

Global Assessment of Vegetation Index & Phenology Lab (VIP) and Global Inventory Modeling and Mapping Studies (GIMMS) Version 3 Products

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1 Abstract

2 Earth observation based long-term global vegetation index products are used by scientists
3 from a wide range of disciplines concerned with global change. Inter-comparison studies are
4 commonly performed to keep the user community informed on the consistency and accuracy of
5 such records as they evolve. In this study, we compared two new records: 1) Global Inventory
6 Modeling and Mapping Studies (GIMMS) Normalized Difference Vegetation Index Version 3
7 (NDVI3g) and 2) Vegetation Index & Phenology Lab (VIP) Version 3 NDVI (NDVI3v) and
8 Enhanced Vegetation Index 2 (EVI3v). We evaluated the two records via three experiments that
9 addressed the primary use of such records in global change research: 1) ~~prediction of the~~ Leaf
10 Area Index (LAI) ~~used in light use efficiency modeling~~; 2) ~~estimation of~~ vegetation climatology
11 ~~in Soil-Vegetation-Atmosphere-Transfer models~~; and 3) trend analysis of the magnitude and
12 ~~phenology timing~~ of vegetation productivity. ~~Experiment one, unlike~~ Unlike previous ~~inter-~~
13 ~~comparison global~~ studies, ~~was performed with~~ a unique Landsat 30 m spatial resolution and *in*
14 *situ* LAI database for major crop types on five continents was used to evaluate the performance
15 of not only NDVI3g and NDVI3v, but EVI3v as well. The performance of NDVI3v and EVI3v
16 was worse than NDVI3g using the *in situ* data, which was attributed to the fusion of GIMMS and
17 MODIS data in the VIP record. EVI3v has potential to contribute biophysical information
18 beyond NDVI3g and NDVI3v to global change studies, but we caution its use due to the poor
19 performance of EVI3v in this study. Overall, ~~the two records showed a high level of agreement~~

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20 ~~both in direction and magnitude on a monthly basis, though VIP values were higher and more~~
21 ~~variable and showed lower correlations and higher error with *in situ* LAI. The the records were~~
22 most consistent at northern latitudes during the primary growing season and southern latitudes
23 and the tropics throughout much of the year, while the records were less consistent at northern
24 latitudes during green-up and senescence and in the great deserts of the world throughout much
25 of the year. These patterns led to general agreement (disagreement) between trends in the
26 magnitude (timing) of NDVI over the study period. Bias in inter-calibration of the VIP record at
27 northernmost latitudes was suspected to contribute most to these discrepancies.- The two records
28 were also highly consistent in terms of trend direction/magnitude, showing a 30+ year increase
29 (decrease) in NDVI over much of the globe (tropical rainforests). The two records were less
30 consistent in terms of timing due to the poor correlation of the records during start and end of
31 growing season.

32
33 **Key words:** Normalized Difference Vegetation Index (NDVI); Leaf Area Index; Enhanced
34 Vegetation Index (EVI); remote sensing; agro-ecosystems

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35 1.0 Introduction

36 The Normalized Difference Vegetation Index (NDVI) (Rouse, 1974) is defined as $(\rho_{\text{NIR}} -$
37 $\rho_{\text{RED}})/(\rho_{\text{NIR}} + \rho_{\text{RED}})$, where ρ_{NIR} and ρ_{RED} are surface reflectance in the Near Infrared (NIR:
38 0.725–1.10 μm) and visible red (0.58–0.68 μm), respectively. As plants become more
39 photoactive, they absorb more visible red light due to the chlorophyll content of leaves and
40 stems, and scatter more in the Near Infrared due to the alignment of cell walls (Tucker et al.,
41 1994). This relationship, detected by remote sensing instruments at the canopy scale, has the
42 effect of making the index increase (decrease) as the density of the canopy increases (decreases)
43 (Tucker, 1979). As such, NDVI has been used widely in global change research with Earth
44 observation remote sensing for three general purposes: 1) the estimation of canopy properties
45 related to light-use efficiency, such as the Leaf Area Index (LAI) and Fraction of
46 Photosynthetically Active Radiation intercepted by the canopy (F_{PAR}) (e.g. Zhu et al. (2013)); 2)
47 representation of vegetation climatology in Soil-Vegetation-Atmosphere Transfer models (e.g.
48 O’ishi and Abe-Ouchi (2009)); and 3) detection of trends in vegetation (e.g. de Jong et al.
49 (2011)) and phenology (e.g. de Jong et al. (2012)). Several agro-ecosystem modeling
50 applications fall into these categories, including: agro-climate forecasting (Funk and Brown,
51 2006); drought monitoring (Karnieli et al., 2006); and crop yield estimation (Xin et al., 2013).
52 Although NDVI is widely used, it is sensitive to atmospheric effects, soil background, and
53 saturates at high LAI. The Enhanced Vegetation Index (EVI) was introduced to overcome these
54 limitations, as it includes a visible blue band to reduce atmospheric effects, calibration terms to
55 reduce the effects of soil background, and does not saturate as severely as NDVI at high LAI
56 (Huete et al., 2002). EVI has also been used in a wide array of global change studies, but post
57 2000, when the Moderate-Resolution Imaging Spectroradiometric (MODIS) satellite sensor

58 began retrieving visible blue reflectance (see Huete et al. (2010) for a review).

59 The Advanced Very High Resolution Radiometer (AVHRR) is the most commonly used
60 sensor for long-term (i.e. pre-MODIS) global change studies, because it began retrieving visible
61 red and NIR reflectance ~~needed to estimate NDVI from in 1981 and thus facilitates 30+ year~~
62 ~~time-series analyses of NDVI~~ (Brown et al., 2006). The AVHRR sensor has been on board eight
63 National Oceanic and Atmospheric Administration (NOAA) satellites: 7 (1981-1985), 9 (1985-
64 1988 and 1994-1995 descending), 11 (1988-1994), 14 (1995-2000), 16 (2000-2003), 17 (2003-
65 2009), 18 (2005-present), and 19 (2009-present). Reflectance data collected from the earlier
66 AVHRR sensors (7, 9, 11, and 14) were difficult to process and synthesize, because they lacked
67 onboard calibration; the NIR channel was sensitive to water, sun glint, glaciers at high latitudes,
68 and clouds; and of orbital drift (Rao and Chen, 1995, 1996). These issues were rectified with the
69 launch of the AVHRR sensors onboard NOAA 16, 17, 18, and 19, but have resulted in
70 radiometric and spectral inconsistencies across sensors that can significantly bias global change
71 analyses (van Leeuwen et al., 2006). Various methods have been developed to make these data
72 continuous and consistent through time, but take different approaches and are frequently
73 updated, necessitating new accuracy assessments to inform the user community as they evolve.

74 The Global Inventory Modeling and Mapping Studies (GIMMS: Tucker et al. (1994)) and
75 Vegetation Index & Phenology Lab (VIP: Didan (2014)) AVHRR products are actively used and
76 frequently updated, but represent fundamentally different approaches to synthesis. The NOAA
77 Global Vegetation Index (Jiang et al., 2010) is a category onto itself, ~~but since it~~ is stationary and
78 therefore not appropriate for change detection. Both GIMMS and VIP are aggregated to a 15-
79 day time step from daily data and are calibrated with higher spatial resolution sensors in the
80 period that overlaps NOAA 7, 9, 11, and 14 and NOAA 16, 17, 18, and 19. However before

81 aggregation, the former undergoes minor radiometric and spectral correction, while the later
82 undergoes rigorous atmospheric correction. Perhaps most importantly, GIMMS is developed
83 solely from AVHRR, while VIP is a blend of the AVHRR 1981-1999 Long-Term Data Record
84 (Nagol et al., 2009; Pedelty et al., 2007) and MODIS 2000-present. Finally, the VIP product
85 includes EVI2 (Jiang et al., 2008), which is a red-NIR version of EVI that has not been widely
86 evaluated and can potentially provide additional biophysical information and improve the
87 accuracy of long-term global change analyses (Rocha and Shaver, 2009). Given these
88 differences, studies have been performed at the global (Beck et al., 2011) and regional (Scheffic
89 et al., 2014) scale to assess the performance of older product versions, ~~while e.~~ Only one recent
90 study compared the latest product versions analyzed in this study globally, but only for the
91 consistency of trends (Tian et al., 2015). There ~~remains~~ no general consensus on which
92 product is superior; however, GIMMS NDVI tends to ~~perform more consistently temporally than~~
93 ~~VIP NDVI, making it be more appropriate than VIP NDVI~~ appropriate for trend analysis, because
94 the combination of poor orbital drift correction and blending between LTDR and MODIS
95 potentially contributes to large interseasonal variations in VIP NDVI, ~~while,~~ VIP NDVI, on the
96 other hand, may be more appropriate for estimating phenology (start of season, length of season,
97 and timing of peak NDVI) and other applications that require absolute NDVI values. In each
98 case, the performance of EVI2 was not evaluated nor was *in situ* data used for intercomparison.

99 The aim of this study was to perform a global assessment of the latest version of GIMMS
100 and VIP over a 30-year period (January 1982 to December 2011) in order to aid the user (global
101 change) community in interpreting results that involve these data. In doing so, we helped resolve
102 the superiority of one product over another. The assessment was performed with three
103 experiments that address the three major themes of global change research that involve Earth

104 observation remote sensing ~~previously introduced~~. Unlike other intercomparison studies, we
105 evaluated EVI2 and used an agro-ecosystem database comprised of relatively high spatial
106 resolution Landsat and *in situ* LAI sample pairs to assess the performance of each product for
107 agro-ecosystems in absolute terms. In addition, unlike other studies, the trend analysis was
108 performed not only on the magnitude of change across the globe on an annual basis, but the
109 change in the timing of NDVI according to the unique phenology in each hemisphere.

