Characterizing Leaf Area Index (LAI) and Vertical Foliage Profile (VFP)
over the United States

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Abstract

Leaf area index (LAI) and vertical foliage profile (VFP) are among the important canopy structural variables. Recent advances in lidar remote sensing technology have demonstrated the capability of accurately mapping LAI and VFP over large areas. The primary objective of this study was to derive and validate a LAI and VFP product over the contiguous United States using spaceborne waveform lidar data. This product was derived at the footprint level from the Geoscience Laser Altimeter System (GLAS) using a biophysical model. We validated GLAS derived LAI and VFP across major forest biomes using airborne waveform lidar. The comparison results showed that GLAS retrievals of total LAI were generally accurate with little bias ($r^2 = 0.67$, bias = $-0.13$, RMSE = 0.75). The derivations of GLAS retrievals of VFP within layers was not as accurate overall ($r^2 = 0.36$, bias = $-0.04$, RMSE = 0.26), and these varied as a function of height, increasing from understory to overstory −0 to 5 m layer: $r^2 = 0.04$, bias = 0.09, RMSE = 0.31; 10 to 15 m layer: $r^2 = 0.53$, bias = $-0.08$, RMSE = 0.22; and 15 to 20 m layer: $r^2 = 0.66$, bias = $-0.05$, RMSE = 0.20. Significant relationships were also found between GLAS LAI products and different environmental factors, in particular elevation and annual precipitation. In summary, our results provide a unique insight into vertical canopy structure distribution across North American ecosystems. This data set is a first step towards a baseline of canopy structure needed for evaluating climate and land use induced forest changes at continental scale in the future and should help deepen our understanding of the role of vertical canopy structure on terrestrial ecosystem processes across varying scales.

1 Introduction

Accurate measurements of three dimensional canopy structure and function play a key role in global carbon dynamics, climate feedbacks as well as biodiversity studies (Heimann and Reichstein, 2008; Loreau et al., 2001; Cramer et al., 2001; Schimel
Spatial variations of ecosystem structure largely inform the geographical patterns of ecological processes, including species richness (Cramer et al., 2001; Goetz et al., 2007; Turner et al., 2003). These structural variables, such as canopy height, leaf area index (LAI) and vertical foliage profile (VFP), have been identified as essential climate variables (ECV), essential biodiversity variables (EBV) or both (Pereira et al., 2013; Aber, 1979; Gower and Norman, 1991; Baret et al., 2013). Yet measurements of these canopy structural data are often limited at field sites, and their spatial distributions over broader geographical areas still remain poorly characterized due to heterogeneity of natural vegetation and inexact measuring techniques (Clark and Kellner, 2012; Asner et al., 2013). Improved spatial characterization of LAI and VFP at large scales may fill this observational gap and help clarify the role of spatial and vertical variability in canopy structure for carbon cycling, biodiversity and habitat quality (Houghton, 2007; Sauer et al., 2008).

Several global scale LAI products have been created from passive remote sensing data for many years (Myneni et al., 2002; Ganguly et al., 2012; Deng et al., 2006; Baret et al., 2007). Most of these products are derived by exploring the correlation between canopy foliage density and the total reflected intensity of electromagnetic radiation at multiple wavelengths. Applications of these LAI products have significantly improved the representation of the dynamics of terrestrial ecosystems and their interactions with the atmosphere (Mu et al., 2007; Zhao et al., 2005; Randerson et al., 2009). However, the overall accuracy of these products does not meet the requirements as specified by Global Terrestrial Observing System (GTOS: http://www.fao.org/gtos/org.html), and a key problem is the saturation of spectral signal over dense forests with high canopy cover (Abuelgasim et al., 2006; Shabanov et al., 2005; Yang et al., 2006). Saturation occurs because the solar flux decreases exponentially as it passes through a dense canopy, and the majority of the returned signal comes from the upper canopy in the form of direct reflectance and multiple scattering (Gower and Norman, 1991; Nilson, 1971). This limits the observational capabilities of passive optical sensors, such as Landsat and MODIS, to estimate LAI over dense forests. Furthermore, deriving the fo-
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H. Tang et al.

