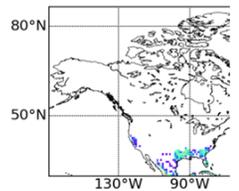
¶

a) Broadleaf Evergre

Simulated (1981-2010



b) Broadleaf Dry De



¶  
Figure 9. Spatial distribution of a) broadleaf evergreen tree, and b) broadleaf dry deciduous tree across North America. Simulated, WANG06 and MODIS distribution are presented in the left, middle and right columns, respectively. ¶

¶

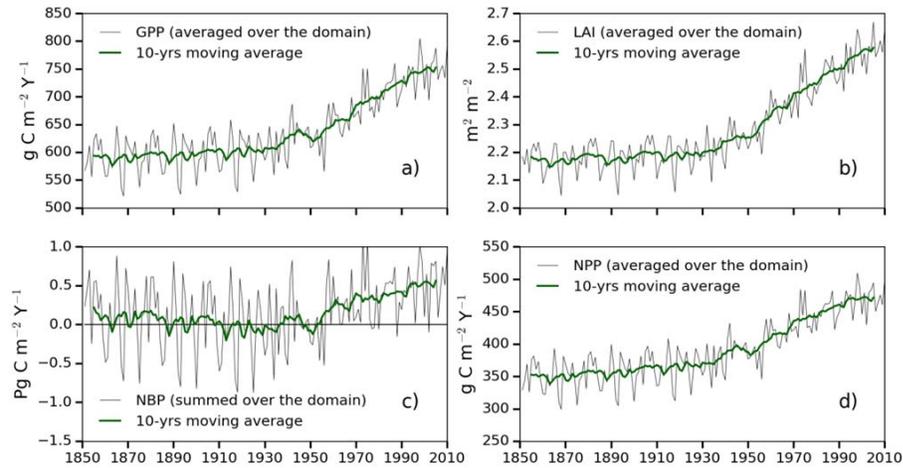
¶

¶

-----Page Break-----

Deleted: 10

94



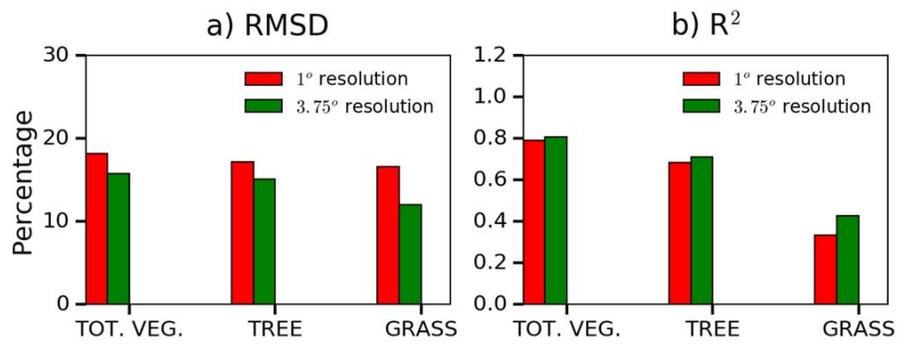
95

96

97 | Figure 10. Time series evolution of a) domain averaged GPP ( $\text{g C m}^{-2} \text{y}^{-1}$ ), b) domain averaged  
98 | LAI ( $\text{m}^2 \text{m}^{-2}$ ), c) domain total NBP ( $\text{Pg C m}^{-2} \text{y}^{-1}$ ), and d) domain averaged NPP ( $\text{g C m}^{-2} \text{y}^{-1}$ ).

99 |

100  
101  
102

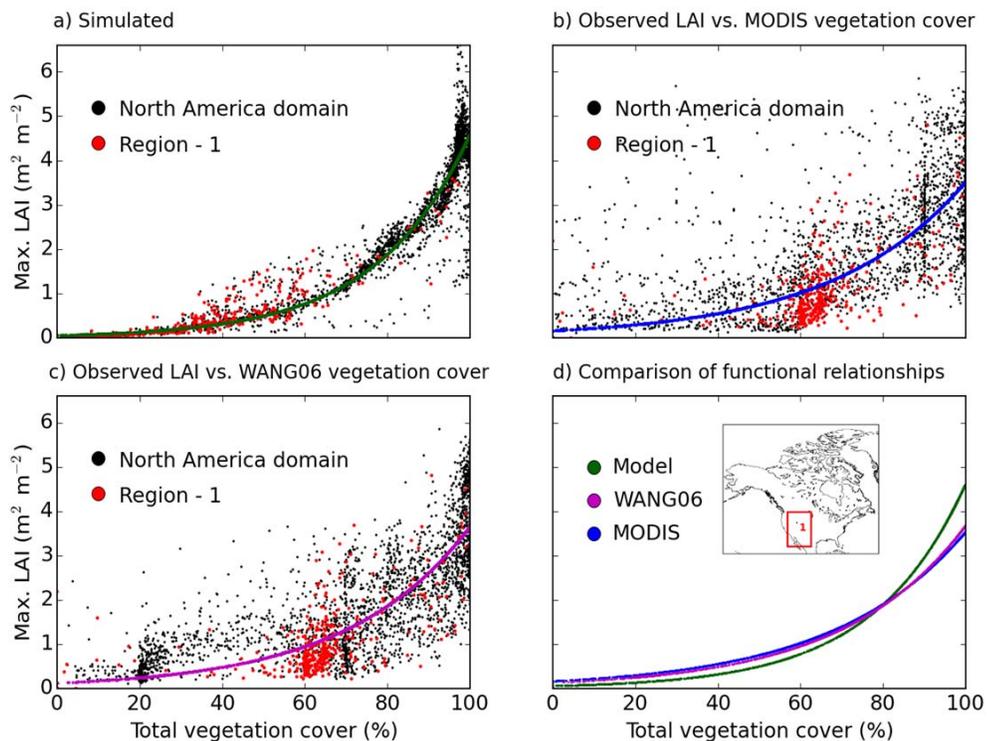


103  
104  
105  
106  
107  
108  
109  
110  
111

Figure 11. Comparison of the performance of the model at the 1° spatial resolution in this study with that at the 3.75° spatial resolution in the Melton and Arora (2016) study. The Melton and Arora (2016) global results were extracted for the North American domain. Spatial correlations and root mean square differences are used as metrics for the comparison between simulated and the observation-based estimate based on the modified WANG06 land cover product for fractional coverage of total vegetation, tree and grass.

112

113



114

115

116 Figure 12. Scatter plots of a) simulated LAI vs. simulated total vegetation coverage, b) observed  
117 LAI vs. MODIS-derived total vegetation coverage, c) observed LAI vs. WANG06 total  
118 vegetation coverage. Plot d) shows a comparison of the fitted curves represented by solid lines,  
119 with an inset map of North America showing the sub-domain of interest bounded by a red  
120 rectangle.

121

122