110 **2.0 Data, processing, and analytical methods**

111 **2.1 Global Inventory Modeling and Mapping Studies (GIMMS) Normalized Difference** 112 **Vegetation Index Version 3 (NDVI3g)**

113 The GIMMS vegetation index record evaluated is version three, which is labelled as
114 NDVI3g for the remainder of the paper. Full details on the product version can be found in
115 Pinzon and Tucker (2014). The new product includes a series of updates since the original
116 GIMMS NDVI and second generation NDVIg (Tucker et al., 2005) products. Like NDVIg, it is
117 a non-stationary NDVI series at 15-day intervals and $1/12^\circ$ (~8km at the equator) resolution;
118 corrected for orbital drift, Rayleigh scattering, and radiometric and spectral inconsistencies over
119 deserts; and takes an empirical (Bayesian) approach to normalize overlapping AVHRR periods
120 with another higher resolution sensor that overlaps the two periods. In addition, daily NDVI data
121 are scaled to 15-day composites using a Maximum Value Compositing (MVC) algorithm
122 (Holben, 1986), which reduces further inconsistencies in the daily data. The most unique
123 development in NDVI3g is the use of Sea-viewing Wide Field-of-view Sensor (SeaWiFS) for
124 intercalibration instead of the System Pour l'Observation de la Terre (SPOT) sensors. This is
125 intended to reduce significant bias in NDVI at extreme northern latitudes that has been observed
126 in SPOT imagery (Guay et al., 2014).

127 **2.2 Vegetation Index & Phenology Lab Version 3 Normalized Difference Vegetation Index**
128 **(NDVI3v) and Enhanced Vegetation Index 2 (EVI3v)**

129 The VIP vegetation index record evaluated is also in its third version, which is labelled as
130 NDVI3v and EVI3v for NDVI and EVI2 data, respectively, for the remainder of the paper.
131 Further information on the product version can be found in Didan (2014). Like previous
132 versions, it is a non-stationary series at 15-day intervals and 1/20° (~5km at the equator)
133 resolution; corrected using radiometric, drift, and cloud screening procedures recommended in El
134 Saleous et al. (2000), and an atmospheric algorithm that reduces the effects of Rayleigh
135 scattering, ozone, aerosols, and water vapor (Vermote et al., 1997); and takes an empirical (linear
136 regression by land cover type) approach for intercalibration. Unlike GIMMS, SPOT is used for
137 intercalibration and daily data are aggregated to 15-day composites using the Constrained View
138 angle - Maximum Value Composite (CV-MVC) approach (Cihlar et al., 1997). Unlike MVC,
139 CV-MVC does not give preference to off-nadir values that may be higher than “true” (at-nadir)
140 values. Version three includes one notable improvement over version two, namely the correction
141 of NDVI and EVI2 for sparsely vegetated areas pre-MODIS era (Scheffic et al., 2014). EVI2 is
142 derived from the following equation and responds similarly to EVI (Jiang et al., 2008):

$$E = 2.5 \frac{\rho_N - \rho_R}{\rho_N + 2.4\rho_R + 1} \quad (1)$$

143 The VIP product contained persistent data gaps due to cloud cover and other noise data
144 and was at a higher spatial resolution than the GIMMS product, so additional steps were taken to
145 process it before the assessment. A MODIS filtering algorithm described in Xiao et al. (2003),
146 Fensholt et al. (2006), and adapted for the tropics in Opiyo et al. (2013) was used to fill ~~some~~
147 data flagged as less than ideal gaps. Data gaps due to cloud cover and poor data quality were not
148 gap-filled. The algorithm was considered a compromise between preserving the actual data as

149 much as possible and filling missing data so that a reasonable comparison could be made.
150 ~~Statistical smoothing could have been used to fill the remaining data gaps, but was not used,~~
151 ~~because it would have risked comparing GIMMS data to a smoother and not actual VIP data.~~
152 **Figure 1** shows the percentage of missing data filled by the filtering algorithm. On a monthly
153 basis, less than 20% of the data was filled for the majority of pixels. Notable exceptions were
154 primarily in the mid and extreme latitudes during wintertime. The most severe case was in south
155 Asia during the monsoon (June – September) where more than 50% of the pixels were filled by
156 the filtering algorithm. After the filter was applied, NDVI3g was resampled to NDVI3v/EVI3v
157 resolution using the gdalwarp utility (<http://www.gdal.org/gdalwarp.html>) with default
158 parameters. Missing values were then made consistent across ~~the datasets~~ GIMMS and VIP, so
159 that the summary statistics (experiment two below) and trends (experiment three below) were
160 captured only for the 15-day values that the two products shared. The datasets were then
161 resampled back to the native NDVI3g spatial resolution for the evaluation. ~~These steps were~~
162 ~~taken to produce more reliable statistics and trends.~~

163 **2.3 First experiment: evaluation of NDVI3g, NDVI3v, and EVI3v with biophysical data**

164 NDVI and EVI are most commonly used in global change studies to capture F_{PAR} , which
165 drives canopy and light interactions in SVATs and other process-based models that estimate
166 plant productivity and evapotranspiration (Glenn et al., 2008). Monsi and Saeki (1953) found
167 that light attenuation in the canopy followed Beer’s Law (Beer, 1852). This means that for a
168 random canopy with a spherical leaf angle distribution, LAI, the second most commonly derived
169 biophysical parameter from NDVI and EVI, can be approximated from F_{PAR} using the following
170 equation (Norman et al., 1995):

$$L = \frac{-\ln(1 - F_{PAR})}{k} \quad (2)$$

171 Where k is an extinction coefficient and LAI is the Leaf Area Index ($m^2 m^{-2}$). Given the
 172 importance of NDVI and EVI in estimating F_{PAR} and LAI, standard regression techniques were
 173 used to measure the relative ability of NDVI3g, NDVI3v, and EVI3v to capture *in situ* LAI
 174 variability. It is difficult to compare these records to *in situ* LAI directly, because the NDVI/
 175 EVI - LAI relationship is typically scale dependent or non-linear (Friedl et al., 1995; Gao et al.,
 176 2000; Hall et al., 1992; Huete et al., 2005). Therefore F_{PAR} derived from Landsat Thematic
 177 Mapper/The Enhanced Thematic Mapper Plus (TM/ETM+) 30 m resolution surface reflectance
 178 data was used intermediately to downscale NDVI3g, NDVI3v, and EVI3v to 30 m resolution to
 179 facilitate the comparison.

180 **2.3.1 Landsat Thematic Mapper/The Enhanced Thematic Mapper Plus (TM/ETM+) and *in*** 181 ***situ* Leaf Area Index (LAI)**

182 The Landsat TM/ETM+ surface reflectance and *in situ* LAI data was extracted from a
 183 database that was developed to determine the ability of Landsat-based NDVI, EVI2, and other
 184 vegetation indices to predict LAI for field crops around the world. Results of the analysis, along
 185 with a full description of the database can be found in Kang et al. (2015). **Figure 2** shows the
 186 distribution of the Landsat-LAI sample pairs in the database. It includes nine major global field
 187 crops (barley, cotton, maize, pasture, potato, rice, soybean, sugar beet, and wheat) and several
 188 less common fields crops classified as "other" for purposes of this analysis. The *in situ* LAI was
 189 determined using ground-based optical (LAI 2000, AccuPar, and hemispherical) and destructive
 190 techniques and compiled from a number of sources. These include: AmeriFlux
 191 (<http://ameriflux.ornl.gov/>) and AsiaFlux (<http://asiaflux.net/>) regional flux networks;
 192 experimental and validation projects (e.g. Marshall and Thenkabail (2015)); the VALidation of

193 European Remote sensing Instruments project (Baret et al., 2014); the Australian Airborne
194 Cal/val Experiments for SMOS project (Peischl et al., 2012); as well as data retrieved from peer-
195 reviewed journals. For each LAI record in the database, Landsat TM/ETM+ radiance was
196 extracted from the United States Geological Survey archive within a ± 15 -day window
197 encompassing the date of *in situ* measurement and converted to surface reflectance with the
198 Landsat Ecosystem Disturbance Adaptive Processing System (Masek et al., 2006). NDVI and
199 EVI2 were computed using the equations above. In rare cases where more than one LAI
200 observation fell in a single Landsat pixel, the LAI values were averaged, so that each *in situ*
201 entry corresponded to a unique Landsat NDVI/EVI2 value. After averaging, the dataset
202 consisted of 2086 LAI-Landsat pairs, which was subsequently reduced to 1459 measurements
203 after further quality control measures [described in Kang et al. \(2015\)](#) were taken to remove
204 inconsistent samples.

205 **2.3.2 Downscaling long-term records with the Fraction Photosynthetically Active Radiation** 206 **intercepted by the canopy (F_{PAR}) and evaluation with *in situ* Leaf Area Index (LAI) data**

207 Downscaling was performed by converting AVHRR and Landsat vegetation indices to
208 F_{PAR} . Unlike the NDVI/ EVI - LAI relationship, the NDVI/EVI - F_{PAR} relationship is quasi scale
209 invariant (Asrar et al., 1992; Friedl et al., 1995; Gutman and Ignatov, 1998; Myneni et al., 2002;
210 Sellers, 1985), meaning a coarse resolution F_{PAR} pixel is approximately equal to the average of
211 overlapping higher resolution F_{PAR} pixels. ~~Hwang et al. (2011), for example, used the quality of
212 scale invariance between NDVI and F_{PAR} to downscale MODIS (1 km and 250 m spatial
213 resolution) NDVI to Landsat spatial resolution NDVI. Since they had access to multiple MODIS
214 and Landsat pixels through time and the linear relationship is land cover dependent, MODIS was
215 downscaled by multiplying each pixel by a ratio of Landsat to MODIS F_{PAR} .~~ In this study, on a

216 per pixel basis, most of the *in situ* LAI was retrieved only once, so using a ratio-based approach
217 was not feasible. Therefore, the AVHRR vegetation indices were downscaled to 30 m spatial
218 resolution by regressing (linearly) Landsat F_{PAR} and NDVI3g, NDIV3v, and EVI3v F_{PAR} . In
219 order to reduce the impact of land cover dependence, the models were developed for each crop.