The objective of this study is to characterize the continental scale variability of canopy structure across the United States using lidar observations from space. First, we implement our existing algorithm at the GLAS footprint level and compare the derived data
with LAI and VFP products from airborne lidar in different forest types. Next we map the aggregated LAI and VFP product according to different ecoregions and land cover types over CONUS. Finally we analyze the distribution of GLAS LAI across different environmental factors, including elevation and precipitation.

2 Methods

2.1 GLAS data

GLAS is a spaceborne, sampling waveform lidar sensor with the working wavelength in the near-infrared band (1064 nm). It emits laser pulses at a frequency of 40 Hz and records the energy reflected from both the ground surface and canopy in an approximately 65 m diameter footprint (Abshire et al., 2005). GLAS samples the Earth surface in transects with individual footprints separated by ~175 m along track, and with between track spacing that varies as a function of latitude (e.g. 30 km spacing between tracks at the equator and 5 km spacing at 80° latitude (Brenner et al., 2012)). As a result of this sampling pattern, GLAS does not provide a wall-to-wall observation of forests. Its spatial allocation of laser footprints is best defined as a pseudo-systematic sampling or cluster sampling strategy (Stahl et al., 2011; Healey et al., 2012). To obtain a spatially continuous estimate of LAI at continental scale, footprint level GLAS data would need to be extrapolated using other remote sensing data (Dubayah et al., 2008; Lefsky, 2010), or can be mapped into appropriate geographic strata such as land cover types or ecoregions.

2.2 Retrieval of GLAS LAI and VFP

We collected a total of 1 100 498 cloud-free GLAS data from Campaigns GLA01 and GLA14 data over the contiguous United States from 2003 to 2007. GLA01 included the complete recorded waveform at a vertical resolution of 15 cm for land surface products, and GLA14 products were comprised of geographical information and various
parameters calculated from the waveform (Harding and Carabajal, 2005). Low energy shots (peak energy < 0.5 Volt) were excluded from data process for retrieval quality control because those waveforms were susceptible to noise contamination. Shots during leaf-off season (November to March) were also filtered out over deciduous forests and mixed forests. LAI and its profiles (0.15 m at vertical resolution) were initially calculated for GLAS footprints based on a Geometric Optical and Radiative Transfer (GORT) model (Ni-Meister et al., 2001), and further corrected for slope effects using an iterative method (Tang et al., 2014a). Canopy VFP were calculated from integration of footprint level LAI profiles at height intervals of 0 to 5, 5 to 10, 10 to 15 and 15 to 20 m.

2.3 Comparison data sets

We validated LAI and VFP data sets using an airborne lidar system, LVIS (Laser Vegetation Imaging Sensor). LVIS is a medium resolution (~25 m diameter) waveform scanning lidar system designed by NASA Goddard Space Flight Center (GSFC) (Blair et al., 1999). It can image the terrestrial surface across a 2 km wide swath and has been deployed to map many different forest structural parameters at regional scales across diverse biomes (Tang et al., 2012; Drake et al., 2002; Swatantran et al., 2012). We calculated both total LAI and VFP at 5 m height intervals from existing LVIS data using our physically based model, which has been validated using different types of field measurements (destructive sampling, LAI-2000 and hemispherical photos) (Tang et al., 2012, 2014a; Zhao et al., 2013). LVIS data used in this study included major forest types from eastern, central and western US, including Maine forests just north of Orono, Maine (2003), Sierra National Forest in California (2008), mixed forests along Baltimore/Washington corridor (2003) and the White River National Wildlife Refuge in Arkansas (2006). These LVIS datasets were all collected during leaf-on season.

We also included a 30 m resolution Landsat LAI map to examine the spatial distribution of GLAS total LAI. Landsat has the longest earth observation history at moderate resolution (30 m), and for decades has provided a consistent and unique measure-
ment of terrestrial ecosystems. The Landsat LAI map was produced using Global Land Survey (GLS) 2005 orthorectified Landsat data (Ganguly et al., 2012).

