Interactive comment on “Seasonal pH and aragonite saturation horizons in the Gulf of Alaska during the North Pacific Survey, 1956–1957” by S. McKinnell and J. R. Christian

Anonymous Referee #2

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General Comments

McKinnell and Christian utilized an extensive data set of pH measurements collected by the Pacific Oceanographic Group of the Fisheries Research Board of Canada in the Gulf of Alaska collected in summer of 1956 and winter of 1957 to estimate aragonite saturation levels in the subarctic North Pacific in summer and winter. Using a number of assumptions and calculations based on this data set they suggested that the aragonite saturation estimates indicated that undersaturated water was a common feature of the surface mixed layer north of 51° N latitude in the winter of 1957. The North Pacific Survey data were compared with the results of a summer 2007 survey of the west coast of North America where pH levels were ~0.1 pH units lower than was found in the summer of 1956. The authors made a number of assumptions in making these calculations that, in some instances, are contradicted by the recent scientific literature. Furthermore, the authors chose to ignore much of the previous research studies on carbonate saturation in the North Pacific, which do not appear to support their findings. I recommend that the manuscript be returned to the authors for major revisions (see specific comments below).

Specific Comments

Page 4591. What is the accuracy of the pH measurements and how was it determined? How did the authors deal with the potential biases caused by the changes in liquid junction potential that occurs when standards of one ionic strength are compared with seawater samples of a significantly different ionic strength? What is the alkalinity-salinity relationship and what are the uncertainties associated with this estimate? What are the major uncertainties in the aragonite saturation calculations and how were they determined? Does the precision in the 1956-57 pH data contribute to the uncertainty in the aragonite saturation calculation? If so, how much? Are there any biases between the 1956-57 datasets relative to the recent data?

Pages 4591-92. The author’s aragonite saturation calculations are based on the assumption that alkalinity does not change over time due to anthropogenic CO2 uptake and carbonate mineral dissolution. This assumption is made despite the fact that the uptake of anthropogenic CO2 is decreasing the saturation state to values below 1.0 in subsurface waters and also in surface waters in wintertime north of 51° N! The authors do not make reference to the work of Sarma et al (2002) that showed that aragonite saturation horizon (ASH) shoaled significantly by 25 to 155 m in the North Pacific over
the last two decades. Over the same time interval, DIC increased from 12.5 to 36.8 
µmol kg\(^{-1}\) in the vicinity of ASH in the Pacific and total alkalinity increased significantly 
by 5 to 10 µmol kg\(^{-1}\) in the Pacific at the ASH, most likely as a result of aragonite dis-
solution. The authors need to address these results that contradict their fundamental 
assumption.

Page 4593. The authors state that their profiles of aragonite saturation as a function 
of depth indicate widespread undersaturation in the Gulf of Alaska in wintertime, es-
pecially in the north where median values of aragonite saturation state were <1 at all 
depths. Furthermore, their data suggest that the upper layer was supersaturated in 
summer in both the northern and southern Gulf of Alaska. Why are these results so 
different from the 1985 results published in Feely et al (1988) for the same region?
Feely et al (1988) show that the surface waters were supersaturated with respect to 
both aragonite and calcite in both winter and summer to depths greater than 100 m? 
Feely et al (1988) go on to suggest that surface water pCO2 values in excess of 650 
uatm are required before undersaturation with respect to aragonite will be achieved.
Did the authors calculate surface water pCO2 values for their 1956-57 data set? Is that 
consistent with wintertime pCO2 values for the region based on their work and other 
results for the Gulf of Alaska in winter?

Page 4593. Discussion. The authors state that “Few surveys of carbonate parameters 
have been made in the Gulf of Alaska which limits our understanding of its state and 
rate of change.” Is it not appropriate then for the authors to compare their results with 
the “few” previous studies that have addressed this problem for this region of the North 
Pacific? Why were the Feely et al., 1982, 1984, 1988; Chen et al., 1988 papers not 
discussed in this paper?

Page 4594. The authors state that “Along the North American upwelling region, Feely 
et al. (2008) found Òarag<1 at pH 7.75 in waters with potential densities of 26.2 at 
relatively shallow depths (40–120 m). In the winter of 1957, this combination of pH 
and potential density was found near the surface (<30 m) in the northwestern Gulf 
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of Alaska. At depths below 30 m, a pH of 7.75 on this potential density layer was 
characteristic of winter pH with average values of 7.65 (Fig. 4). In the summer of 1956, 
42 of 44 pH values were >7.75 on this density layer. Considering only the samples 
collected south of 51 N in summer, the minimum pH observed on the 26.15–26.25 layer 
at all stations was 7.83 (n=23, mean pH=8.00; mean depth of layer=136 m). These 
results suggest that the pH measured on samples taken in July 2007 would have been 
observed very rarely, if at all, during the summer of 1956.” If the water that is upwelled 
on the coast of North America does not come from the same source as the water 
on the same density surface in the Gulf of Alaska, I am not sure that this comparison 
is valid. I would think that the two water masses with different sources could have 
distinctly different pH, DIC and TA values before even considering the anthropogenic 
CO2 additions over time.

Figure 6 and in the text. What are the uncertainties of the aragonite saturation esti-
mates calculated by the procedures employed in this paper?

References cited
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influencing the degree of saturation of the surface and intermediate waters of the North 

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