220 The Fraction of Photosynthetically Active Radiation intercepted by the canopy was
221 computed using the ratio method first proposed in Gutman and Ignatov (1998):

$$F_p = \frac{V - V_{min}}{V_{max} - V_{min}} \quad (4)$$

222 Where VI_{min} is the vegetation index (NDVI or EVI2) for bare soil ($LAI = 0$), and VI_{max} is the
223 vegetation index (NDVI or EVI2) for dense vegetation ($LAI = 1$). VI_{min} and VI_{max} for NDVI
224 and EVI2 were set to 0.05 and 0.95 (Fisher et al., 2008; Mu et al., 2007). These limits are
225 sometimes considered dependent on the spatial and temporal resolution and land cover type
226 (Zeng et al., 2000). The limits proved arbitrary for downscaling purposes however, and using
227 the range 0.05 to 0.95 guaranteed that fractions ranged from zero to one.

228 Once NDVI3g, NDIV3v, and EVI3v F_{PAR} were downscaled to corresponding Landsat
229 data, their performance was evaluated by regressing them (linearly) with the *in situ* LAI data.
230 Since the relationship between F_{PAR} and LAI is logarithmic, as shown in **Equation 2**,
231 standardized residual plots (not shown) were made and linear transformations were performed to
232 verify that the assumptions of normality were met. In most cases, transformations were not
233 required. The performance of the final model selected in each case was characterized by the
234 coefficient of determination (R^2), significance tests, and root-mean-square error (RMSE).

235 Of the original 1459 Landsat – LAI data pairs, only 242 were used for the final analysis.
236 The majority of the data loss was due to considerable overlap of LAI data in space and time,
237 because they were collected without remote sensing applications in mind: 1) LAI values that

238 were captured by the same coarse resolution pixels were averaged along with Landsat
239 NDVI/EVI2 and 2) due to the presence of missing values in the long-term records. LAI and
240 Landsat NDVI/EVI2 were averaged on a 15-day basis. These reductions led to small sample
241 sizes for each crop. The sample sizes for cotton and rice were so small that they were omitted to
242 avoid over-fitting. In order to increase the sample size on a per-crop basis, two aggregations
243 based on the presumed similarity of crop spectral/canopy characteristics were made: 1) barley
244 and wheat (winter and spring varieties) were classified as wheat and 2) garlic, onion, potato, and
245 sugar beet were classified as tuber.

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246 **2.4 Second experiment: comparison of NDVI3g and NDVI3v climatology used to** 247 **parametrize SVAT models**

248 SVAT models traditionally were stand-alone and used to simulate the interaction of
249 incoming solar radiation with the canopy driven by F_{PAR} and ~~other biological and chemical~~
250 ~~canopies biogeochemical~~ processes for a single location, but are becoming increasingly coupled
251 to regional and global scale climate models and run over regularized grids, given the importance
252 of vegetation feedbacks on the atmosphere (Quillet et al., 2010). ~~With the exception of newer~~
253 ~~SVATs that include a dynamic vegetation component (see Scheiter et al. (2013) for a review);~~
254 ~~the vast majority of SVATs assume vegetation varies throughout the year without interannual~~
255 ~~variations.~~—A common dataset used to parameterize the F_{PAR} ~~vegetation~~ component of SVATs is
256 the 0.15° resolution monthly climatology of F_{PAR} ~~derived from~~ AVHRR NDVI (Gutman and
257 Ignatov, 1998). Given the importance of the F_{PAR} climatology ~~NDVI in representing vegetation~~
258 in SVATs, long-term summary statistics for NDVI3g and NDVI3v were computed as part of the
259 assessment. EVI3v was not included in this ~~phase of the analysis~~ experiment, because it does not
260 have a GIMMS counterpart to compare it to, has different and well-documented statistical

261 properties than NDVI, and it is derived from the same visible red and NIR channels and
262 underwent the same corrections as NDVI3v making its comparison redundant. The summary
263 statistics were computed from the 15-day data, but the results are presented here on a monthly
264 basis to reflect the NDVI climatology used in SVATs. The summary statistics included: mean,
265 standard deviation, coefficient of determination (R^2) from linear regression, and slope from
266 linear regression. The mean and standard deviation statistics are most critical for understanding
267 the differences in NDVI climatology, while R^2 and slope indicate the strength, magnitude, and
268 direction of the correlation between the two datasets. All summary statistics are presented with
269 significance (p) < 0.05 . Non-linear correlation statistics were also computed, but were not
270 included, because they showed similar spatial patterns as the linear statistics.

271 **2.5 Third experiment: comparison of NDVI3g and NDVI3v trends in magnitude and timing** 272 **(phenology)**

273 Changes in the magnitude and timing (phenology) of plant productivity are important for
274 understanding how ecosystems respond to climate change (Nemani et al., 2003). In North
275 America, for example, trend analysis of these changes has revealed that global warming is
276 driving an increase in plant productivity and a lengthening of the growing season (i.e. earlier
277 green-up in the spring and later senescence in autumn) (Barichivich et al., 2013). The
278 characterization of the magnitude and phenology of productivity over a year is typically
279 estimated with empirical methods that include NDVI and other bioclimatic predictors such as
280 temperature and relative humidity (e.g. Brown et al. (2012)). In order to avoid confounding the
281 assessment of GIMMS and VIP with other variables, harmonic regression (Eastman et al., 2009;
282 Jakubauskas et al., 2001) was performed on the vegetation index records to measure the
283 magnitude and timing of NDVI on an annual basis. As with experiment two, EVI3v was not

284 evaluated in this experiment. A trend analysis was then performed on the regression parameters
285 to compare NDVI3g and NDVI3v as surrogates for the change in magnitude and timing of plant
286 productivity over time.

287 The primary parameters of harmonic regression are the amplitude (in this case the
288 difference between peak and mean NDVI) and phase (in this case timing of NDVI peaks and
289 troughs). Amplitude and phase are computed by fitting a series of sinusoidal functions to the
290 time series (Eq. 3). The harmonic regression was performed on a monthly basis for each year.
291 Monthly values were determined by taking the maximum NDVI of the two 15-day composites
292 per month.

$$NDVI_t = NDVI_0 + \sum_{i=1}^j A_i \cos\left(\frac{2\pi}{N} t\right) + B_i \sin\left(\frac{2\pi}{N} t\right) \quad (3)$$

293 Where $NDVI_t$ is the predicted Normalized Difference Vegetation Index at month (t), $NDVI_0$ is
294 the annual monthly mean, i is the number of harmonics up to the jth harmonic, N is the number
295 of samples (months) in the year, and A and B are coefficients used to compute the amplitude and
296 phase. The regression was performed for the first harmonic, which represents the primary
297 growing season, because multimodal systems (harmonics > 1) are uncommon and capturing
298 them risks over-fitting.

299 The change in amplitude and phase over time was quantified using the Theil-Sen
300 technique (Gilbert, 1987). The Theil-Sen technique takes the median non-parametric slope over
301 all possible pairwise slopes through time. Unlike linear regression, it does not require normality
302 or homoscedasticity, making it appropriate for trend analyses involving NDVI data (de Beurs
303 and Henebry, 2005). The significance of the amplitude and phase trends ($p < 0.05$) was
304 identified using the non-parametric Mann-Kendall test. Since the primary growing season in the

305 southern hemisphere occurs over two given calendar years, the trend analysis was repeated for
306 the southern hemisphere by advancing the regression six months ahead each year. This resulted
307 in one less year or a 29-year trend analysis for the southern hemisphere.

308 3.0 Results

309 3.1 First Experiment: performance of long-term records using Landsat F_{PAR} and *in situ*

310 LAI

311 ~~Of the original 1459 Landsat – LAI data pairs, only 242 were used for the final analysis.~~
312 ~~A small portion of the data loss was due to the fact that they were collected after the long-term~~
313 ~~records ended. Most of the data loss was due to considerable overlap of LAI data in space and~~
314 ~~time, because they were collected without remote sensing applications in mind: 1) LAI values~~
315 ~~that were captured by the same coarse resolution pixels were averaged along with Landsat~~
316 ~~NDVI/EVI2 and 2) due to the presence of missing values in the long-term records, LAI and~~
317 ~~Landsat NDVI/EVI2 were averaged on a 15-day basis. These reductions led to small sample~~
318 ~~sizes for each crop. The sample sizes for cotton and rice were so small that they were omitted to~~
319 ~~avoid over-fitting. In order to increase the sample size on a per-crop basis, two aggregations~~
320 ~~based on the presumed similarity of crop spectral/canopy characteristics were made: 1) barley~~
321 ~~and wheat (winter and spring varieties) were classified as wheat and 2) garlic, onion, potato, and~~
322 ~~sugar beet were classified as tuber.~~

323 The accuracy of each long-term record when compared to *in situ* LAI was mixed, but
324 NDVI3g performed moderately better than NDVI3v and EVI3v. The scatterplots of predicted
325 (downscaled) NDVI3g, NDVI3v, and EVI3v F_{PAR} versus Landsat F_{PAR} for wheat and pasture are
326 shown in **Figure 3**, while the summary statistics of the linear models used to downscale the
327 records for all crops with sufficient samples sizes and reasonable correlations are shown in **Table**

328 **1.** The models used to downscale NDVI3g yielded higher correlations and lower error than the
329 models used to downscale NDVI3v for maize and wheat, while NDVI3v yielded higher
330 correlations and lower error for soybean and pasture, and EVI3v was the most difficult to
331 downscale of the three. Specifically, R^2 for NDVI3g over NDVI3v was 0.04 for maize and
332 0.18 for wheat, while R^2 for NDVI3v over NDVI3g was 0.06 and 0.04 for pasture and soybean.
333 It is important to note however that the strength of the relationships were low across all records
334 with the exception of pasture, which could be due to the homogeneity (consistent clumping) of
335 pasture over large areas. The relationship for tuber was so poor that it was not included in the
336 LAI evaluation. The relationship between the downscaled NDVI3g, NDVI3v, and EVI3v F_{PAR}
337 and *in situ* LAI are shown for wheat and pasture in **Figure 4**, while the model statistics and
338 transformation for a linear comparison, are presented in **Table 2**. The NDVI3g-LAI models
339 captured *in situ* variability better than NDVI3v and EVI3v for maize ($R^2 = 0.06$), pasture ($R^2 =$
340 0.11), and wheat ($R^2 = 0.10$), with comparable results between NDVI3g and NDVI3v for
341 soybean. EVI3v tended to perform better than NDVI3v for two of the crops: pasture ($R^2 =$
342 0.05) and wheat ($R^2 = 0.04$). As can be seen in **Figure 4**, however, the predictive power of
343 EVI3v could be inflated by leveraging at high LAI, i.e. EVI3v tends to be more variable than
344 NDVI3v at higher LAI.