2.4 Analysis

The comparison between LVIS and GLAS was performed at the GLAS footprint level. LVIS shots falling within a 32.5 m radius from a GLAS shot center were selected. We filtered GLAS footprints to have a minimum of 3 coincident LVIS shots to increase the likelihood that the LVIS data covered a sufficient portion of the larger GLAS footprints. Both LAI and the 5 m interval VFP of LVIS shots were averaged onto each coincident GLAS footprint for comparison. We also made a footprint level comparison between GLAS LAI and the Landsat LAI map. A 3 × 3 window was applied to each GLAS footprint center to extract the averaged Landsat LAI pixels. Pixels with invalid values (e.g. retrieval failure or non-vegetation pixel) were excluded in the comparison. Agreements of different LAI datasets were assessed by coefficient of determination, bias and RMSE (Root Mean Square Error):

\[
\text{bias} = \frac{1}{n} \sum_{i=1}^{n} (\text{GLAS}_i - \text{Ref}_i) \tag{1}
\]

\[
\text{RMSE} = \sqrt{\frac{1}{n} \sum_{i=1}^{n} (\text{GLAS}_i - \text{Ref}_i)^2} \tag{2}
\]

In Eqs. (1) and (2), \(\text{GLAS}_i\) is GLAS LAI (or VFP) value at footprint level and \(\text{Ref}_i\) is that extracted from LVIS or Landsat.

Next, we aggregated the footprint level GLAS data into terrestrial ecoregions based on subset of a global map (Olson et al., 2001). Statistical analysis of total LAI and LAI strata (VFP aggregated at every 10 m height interval) was performed subsequently for each ecoregion. We also analyzed the GLAS LAI and VFP distribution across different environmental gradients throughout CONUS. GLAS footprints were categorized according to different environmental factors, including vegetation type, topographic data...
and annual measurements of climate variables. The vegetation map was derived from the MODIS Land Cover Type product (MCD12Q1) at 500 m resolution following the IGBP scheme (Friedl et al., 2010). Elevation data was extracted from the void-filled 90 m resolution SRTM (Shuttle Radar Topography Mission) DEM data (Reuter et al., 2007). Precipitation, temperature and vapor pressure deficit information originated from the 800 m resolution 30 yr annual normal climate data developed by the PRISM Climate Group (PRISM, 2013).

3 Results

This section includes three major parts: the first part focuses on the validation and comparison of GLAS LAI and VFP data with existing products; the second presents the geographical distribution of GLAS LAI and VFP, and; the last part shows their relationship with environmental factors.

3.1 GLAS LAI and VFP comparisons with LVIS and Landsat

The footprint level comparison between GLAS LAI and LVIS LAI had an overall $r^2$ of 0.60, bias of $-0.23$, and RMSE of 0.82 (Fig. 1). Except for a few outliers at the lower range of LAI, most of the comparison points were distributed along the 1 : 1 line suggesting no systematic difference between the two data sets. No significant bias was found across individual sites either.

The agreement of the 5 m height interval VFP distributions between the two data sets was lower than that of total LAI ($r^2 = 0.36$, a bias $= -0.04$ and RMSE $= 0.26$). Although there was no systematic bias observed when all sites and vertical intervals are considered (Fig. 2), examination by layer showed that GLAS overestimated understory LAI (0 to 5 m) ($r^2 = 0.04$, bias $= 0.09$, RMSE $= 0.31$) when compared with LVIS LAI (Fig. 3) but agreement improved as the vertical height interval considered moved higher in the canopy (5 to 10 m, $r^2 = 0.33$, bias $= -0.13$, RMSE $= 0.29$; and 10 to 15 m, $r^2 = 0.53$,
bias = −0.08, RMSE = 0.22), reaching a maximum at the top of the canopy (15 to 20 m, 
\( r^2 = 0.66 \), bias = −0.05, RMSE = 0.20).

The comparison between Landsat LAI and GLAS LAI had a much lower agreement
than that of LVIS \( (r^2 = 0.18, \text{bias} = 0.18 \text{ and RMSE} = 2.02) \) (Fig. 4). Even though the
two data sets agreed well at lower LAI values, Landsat overestimated LAI at the middle
range (from LAI values of 1 to 3) and then saturated above a value of about 4 to 5
against GLAS data.