345 **3.2 Second experiment: similarity of NDVI3g and NDVI3v climatology**

346 On a monthly basis, NDVI3g and NDVI3v showed a high level of consistency in terms of
347 relative magnitude expressed as R^2 (**Figure 5**) and direction expressed as slope (**Figure 6**). Both
348 metrics were computed with the slopes forced through the origin (0, 0). In the northern
349 hemisphere, R^2 approached one after green-up (May) and progressively got stronger over the
350 boreal summer months (June, July, and August). The poorest correlations ($R^2 < 0.7$) were seen

351 primarily at the northern-most latitudes during the transition between boreal winter and spring.
352 Correlations were more consistent in the Southern Hemisphere where snow and cloud cover was
353 notably less than in the north. A glaring exception however was the Strut Stony Desert of South
354 Central Australia, which showed poor correlations during the transition between Austral summer
355 (December, January, and February) and fall. The tropics showed high and significant
356 correlations throughout most of the year as well. The slopes followed a similar pattern as the
357 correlations, with values approaching a one-to-one relationship (slope=1.0) after the transition
358 from winter to spring in the northern hemisphere and consistently over much of the year in the
359 tropics and southern hemisphere. The great deserts of the world and sparsely vegetated areas had
360 slopes approaching zero throughout the year. Since the slopes were expressed with NDVI3v as
361 the dependent variable and the slopes were always less than one, NDVI3g was always less than
362 NDVI3v. The difference in NDVI3g and NDVI3v magnitudes is more clearly shown in **Figure**
363 **7**, which illustrates the monthly latitudinal mean and standard deviation for both. Mean NDVI3v
364 was always higher and more variable than NDVI3g. In addition, large divergence in means
365 between the two records occurred during senescence in the northern hemisphere. Other patterns
366 were more consistent: NDVI3g and NDVI3v were high in the tropics throughout the year and
367 peak and decline following the seasons in the northern and southern hemispheres; and the
368 standard deviations for both were higher in the northern hemisphere than the southern
369 hemisphere due to continentally.

370 **3.2.3 Third experiment: similarity of NDVI3g and NDVI3v trends in magnitude and** 371 **phenology**

372 The two NDVI records exhibited a high level of correspondence in maximum primary
373 season NDVI (1st harmonic amplitude), both in direction and location (**Figure 8**). In terms of

374 magnitude trends, however, NDVI3v was higher than NDVI3g. The figure was masked for
375 pixels that had complete NDVI records to ~~guarantee accurate~~ facilitate curve-fitting in a given
376 year and then again for trends that were statistically significant over the 30-year period. This
377 resulted in no trends over much of the northern latitudes. In addition, NDVI amplitudes 0.03
378 per year (or 1.0 over the 30-year period) and NDVI amplitudes -0.03 (or -1.0 over the 30-year
379 period) were flagged as missing, since NDVI ranges from -1 to 1 . In most cases, however, the
380 increase in absolute amplitude per year was less than 0.01 or 0.3 over the 30-year period.
381 Overall, the positive NDVI3g trends appeared to be more consistent spatially in several
382 important cropping and grazing regions, including: the Great Plains of the United States; the
383 Region del Norte Grande of Argentina; the Iberian Peninsula (particularly Portugal); Lesotho,
384 South Africa (east), and Swaziland; Ganges (India) and Indus (Pakistan) Plains; the Sahel of
385 West Africa; and Cape York Peninsula (Australia). Negative trends (also more consistent in
386 NDVI3g) appeared to be primarily in the great deserts of the northern hemisphere. In the
387 southern hemisphere, however, some negative trends were seen in the tropical forests of the
388 Amazon and Congo River basins.

389 The two records in terms of primary season timing (1^{st} harmonic phase) showed a lower
390 level of correspondence than for amplitude (**Figure 9**). As above, trends were not seen over
391 much of the northern hemisphere. In addition, the NDVI phases 0.07 per year (or ~ 2 months
392 over the 30-year period) and NDVI phases -0.07 (or ~ -2 months over the 30-year period) were
393 flagged as missing, because changes of more than two months were deemed aberrant. In most
394 cases however, the absolute change in timing was less than two months. As with trends in
395 amplitude, the trends in phase were more consistent spatially over both hemispheres from
396 NDVI3g. Earlier green-up (negative trend) represented the majority of trends in the two

397 datasets, though considerably less than the increase in amplitude shown in **Figure 8**. Negative
398 trends were seen over many important cropping and grazing areas: California and the
399 Southwestern United States; the Iberian Peninsula; the Sahel of sub-Saharan Africa; Iran (east);
400 South Africa (west); Turkmenistan (north); and over much of the areas bordering the deserts of
401 Australia. Later green-up (positive trend) was primarily concentrated in the great deserts (e.g.
402 the Great Sandy and Gibson deserts of northwestern Australia).

403 **4.0 Discussion**

404 This study assessed the latest versions of two non-stationary and long-term vegetation
405 index records used in global change studies. The assessment was performed with three
406 experiments that addressed ~~the primary important~~ global change applications, namely: ~~the~~
407 ~~estimation of~~ F_{PAR} and LAI; ~~estimation of SVAT~~ vegetation climatology; and trend analysis of
408 vegetation ~~productivity~~ magnitude and ~~phenology~~ timing. The results of the analysis highlight
409 important similarities and differences between the two records that the global change community
410 should be aware of before using them for these applications: 1) NDVI3v was consistently higher
411 and more variable than NDVI3g, which in Tian et al. (2015) has been attributed to artificial
412 jumps in the record between AVHRR and MODIS periods and may contribute to relatively lower
413 correlations and higher errors with *in situ* LAI; 2) the performance of EVI3v with *in situ* LAI
414 compared to NDVI3g was unexpectedly poor; 3) correlations between GIMMS and VIP were
415 highest during the primary growing season, so trends in peak NDVI were fairly consistent
416 between the two, both showing increases over much of the globe and decreases in tropical
417 rainforests; and 4) correlations between GIMMS and VIP were lower during green-up and
418 senescence, which were most pronounced at high latitudes where the NDVI3g product is
419 expected to have much lower bias due SeaWifs inter-calibration. so trends in NDVI timing were

420 ~~less consistent between the two, however, both showed earlier green-up over much of the globe,~~
421 ~~particularly in the driest regions of the world.~~Overall, we recommend using NDVI3g over
422 NDVI3v and EVI3v for vegetation climatology and trend analysis, because it is spatially and
423 temporally more consistent. Unlike previous studies, however, the *in situ* LAI experiment
424 revealed that NDVI3g is better suited for absolute measurements as well.

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425 4.1 First Experiment: performance of long-term records using Landsat FPAR and in situ LAI

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426 Unlike previous inter-comparison studies, a unique moderate resolution remote sensing
427 and *in situ* LAI database for agro-ecosystems was used for accuracy assessment. ~~Although there~~
428 ~~was a spatial mismatch between *in situ* and AVHRR data, and the *in situ* data had a small sample~~
429 ~~size with a limited geographic extent.~~In most cases, NDVI3g ~~appeared to be~~was more accurate
430 than NDVI3v or EVI3v. EVI3v performed considerably worse than NDVI3g, which is
431 surprising, because EVI tends to be better correlated than NDVI from other sensors with canopy
432 structural properties (Huete et al., 2002). Earlier studies have suggested that the LTDR NDVI
433 from which MODIS data is merged in the VIP product is more appropriate for modeling
434 applications requiring absolute values (Beck et al., 2011), meaning NDVI3v should reproduce
435 more accurate estimates of F_{PAR} and LAI than NDVI3g, but this was not the case in this study.
436 Tian et al. (2015) assessed ~~the blended and smoothed LTDR and MODIS product~~NDVI3v. They
437 attributed ~~d- jumps in the NDVI3v record~~ ~~the relatively high and variable NDVI3v mainly~~ to poor
438 orbital drift correction and the break in the LTDR and MODIS records in 2000. The reason for
439 the poor performance of EVI2 is less clear, but clearly needs to be addressed in future work,
440 given its potential importance to advancing global change research. ~~However, since the LTDR~~
441 ~~data appears to reproduce more accurate absolute values than GIMMS and a smoother was not~~
442 ~~used and there was a high level of correlation between NDVI3g and NDVI3v in this study,~~

443 ~~orbital drift correction is likely not the culprit. Therefore, the blending of MODIS and LTDR is~~
444 ~~most likely the most important factor impacting the accuracy of biophysical estimates in~~
445 ~~NDVI3v and EVI3v and should be addressed in later product versions.~~

446 ~~At the time of writing this manuscript, a VIP Version 4 is forthcoming. It will be~~
447 ~~interesting to see if this new version will produce more accurate results using the LAI Landsat~~
448 ~~database. In the meantime, however, if users require the higher spatial resolution offered by~~
449 ~~VIP and the added biophysical information afforded by EVI3v for application purposes, several~~
450 ~~options exist for improving their accuracy. Perhaps the most important would be to fill the~~
451 ~~remaining data gaps in the filtered VIP datasets generated here with a smoothed data (see~~
452 ~~Kandasamy et al. (2012) for examples), which will address some of the noise in the data~~
453 ~~observed in Tian et al. 2015 and this study. NDVI3g has undergone extensive statistically~~
454 ~~smoothing. Another option widely used in the climate modeling community, that could be~~
455 ~~combined with this option would be to generate an ensemble mean of NDVI3v and NDVI3g to~~
456 ~~account for some of the bias and uncertainties in each product. Finally, instead of using EVI3v,~~
457 ~~the red and NIR channels included in the VIP database could be used to calculate the Soil~~
458 ~~Adjusted Vegetation Index (SAVI) (Huete, 1988) instead. The evaluation of EVI2 has so far~~
459 ~~been limited, whereas Unlike EVI2, SAVI has undergone extensive evaluation.~~