### 3.2 Aggregated GLAS LAI and VFP within Ecoregions

We next mapped GLAS LAI across US ecoregions (Fig. 5). Highest LAI values were
found along northern Pacific Coast while lowest values occurred in the basin and range
province and the arid rains shadow region east of the Rocky Mountains. Northern
California coastal forests (Pacific temperate rainforests) were found to have the high-
est mean LAI value of 5.24. In the eastern US, the mixed deciduous forests of the
Appalachian–Blue Ridge province had the highest value of 3.95 while other ecogreions
around north–south direction of Appalachian Mountains had similar LAI values around
3 ~ 4 (Table 1). Forest ecogreions with lowest LAI values (excluding desert, shrub-
land and grassland) were located in Arizona mountains forests (1.15) and Great Basin
montane forests (0.90). Differences between these ecoregion-level LAI were signifi-
cant based on a bonferroni adjusted \( t \) test, except for those among Willamette Valley
forests, Appalachian–Blue Ridge forests, Puget lowland forests and Appalachian mixed
mesophytic forests (\( p \) values > 0.05).

LAI strata formed by VFP at each 10 m height interval were also averaged and
mapped across the US (Fig. 6). We chose the 10 m height interval rather than that
of 5 m because LAI strata aggregated at 10 m height interval represented a more ac-
curate and reliable description of vertical canopy structure given the relatively lower
measurement accuracy in the understory (< 5 m) we found in comparison to LVIS data.
Each strata showed a generally similar geographic pattern as that of total LAI with the
decreasing trend from coast to interior lands, but the specific patterns among strata
Northwestern forests were observed to have the highest total LAI values as well as LAI strata values. Northern California coastal forests exhibited the largest total LAI value as well as highest foliage density under 20 m height, while British Columbia mainland coastal forests showed the highest foliage density (1.13) above 20 m height with a lower total LAI value (4.74).

The distribution of GLAS total LAI and profiles were examined across different land cover types (Fig. 7 and 8). Not surprisingly, forests were found to have a consistently greater value than non-forest biomes in both total LAI and its strata. For example, deciduous broadleaf forests had the highest value of total LAI (mean = 4.03) as well as that of middle and upper LAI strata (height > 10 m), while open shrubland showed the lowest total LAI values of 0.77. However, vertical LAI distributions of most forests and non-forests were similar with peak foliage density distributed around a height of 2 ~ 4 m. The only exception was deciduous broadleaf forest, of which most of leaves were distributed at middle-story level with a peak height at about 8 m. Its VFP values did not decrease significantly until reaching a height of 15 m.

### 3.3 GLAS LAI Distributions by Environmental Factors

A linear regression analysis between GLAS LAI and SRTM DEM showed that increasing altitude led to an overall decreasing, but non-monotonic, trend in LAI values (LAI = 3.60 - 0.686 × Elevation (km), r² = 0.59, all P < 0.01) (Fig. 9). GLAS LAI values increased with DEM at the elevation range from 0 to 750 m and 2000 to 3000 m. The variation in the LAI-DEM relationship agreed well with Forest Ratio (LAI = 0.112 + 3.18 × Forest Ratio, r² = 0.45, P < 0.01). Here Forest Ratio was, defined as the percentage of footprints classified as forests in total GLAS shots (forest and non-forest). A multiple linear regression analysis showed that about 87 % of total variance could be explained by a simple combination of elevation groups and Forest Ratio values: LAI = 2.59 × Forest Ratio - 0.595 × Elevation (km) + 1.58.

We also analyzed GLAS LAI by 30 yr normal annual climate data using linear regression models (Fig. 10). It was observed that increasing precipitation signifi-
cantly increased LAI values ($\Delta$LAI = 1.84 per 1000 mm precipitation increase) but only at low and moderate precipitation levels (<2400 mm): $\text{LAI} = 1.84 \times \text{precipitation (mm)} \times 10^{-3} + 0.774$, $r^2 = 0.96$, adj $r^2 = 0.95$, $P < 0.01$. It contributed little when exceeding that threshold ($\text{LAI} = 0.22 \times \text{precipitation (mm)} \times 10^{-3}$, $r^2 = 0.40$, adj $r^2 = 0.30$, $P = 0.09$), as we found no significant LAI increase among groups greater than 2400 mm using a bonferroni adjusted $t$ test. GLAS LAI was also negatively but slightly correlated with minimum (maximum) vapor pressure deficit with a Pearson’s correlation coefficient of $-0.29$ ($-0.15$). The correlation coefficients between GLAS LAI and annual mean/minimum/maximum temperature were even lower with values of 0.13, 0.18 and 0.08 respectively.