460 ~~The LAI Landsat database should be combined with other databases in the future, such as~~
461 ~~the LAI for woody plant database (Lio et al., 2014), so that a large amount of data over multiple~~
462 ~~biomes are used to develop robust evaluations (Weiss et al., 2014). New databases should aim to~~
463 ~~extend the temporal ranges of biophysical data on a per pixel basis, so that the ratio-based~~
464 ~~approach to downscaling as suggested in Hwang et al. (2011) can be performed, instead of the~~
465 ~~linear regression by crop type approach taken here. The downscaling procedure can also be~~

466 ~~improved. In the Hwang et al. (2011) study, F_{PAR} was used to downscale MODIS data to~~
467 ~~Landsat resolution, representing a ratio of approximately 8 : 1 (250 m : 30 m), whereas in this~~
468 ~~study, Landsat F_{PAR} was used to downscale AVHRR data, representing a ratio of approximately~~
469 ~~266 : 1 (8000 m : 30 m). The large discrepancy in resolution in this study could be resolved in~~
470 ~~the future by first downscaling AVHRR with MODIS F_{APAR} and then downscaling again using~~
471 ~~Landsat F_{APAR} .~~

472 4.2 Second experiment: similarity of NDVI3g and NDVI3v climatology

473 NDVI3g and NDVI3v showed a high level of agreement with one another at mid-
474 latitudes during the primary growing season and in the densely vegetated tropics throughout
475 most of the year, and a low level of agreement at high latitudes during winter months and in the
476 sparsely vegetated sub-tropics throughout most of the year. The high level of agreement is
477 expected, because data gaps, cloud contamination, and atmospheric water vapor, is less at mid-
478 latitudes during summer months (Beck et al., 2011; Moulin et al., 1997). The high level of
479 agreement in the tropics was more surprising, because data gaps and cloud contamination are
480 persistent there throughout much of the year, typically leading to large discrepancies among
481 records (Brown et al., 2006). However, as previously stated, ~~the standard smoothed VIP data~~
482 ~~was not used this study, so many of the potentially smoothed and many contaminated pixels were~~
483 omitted from the analysis. The large discrepancy at high latitudes could have been due to factors
484 other than cloud contamination and other noise data gaps, including the 1) presence of snow
485 cover; 2) high frequency of off-nadir pixels, which would impact the results of the compositing
486 algorithm (MVC versus CV-MVC); and perhaps most importantly, 3) use of SeaWiFS over
487 SPOT for GIMMS inter-calibration (Hall et al., 2006). The large discrepancy in deserts and
488 sparsely vegetated areas on the other hand was most likely due to the dominance of soil in the

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489 signal and sensitivity of NDVI to soil wetness (Jiang et al., 2006). With the high level of
490 correlation during the primary growing season and higher and more variable NDVI3v, users
491 should expect NDVI3v climatology during the primary growing season to be higher at mid-
492 latitudes and in the tropics throughout most of the year, but consistent with changes in NDVI3g.
493 During winter months, especially at high latitudes and in semi-arid to arid subtropical regions,
494 where SeasWiFS inter-calibration is less biased, NDVI3v will be higher, more variable, and less
495 consistent with accurate than NDVI3g.

496 4.3 Third experiment: similarity of NDVI3g and NDVI3v trends in magnitude and timing

497 NDVI3g and NDVI3v both showed greening (positive NDVI amplitude) globally, with
498 localized browning (negative NDVI amplitude) over a 30+ year time frame, but the magnitude of
499 the trends in the latter was higher. Therefore, trend analyses of peak NDVI or annual means will
500 be higher in NDVI3v than NDVI3g, but the direction will be the same. The direction of change
501 in general corroborated previous global studies. The gain or loss of plant productivity is
502 generally attributed to biophysical drivers (temperature and precipitation), human-related
503 change, and discontinuities in the long-term record (de Jong et al., 2012). At mid-latitudes,
504 warming (cooling) at the beginning of the growing season can lead to greening (browning) in
505 areas where water supplies are ample. In North America east of the Great Plains, for example,
506 greening was observed in NDVI3g and NDVI3v, which has been attributed to temperature-
507 driven increases in plant productivity in previous studies (Wang et al., 2011). Increased rainfall
508 (droughts) proceeding or during the growing season can lead to greening (browning) particularly
509 in water-limited regions such as the Sahel. As shown here, the Sahel has experienced greening
510 over the past 30+ years. This greening, typically referred to as the “re-greening of the Sahel” is
511 defined in other studies as the increase in woody biomass (Brandt et al., 2015) that followed the

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512 recovery of rains in the 1990's after two decades of severe droughts driven by below normal sea
513 surface temperatures in the subtropical North Atlantic (Giannini et al., 2013). Deforestation is
514 perhaps the most ~~recognized~~ appreciated human driver of plant productivity. Browning in the
515 Amazon and Congo River basins, as was shown in this study, has been attributed to widespread
516 deforestation in previous studies (Hansen et al., 2010; Mayaux et al., 2013), though other drivers,
517 such as shift in Walker circulation potentially contribute to the loss as well (Zhou et al., 2014).
518 Greening was observed in the areas tropical rainforests as well, but this has been attributed in
519 previous studies to rapid regrowth after deforestation, the way VIs are composited, and the
520 methods by which trends are detected (Beck et al., 2011). Some of the trends disagree with
521 previous research and should be addressed in future studies. Most prominent were that no trend
522 was detected at extreme northern latitudes, though previous studies have shown summer
523 drought-driven declines in boreal forest productivity (Goetz et al., 2005), and positive trends
524 were detected for the Region del Norte Grande of Argentina, though previous studies have
525 shown negative trends attributed to the rapid encroachment of agriculture into subtropical forests
526 of the region (Paruelo et al., 2004).

527 NDVI3g and NDVIv both showed earlier green-up (negative NDVI phase) more than
528 later green-up (positive NDVI phase), but they were less consistent with one another compared
529 to trends in peak NDVI. NDVI3g and NDVI3v showed low correlations during green-up and
530 diverging climatology during senescence, which could lead to discrepancies in the timing of start
531 of season (SOS) and end of season (EOS). Global studies seldom analyze trends in vegetation
532 timing. On a regional basis, however, tThe findings appear to be less consistent with previous
533 studies with the timing trends in other studies. Over the majority of northern regions, for
534 example, ~~the start of season (SOS)~~ has been retreating as shown, however unlike this study,

535 | previous studies have shown that ~~the end of the season (EOS)~~ has been advancing. The
536 | combination of the two processes has led to a longer growing season attributed primarily to
537 | asymmetric and rising global temperatures. One of the limitations of the harmonic approach
538 | taken in this study is that it is rigid, i.e. it assumes that the time series oscillates at a regular
539 | interval over each year. In the future, a harmonic or other phenological model that accounts for
540 | SOS and EOS asymmetry may be more appropriate for accurate trend analysis.

541 | **5.0 Conclusion**

542 | This paper revealed important similarities and differences of two new long-term
543 | vegetation databases: Global Inventory Modeling and Mapping Studies Normalized Difference
544 | Vegetation Index Version 3 (NDVI3g) and 2) Vegetation Index & Phenology Lab Version 3
545 | NDVI (NDVI3v) and Enhanced Vegetation Index 2 (EVI3v). Overall, NDVI3g performed better
546 | and more consistently than NDVI3v and EVI3v in three experiments designed to evaluate the
547 | two products in absolute terms and changes in magnitude and timing. ~~when downscaled with~~
548 | Landsat 30 m resolution fraction of photosynthetically active radiation intercepted by the canopy
549 | and compared to *in situ* Leaf Area Index (LAI). ~~VIP processing and the approach taken to~~
550 | synthesize data streams contributed to higher and more variable values that adversely affected
551 | the predictive ability of the database. ~~VIP tended to be higher in magnitude, more variable, and~~
552 | less consistent in terms of trends, due primarily to the blending of two sensors with different
553 | attributes (AVHRR with MODIS). GIMMS, on the other hand only uses AVHRR. ~~However,~~
554 | ~~the~~The two databases showed a high level of consistency during the primary growing season,
555 | which contributed to similar changes in the relative magnitude and direction of plant productivity
556 | climatology and dynamics, which are critical to global change research. The two products were
557 | less consistent in timing, especially at the start and end of the primary growing seasons at high

558 | latitudes. It is suspected that these poor correlations are attributed to the higher resolution
559 | sensors each product uses for intercalibration. due in part to their poorer correlation at the start
560 | and end of growing season. New opportunities exist for improving the two products that can
561 | account for the discrepancies highlighted here. In the meantimeIn conclusion, it is suggested
562 | users requiring a long-term product to measure biophysical parameters, vegetation climatology,
563 | and trends in plant productivity magnitude and timing to use NDVI3g and to avoid using EVI3v.

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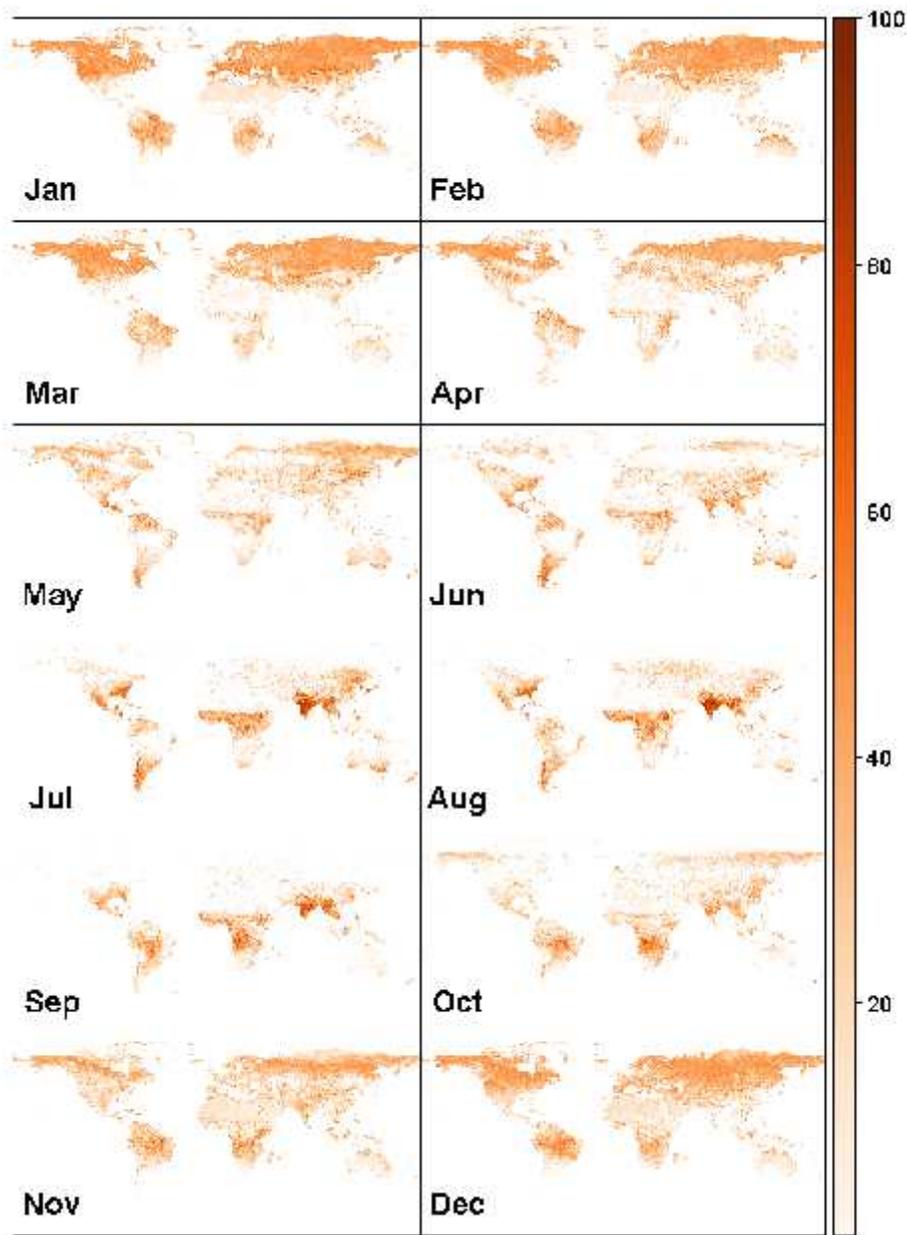


Figure 1. Percentage increase in pixels added (i.e. gaps filled) after applying the temporal filter to Vegetation Index & Phenology Lab Version 3 records.

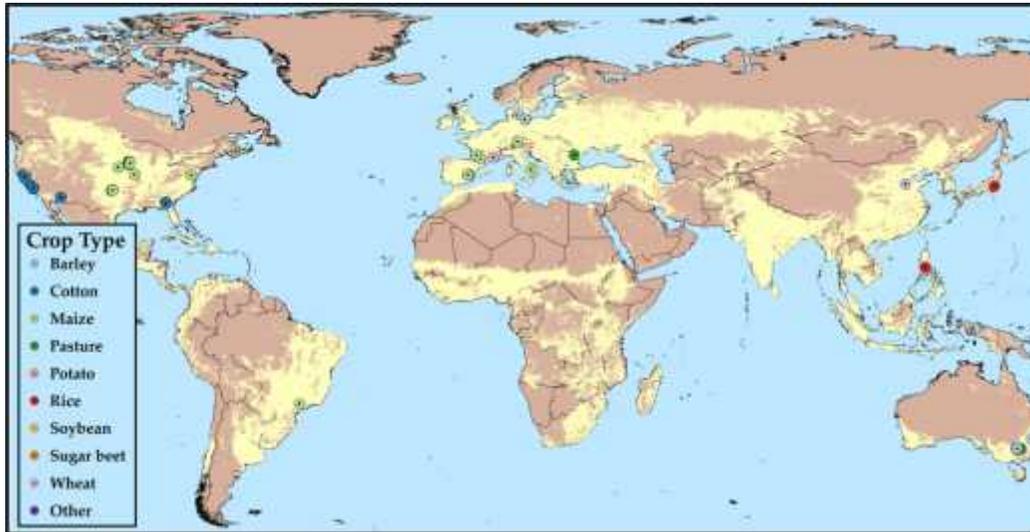


Figure 2. Sites where *in situ* (destructive or optical) measurements and Landsat Thematic Mapper/The Enhanced Thematic Mapper Plus ground reflectance data were compiled, resulting in more than 1,400 data pairs. The sites are overlaid with 1 km grid cells that contain 5% or more crop area (Ramankutty et al., 2008).

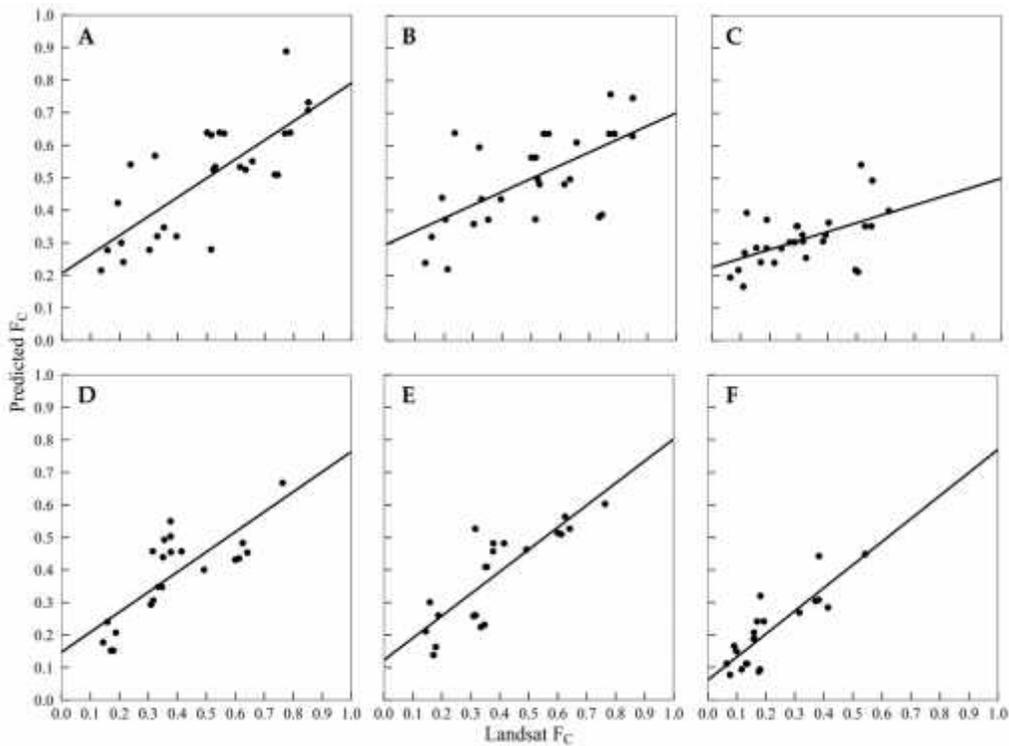


Figure 3. Scatterplots of the Fraction Absorbed of Photosynthetically Active Radiation (F_{APAR}) Landsat versus F_{APAR} for wheat (**a-c**) and pasture (**d-f**) estimated by the Global Inventory Modeling and Mapping Studies Normalized Difference Vegetation Index Version 3; Vegetation Index & Phenology Lab Version 3 Normalized Difference Vegetation Index; and Vegetation Index & Phenology Lab Version 3 Enhanced Vegetation Index 2, respectively. The solid lines represent the linear model used to downscale the vegetation record for evaluation with *in situ* leaf area index.

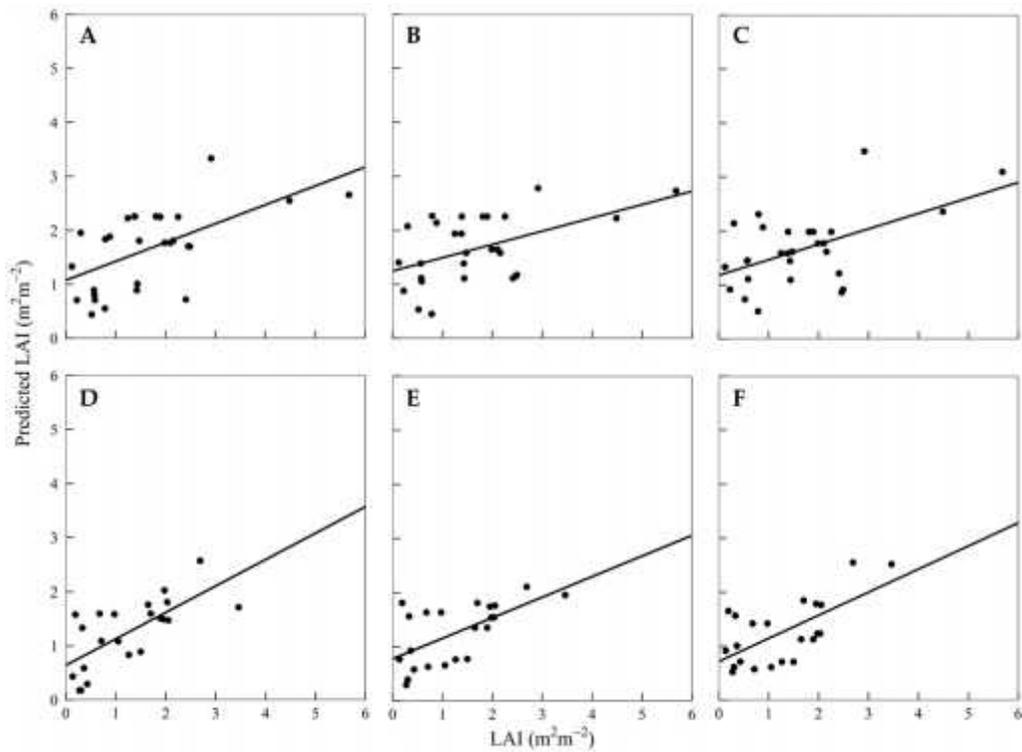


Figure 4. Scatterplots of *in situ* leaf area index for wheat (a-c) and pasture (d-f) versus corresponding Landsat resolution pixels downscaled from the Global Inventory Modeling and Mapping Studies Normalized Difference Vegetation Index Version 3; Vegetation Index & Phenology Lab Version 3 Normalized Difference Vegetation Index; and Vegetation Index & Phenology Lab Version 3 Enhanced Vegetation Index 2 datasets, respectively. The solid lines represent the best model fit.