Finally, we applied multiple linear regression analysis to illustrate the combined environmental effects of altitude and precipitation the on distributions of LAI and VFP. The regression analyses were conducted at both GLAS footprint level and aggregated scale on altitude and precipitation groups. At footprint level, altitude and precipitation together explained about 30% of variance of total LAI ($\text{LAI} = 2.73 - 0.69 \times \text{Elevation (km)} + 0.58 \times \text{precipitation (mm)} \times 10^{-3}$, $r^2 = 0.29$, adj $r^2 = 0.29$, $P < 0.01$). However, their correlations with footprint level VFP (0–10 m, 10–20 m and >20 m height intervals) were not significant with $r^2$ of 0.07, 0.12 and 0.08 respectively. At the aggregated scale, there was a better relationship between averaged LAI (VFP) values and environmental factors. The combination of altitude and precipitation can explain more than 60% variance in both total LAI and VFP, but explains only about 36% of variance on LAI for canopies less than 10 m height.

4 Discussion

In this study, we generated GLAS estimates of LAI and VFP across the United States, validated with an airborne lidar sensor, LVIS. Comparisons between LVIS and GLAS LAI and VFP estimates in different forest types across the United States show that GLAS generally provides accurate LAI and VFP estimates at footprint level. Consid-
er the temporal offset and spatial resolution differences between LVIS and GLAS, their overall agreements on LAI and VFP are reasonable ($r^2 = 0.60$, bias = $-0.23$, and RMSE = 0.82; and $r^2 = 0.36$, bias = $-0.043$, and RMSE = 0.26). Our comparisons further demonstrate the efficacy of our retrieval methods over continental scales that encompass large gradients in environmental factors and variability in forest types.

Measurement accuracy of GLAS VFP was lower compared to total LAI but ($r^2 = 0.36$, bias = $-0.043$, and RMSE = 0.26). Accuracies decreased for the lowest canopy layers, with the $r^2$ values falling from a peak of 0.66 at upper-story (15 to 20 m) to 0.33 at middle-story (5 to 10 m), to essentially no relationship in the lowest 5 m in the understory. There may be multiple factors contributing to this trend. First, a slope effect may reduce measurement accuracy of GLAS (Tang et al., 2014a). Slopes can blur the boundary between vegetation and topography signals in a lidar waveform, making their separation difficult and potentially leading to the error in LAI and VFP estimates. Despite methods to correct for topography (Lee et al., 2011; Tang et al., 2014a; Park et al., 2014), this effect cannot be fully mitigated, especially over steep slopes, and consequently may introduce errors and uncertainties into VFP estimates. Additionally, topographical effects can lead to a vertical misalignment of VFP between LVIS and GLAS. GLAS measures the terrestrial surface at a larger footprint with higher topographical variations, and a direct average of LVIS VFP can possibly result in a mismatch of vertical foliage distribution up to several meters. For example, consider two adjacent LVIS shots with the same VFP distribution but a 1 m difference in ground elevation (like a signal lag in the waveform). Adding the two waveforms along the geodetic altitude would lead to a 1 m vertical offset in the averaged waveform (pseudo-GLAS waveform) and produce a different VFP using the direct average method in a normalized coordinate system. But their total LAI values remain the same as long as the total energy from ground and vegetation can be separately correctly. Reducing vertical resolution of VFP can partially mitigate the mismatch effect because a lower vertical resolution requires integration over longer vertical axis which is more tolerant to ground mismatch. Take the above example again, the two VFP, at 1 m vertical resolution, do not match
each other at all along the entire waveform due to the offset. However, integration at every 5 m creates a signal overlap of 4 m in each height bin with a maximum of 20 % measurement error. Thus there is ultimately a tradeoff between vertical resolution and VFP accuracy. It also explains the higher agreement of total LAI (essentially an integration of VFP over the entire canopy) in the comparison between LVIS and GLAS. Lastly, measurement of near-ground understory vegetation by GLAS is difficult. By default GLAS waveforms are processed by a Gaussian decomposition method to get an approximate fit comprised of a series of Gaussian functions where the last one usually represents the ground (Hofton et al., 2000). The upper tail of the ground Gaussian peak may be mixed with signals from lower understory, and their separation is problematic, especially over slopes. All of these factors, plus the nature of high complexity and heterogeneity in canopy understory (Aubin et al., 2000; Valladares and Niinemets, 2008), may help explain the lower agreement on understory VFP between LVIS and GLAS.