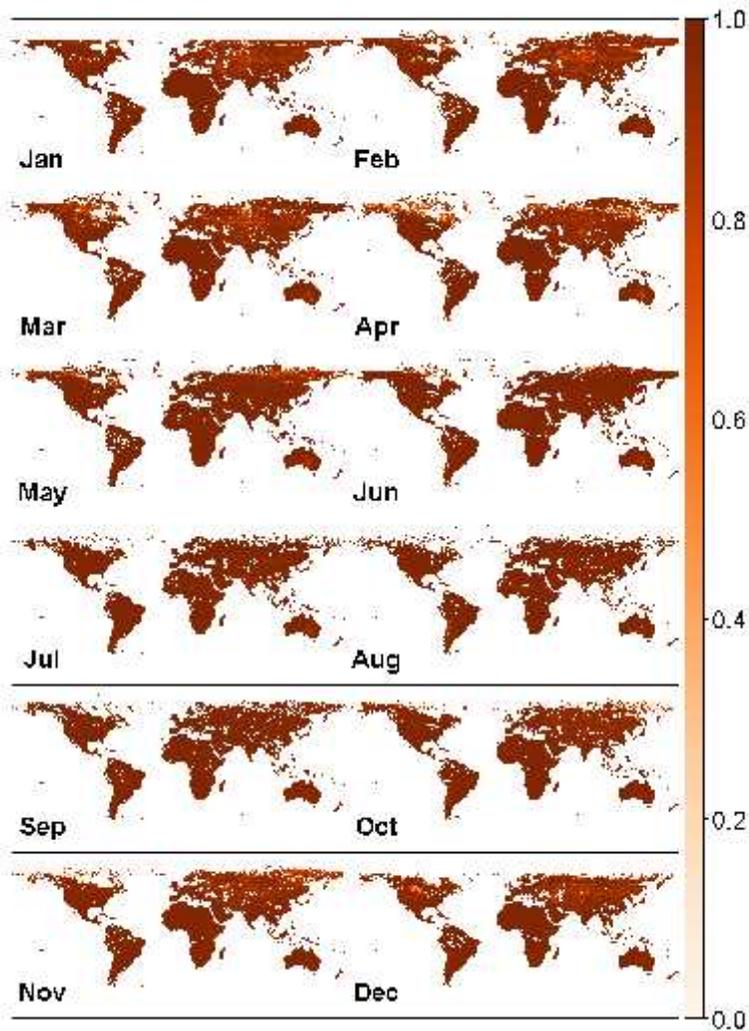


Figure 5. The coefficient of determination (R^2) on a per-pixel basis for the Vegetation Index & Phenology Lab Version 3 Normalized Difference Vegetation Index versus the Global Inventory Modeling and Mapping Studies Normalized Difference Vegetation Index Version 3. R^2 was determined using a 30-year time series of 15-day composites for each month. The images have been masked for significance > 0.05 and latitudes ranging from 60°N - 60°S .

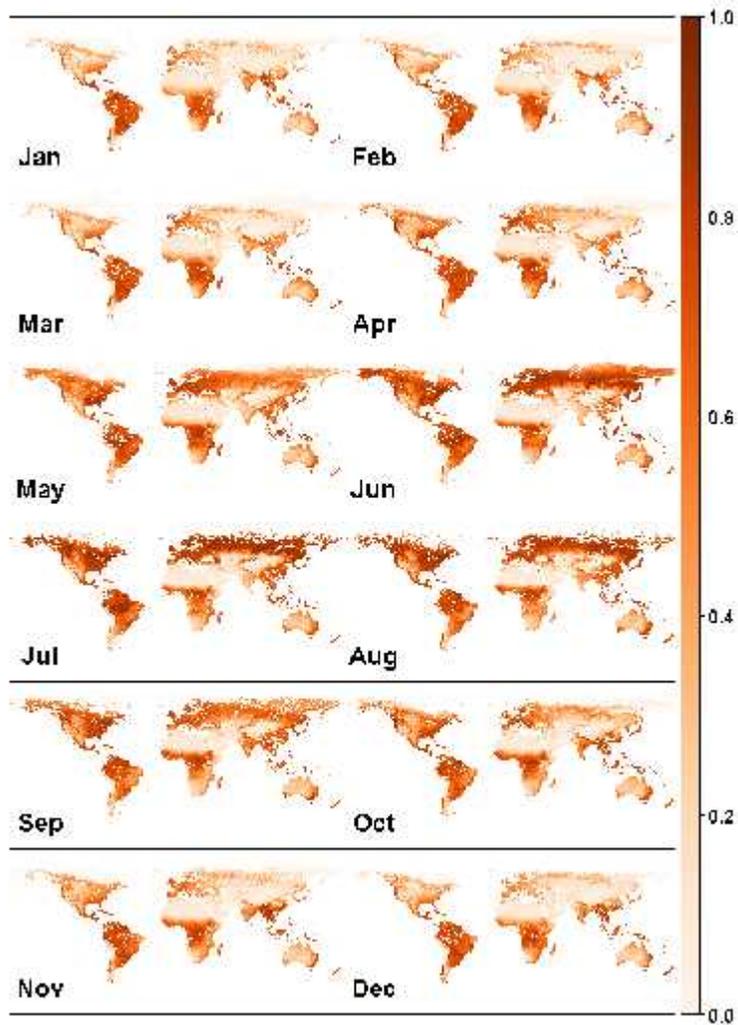


Figure 6. The slope (intercept = 0) determined from linear regression on a per-pixel basis for the Vegetation Index & Phenology Lab Version 3 Normalized Difference Vegetation Index versus the Global Inventory Modeling and Mapping Studies Normalized Difference Vegetation Index Version 3. Slope was determined using a 30-year time series of 15-day composites for each month. The images have been masked for significance ≥ 0.05 and latitudes ranging from 60°N - 60°S.

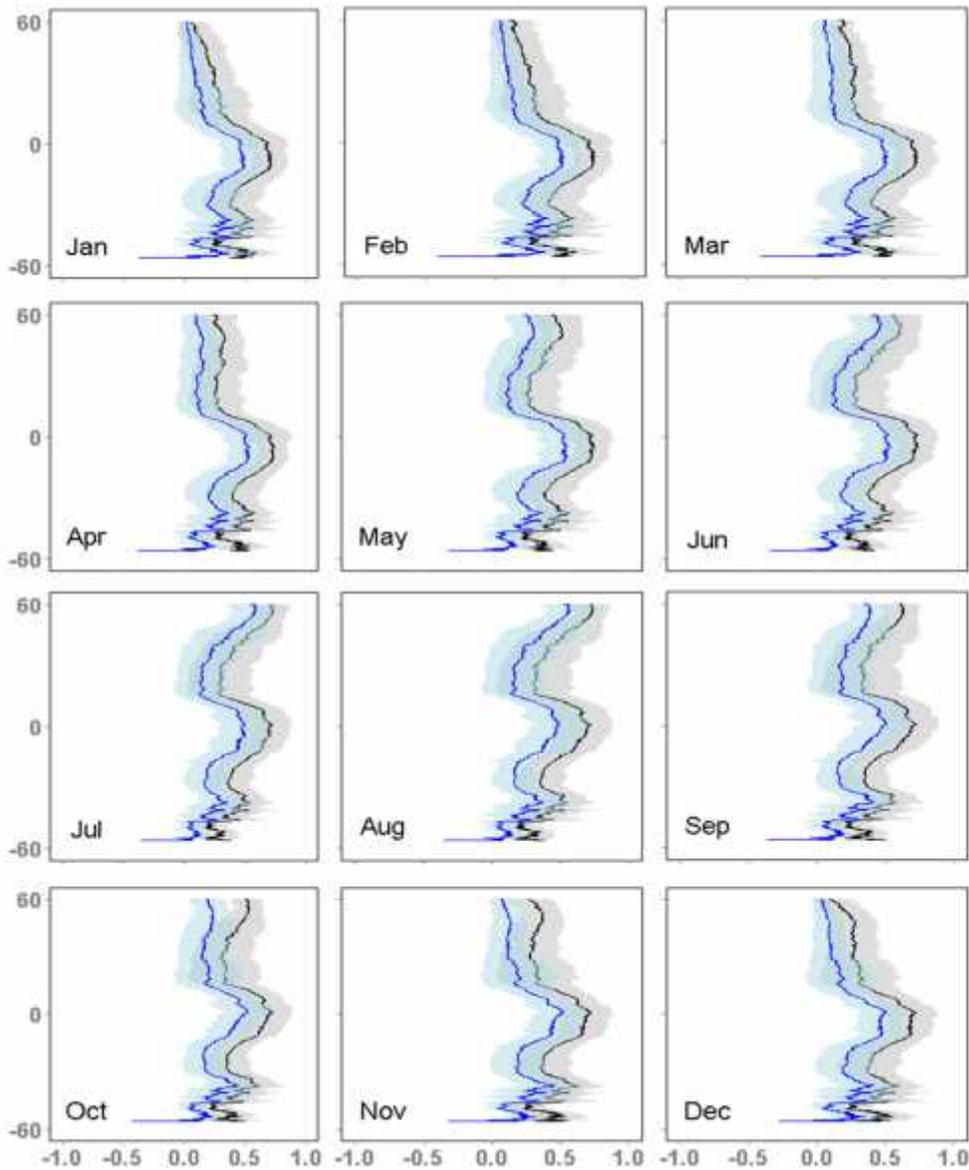


Figure 7. The latitudinal mean (solid line) and standard deviation (ribbon) of the Global Inventory Modeling and Mapping Studies Normalized Difference Vegetation Index Version 3 (blue) and Vegetation Index & Phenology Lab Version 3 Normalized Difference Vegetation Index (black) over 30 years. Values are shown from 60°N - 60°S.