Comparison between GLAS and Landsat displayed a much lower agreement than that of LVIS, was somewhat biased, and showed clear signals of saturation beyond LAI values of about 5. This result, along with all previous studies (Tang et al., 2012, 2014b), clearly showed the non-saturation advantage of lidar data against passive remote sensing in observing high LAI forests. On the low end of LAI spectrum, GLAS values were lower as compared with Landsat. There are different factors (some in the LVIS comparison too) could possibly lead to their difference in LAI estimates such as geolocation errors of GLAS shots, observation scale difference (65 m vs. 30 m) and misclassifications from MODIS land cover types (mainly impacting the correction of clumping effect). But this underestimation should be largely due to the fact that GLAS may not be able to adequately capture LAI values of short grassland with limited vertical structure or areas of sparse canopy cover, whereas Landsat is able measure such areas based on their total spectral response (tree and grass).

Analysis of GLAS LAI and VFP across ecoregions displayed a reasonable and expected geographical distribution. The great advantage of lidar based estimates is that they can produce LAI vertical strata maps, providing a view of canopy variability across
ecosystem types over large areas. Specifically, we can identify the foliage concentrations at various vertical bins and at spatial resolutions of interest (Figs. 6 and 8 and Table 1). This approach may reduce errors that arise from assumptions of uniformly distributed foliage within canopy, and could potentially be a contribution towards continental scale ecological and biological studies of forest structure and dynamics.

LAI and VFP also varied across different landscapes represented by various land cover types. As expected, we found both total LAI and maximum value of foliage density significantly increase along the vegetation gradient described by the transition from shrubland to savanna to woody savanna to forests (Figs. 7 and 8). In particular, we found deciduous broadleaf forest showing a different pattern with its foliage more evenly distributed in understory and mid-story when compared with all other forests. Our results suggest the existence of canopy layering, and highlight the feasibility of quantifying these layers across landscapes (Whitehurst et al., 2013). Regardless, of whether the data are conceptualized as layers or as continuously varying profiles, they nonetheless provide the actual vertical structure, and thus should help refine current empirical assumptions about vegetation structure of different land cover types in current LAI inversion algorithms (e.g. MODIS) and in ecosystem models (Hurtt et al., 2010; Antonarakis et al., 2014).

Elevation and precipitation were found to be significantly correlated with LAI at both footprint level and across aggregated groupings by elevation and forest ratio. LAI decreased with elevation and this trend was consistent with previous studies (Luo et al., 2004; Moser et al., 2007; Pfeifer et al., 2012). Variations of the trend can be largely explained (about 45 % of total variance) by the Forest Ratio (defined in Sect. 3.3). A combination of the two factors (elevation groups and Forest Ratio) explained almost 90 % variance of average LAI spatial distribution. We also found a significant but nonlinear relationship between GLAS LAI and annual precipitation (Fig. 10). This non-linear relationship agrees with previous studies in the tropics (Pfeifer et al., 2014; Spracklen et al., 2012). However, we found no significant variation of GLAS LAI with either temperature or vapor pressure deficit variables. A combined effect of elevation and precipitation
explains about 30% of LAI variation at GLAS footprint level, suggesting the natural complexity highly spatial variability of LAI distribution.

As a direct quantification of 3-D foliage distribution, GLAS LAI profiles are thus far the best representations of terrestrial ecosystem structure over broad geographical areas and suggest that ecological applications of these profiles are worth exploring. First, this data could refine large scale modeling of plant respiration and photosynthesis and consequently and improve ecosystem modeling (Houghton, 2007). Previous studies have reported a potential 50% underestimate of GPP values when vertical foliage stratification is not considered (Kotchenova et al., 2004; Sprintsin et al., 2012). A consistent, global data set of VFP should thus improve initialization of ecological models (Hurtt et al., 2004), and refine estimation of GPP, in conjunction with passive remote sensing data (Turner et al., 2006). Secondly, these profiles may be important descriptors of habitat as related to biodiversity and habitat quality. Many studies have confirmed the general relationship between species richness, habitat heterogeneity and forest structural complexity across different landscapes (Swatantran et al., 2012; Goetz et al., 2010; Schut et al., 2014; Ferger et al., 2014). The inclusion of LAI profiles provides spatially explicit vegetation structure data and may potentially improve current observations of species distribution at continental scale, e.g. for avian species (Sauer et al., 2008; Culbert et al., 2013), and lead to entirely new biodiversity metrics (e.g. see Huang et al., 2014). For example the concept of an “edge” has been traditionally defined as the boundary between forest and non-forest areas. LAI profiles provide a means of defining new edges based on differences in LAI as a function of height, so the edge is now the boundary between a rapid change in foliage density at a particular height.

5 Conclusion

Accurate representation of canopy vertical structure and its dynamics has long been recognized as a priority because it represents a key interface between terrestrial sur-
face and atmosphere and impacts the water and carbon cycles, and their transfer of energy and mass. Foliar profiles are also increasingly recognized as important determinants for habitat quality, species distribution, diversity and abundance. As ecosystems come under increasing pressure from climate and land use change, global data sets of canopy structure are needed to help better understand the consequences of these changes on ecosystem form, function and services.

In this paper we have demonstrated the potential for global mapping of key canopy structures, LAI and VFP, from space. While imperfect, given their large footprint and sparse sampling, the waveforms from ICESat are currently the only such global data set of structure. Our ability to produce this data set is the end result of a series of research experiments that linked various types of observations, from destructive profiles, to ground based optical methods, to airborne lidar, to passive optical retrievals. This background gives us confidence that meaningful and useful data on LAI and VFP can be derived from future spaceborne lidar. There are still hurdles to overcome related to topography, understory accuracy, model assumptions and parameterizations, such as ground/canopy reflectance ratios and foliage clumping, among others, to achieving higher accuracy. We anticipate these will be resolved in time and lead to an even more capable model suitable for the next generation of waveform lidar observations from space, such as NASA’s Global Ecosystem Dynamics Investigation (GEDI) (Dubayah et al., 2014) and potentially ICESat-2 (Abdalati et al., 2009).

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Characterizing Leaf Area Index (LAI) and Vertical Foliage Profile (VFP)

H. Tang et al.


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### Table 1. Ecoregions with highest total LAI values (unit: m² m⁻²).