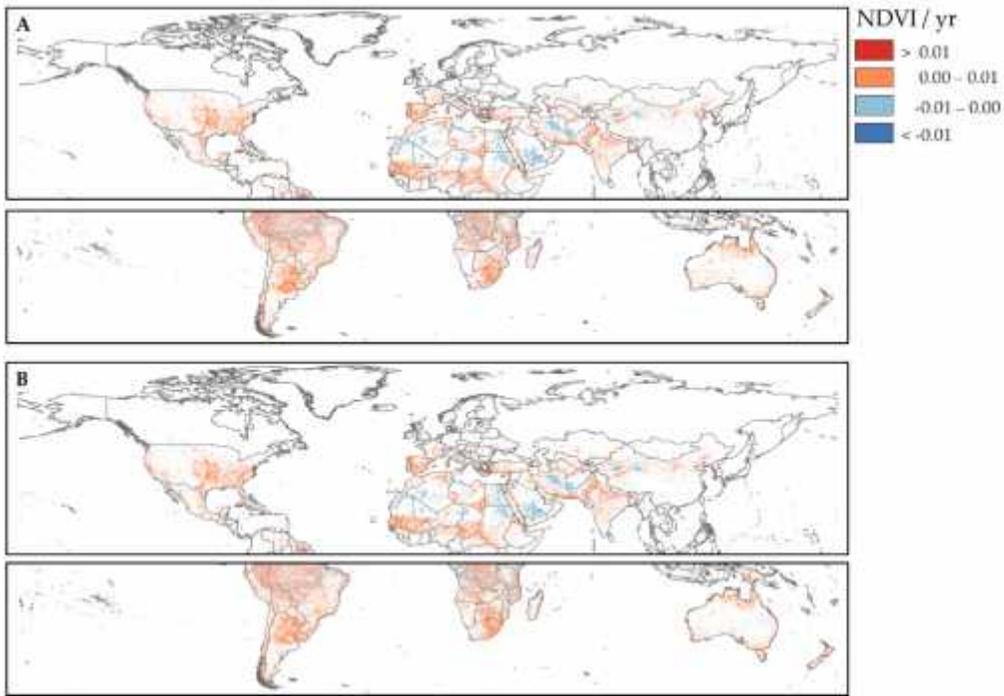


Figure 8. The change in maximum Normalized Difference Vegetation Index (NDVI) per year (yr) from the a) Global Inventory Modeling and Mapping Studies (GIMMS) and b) Vegetation Index & Phenology Lab (VIP) records. The upper panels represent the northern hemisphere (30 year change) and the lower panels represent the southern hemisphere (29 year change). The trends have been masked for significance ≥ 0.05 .

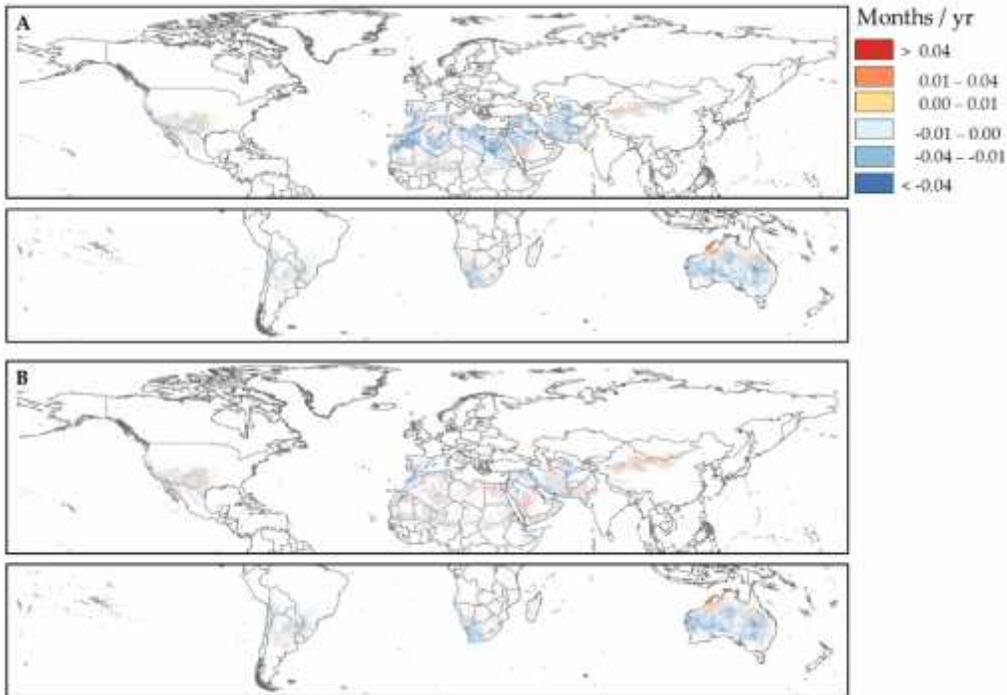


Figure 9. The change in timing of the Normalized Difference Vegetation Index (NDVI) per year (yr) from the a) Global Inventory Modeling and Mapping Studies (GIMMS) and b) Vegetation Index & Phenology Lab (VIP) records. The upper panels represent the northern hemisphere (30 year change) and the lower panels represent the southern hemisphere (29 year change). Negative values indicate earlier green-up/scenence, while positive values indicate later green-up/scenence. The trends have been masked for significance $\alpha = 0.05$.

Table 1. Summary statistics (R^2 = coefficient of determination, m = slope, b = intercept, p = significance, and **RMSE** = root-mean-square error) of the linear relationships between the Fraction of Photosynthetically Active Radiation intercepted by the canopy (F_{PAR}) estimated by Landsat Thematic Mapper or Enhanced Thematic Mapper Plus and F_{PAR} estimated by the long-term vegetation records (NDVI3g = Global Inventory Modeling and Mapping Studies Normalized Difference Vegetation Index Version 3, NDVI3v = Vegetation Index & Phenology Lab Version 3 Normalized Difference Vegetation Index, and EVI3v = Vegetation Index & Phenology Lab Enhanced Vegetation Index 2).

Crop	Product	R²	m	b	p	RMSE
<i>Maize</i> N = 98	NDVI3g	0.33	0.61	0.416	<0.001	0.178
	NDVI3v	0.29	0.73	0.201	<0.001	0.183
	EVI3v	0.26	0.65	0.178	<0.001	0.163
<i>Pasture</i> N = 22	NDVI3g	0.62	0.72	0.106	<0.001	0.110
	NDVI3v	0.68	0.85	-0.100	<0.001	0.101
	EVI3v	0.71	0.81	-0.038	<0.001	0.071
<i>Soybean</i> N = 39	NDVI3g	0.40	0.82	0.146	<0.001	0.168
	NDVI3v	0.47	1.09	-0.212	<0.001	0.158
	EVI3v	0.40	0.86	0.086	<0.001	0.125
<i>Wheat</i> N = 28	NDVI3g	0.59	0.86	0.222	<0.001	0.148
	NDVI3v	0.40	0.84	0.058	<0.001	0.177
	EVI3v	0.27	0.74	0.096	0.004	0.140

Table 2. Summary statistics (R^2 = coefficient of determination, m = slope, b = intercept, p = significance, and **RMSE** = root-mean-square error) of the relationships between *in situ* Leaf Area Index (LAI) and Fraction of Photosynthetically Active Radiation intercepted by the canopy (F_{PAR}) estimated by the downscaled long-term vegetation records (NDVI3g = Global Inventory Modeling and Mapping Studies Normalized Difference Vegetation Index Version 3, NDVI3v = Vegetation Index & Phenology Lab Version 3 Normalized Difference Vegetation Index, and EVI3v = Vegetation Index & Phenology Lab Enhanced Vegetation Index 2). A logarithmic transformation was performed for soybean to meet the assumptions of normality, while the *in situ* LAI from the other crops were not transformed.

Crop	Product	R²	m	b	p	RMSE	Transformation
<i>Maize</i> N = 98	NDVI3g	0.28	7.02	-1.942	<0.001	1.405	Linear
	NDVI3v	0.22	6.67	-1.695	<0.001	1.461	Linear
	EVI3v	0.21	7.87	-0.739	<0.001	1.474	Linear
<i>Pasture</i> N = 22	NDVI3g	0.49	4.65	-0.532	<0.001	0.665	Linear
	NDVI3v	0.38	3.90	-0.244	0.002	0.733	Linear
	EVI3v	0.43	5.46	0.097	<0.001	0.704	Linear
<i>Soybean</i> N = 39	NDVI3g	0.50	5.56	-3.264	<0.001	0.756	Logarithmic
	NDVI3v	0.51	5.12	-2.991	<0.001	0.753	Logarithmic
	EVI3v	0.39	6.89	-2.713	<0.001	0.838	Logarithmic
<i>Wheat</i> N = 28	NDVI3g	0.35	4.29	-0.482	<0.001	1.029	Linear
	NDVI3v	0.25	4.34	-0.504	0.007	1.107	Linear
	EVI3v	0.29	7.92	-0.806	0.003	1.077	Linear