<table>
<thead>
<tr>
<th>Ecoregions</th>
<th>Total LAI Mean(±SD)</th>
<th>LAI 0–10 m Mean(±SD)</th>
<th>LAI 10–20 m Mean(±SD)</th>
<th>LAI &gt; 20 m Mean(±SD)</th>
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</thead>
<tbody>
<tr>
<td>Northern California coastal forests</td>
<td>5.24 ± 2.11</td>
<td>2.06 ± 1.32</td>
<td>1.67 ± 1.09</td>
<td>1.08 ± 1.15</td>
</tr>
<tr>
<td>Central Pacific coastal forests</td>
<td>5.00 ± 2.14</td>
<td>1.52 ± 1.61</td>
<td>1.10 ± 1.16</td>
<td>0.84 ± 1.25</td>
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<tr>
<td>British Columbia mainland coastal forests</td>
<td>4.74 ± 2.26</td>
<td>1.48 ± 1.31</td>
<td>1.23 ± 1.08</td>
<td>1.13 ± 1.13</td>
</tr>
<tr>
<td>Central and Southern Cascades forests</td>
<td>4.31 ± 2.34</td>
<td>1.06 ± 1.35</td>
<td>0.79 ± 1.02</td>
<td>0.64 ± 1.07</td>
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<tr>
<td>Klamath–Siskiyou forests</td>
<td>4.31 ± 2.31</td>
<td>1.26 ± 1.30</td>
<td>0.99 ± 1.07</td>
<td>0.73 ± 0.99</td>
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<td>Willamette Valley forests</td>
<td>3.99 ± 2.24</td>
<td>0.73 ± 1.09</td>
<td>0.60 ± 0.89</td>
<td>0.75 ± 1.31</td>
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<td>Appalachian–Blue Ridge forests</td>
<td>3.95 ± 2.03</td>
<td>1.04 ± 1.27</td>
<td>0.82 ± 0.99</td>
<td>0.47 ± 0.82</td>
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<td>Puget lowland forests</td>
<td>3.91 ± 2.25</td>
<td>0.98 ± 1.39</td>
<td>0.71 ± 1.08</td>
<td>0.40 ± 0.81</td>
</tr>
<tr>
<td>Appalachian mixed mesophytic forests</td>
<td>3.86 ± 2.04</td>
<td>1.06 ± 1.29</td>
<td>0.77 ± 0.93</td>
<td>0.48 ± 0.83</td>
</tr>
<tr>
<td>North Central Rockies forests</td>
<td>3.67 ± 2.27</td>
<td>1.61 ± 1.55</td>
<td>0.84 ± 0.89</td>
<td>0.47 ± 0.72</td>
</tr>
</tbody>
</table>
**Figure 1.** A comparison between two lidar derived Leaf Area Index (LAI) datasets at different sites across the US ($N = 318$), produced from the Laser Vegetation Imaging Sensor (LVIS) and the Geoscience Laser Altimeter System (GLAS) respectively. Each point represents a comparison at GLAS footprint while different colors and shapes indicate different sites (AR: White River National Wildlife Refuge in Arkansas; CA: Sierra National Forest in California; MD: Baltimore/Washington corridor in Maryland; ME: Maine forests to the north of Orono, Maine). The comparison produces $r^2$ of 0.60, bias of $-0.23$, and RMSE of 0.83. Dashed line is the 1:1 line.
**Figure 2.** A comparison of Vertical Foliage Profile (VFP) density derived from LVIS and GLAS over different sites in the US (same sites as Fig. 1 but with $N = 1272$). Each VFP point represents an integrated value of foliage density at 5 m height interval.
Figure 3. Comparison between LVIS and GLAS VFP density integrated at every 5 m height interval (from ground to canopy top).
Figure 4. Comparison between Landsat LAI and GLAS LAI over contiguous US: (a) density scatter plot of Landsat and GLAS LAI ($r^2 = 0.18$, bias = 0.18 and RMSE = 2.02); (b) Difference between Landsat and GLAS LAI. Darker kernel density color refers to more clustered distribution of LAI pairs.
Figure 5. GLAS LAI distributions by ecoregion.
**Figure 6.** LAI strata distributions by WWF ecoregions. Despite similar total LAI values, the southeastern forests show different LAI values at stratified height intervals.
Figure 7. Distribution of total GLAS LAI across different land cover types. The width of the boxes is proportional to the number of observations for each type ($N =$ Evergreen Needleleaf: 45 207, Evergreen Broadleaf: 438, Deciduous Needleleaf: 123, Deciduous Broadleaf: 48 283, Mixed Forest: 62 053, Closed Shrubland: 4087, Open Shrubland: 7364, Woody Savanna: 43 536, Savanna: 3051). Notches show the approximate 95% confidence interval of the median.
Figure 8. Averaged GLAS VFP for different land cover types across US: non-forest vegetation types (left) and forest types (right). Mean values are central lines within the color-filled 95% CI envelope.
Figure 9. Distribution of GLAS LAI (left axis) and Forest Ratio – GLAS shots over forest divided by total shot numbers – (right axis). Overall, there is a decreasing trend of LAI values as elevation increases, but deviations occur from this trend that are associated with elevational variation in Forest Ratio.
Figure 10. Distribution of GLAS LAI as a function of precipitation. A linear regression analysis of LAI values averaged by precipitation groups shows an increasing trend up to areas of about 2400 mm (blue line). Beyond this value the rate of change slows considerably (magenta line) but the trend is only weakly significant ($P = 0.09$).