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C.W. Hunt (Referee)

chunt@unh.edu Received and published: 18 August 2014

Review of Abril et al. "Technical Note: Large overestimation of pCO<sub>2</sub> calculated from pH and alkalinity in acidic, organic-rich freshwaters."

Comment 1 - "GENERAL COMMENTS

I read this manuscript with great interest, as it addresses a topic my colleagues and I have been considering for quite a while, with large implications for the current understanding of freshwater air-water CO<sub>2</sub> fluxes. The featured dataset of concurrent pH/TA/pCO<sub>2</sub> measurements represents a valuable first step in examining potential errors in large-scale pCO<sub>2</sub> flux estimates based on TA and pH measurements, and the decision to structure the manuscript as a technical note seems sound. However, the data also present an opportunity to explore possible ways to refine or even correct estimates of pCO<sub>2</sub> from the TA/pH pairing. While such refinements or corrections would most likely be variable and site-dependent, they offer a chance to substantially improve current regional or global CO<sub>2</sub> flux estimates.

The scientific quality of this paper is high. The authors employed well-documented methods for field and laboratory measurements, and a reasonable approach to derive pCO<sub>2</sub> from TA/pH. Except for some suggestions to improve language usage listed in under Technical Corrections, the authors do a fine job presenting their data and findings."

Reply 1 – We thank C.W. Hunt for his very positive overall evaluation of our MS. Indeed, one crucial question raised by the three referees can be summarized as: "Is it possible to correct the bias in calculated pCO<sub>2</sub> data?". In fact, we tested several methods to derive empirical relationships that could be used to correct the pCO<sub>2</sub> calculated from pH and TA. Unfortunately, we found no reliable consistent quantitative relationships to allow correcting for the bias in pCO<sub>2</sub> when values are derived from pH, DOC, and TA.

The first approach consisted in calculating organic alkalinity from pH and DOC using the models of Driscoll et al (1989) –which assumes a single apparent pK value for organic acids- and the model of Hruska et al. (2003) – a triprotic model which assumes three apparent pK values-. These two organic acid models applied to our data led to very similar organic alkalinity values (See attached figure 1A). The organic alkalinity was then subtracted from the TA and the pCO<sub>2</sub> was re-calculated from the measured pH and the TA value corrected from organic acids. pCO<sub>2</sub> values corrected that way were, however, still very different from those measured in the field (See figure 1B), being sometimes higher and sometimes lower than the measured values.

The second approach consisted in subtracting from the measured TA, the alkalinity calculated with the CO<sub>2</sub>sys program using as input parameters the measured pH and pCO<sub>2</sub>, in order to derive a non-carbonate alkalinity (NCA). Besides the fact that NCA derived that way was often negative (probably due to large sensitivity of calculation from the pCO<sub>2</sub>/pH pairing), it was neither (or weakly) correlated with DOC, nor with pH (See Figure 2). Consequently, we could not derive any empirical relationship useful to correct for the bias in pCO<sub>2</sub> calculation.

Change 1 – In the revised version of our MS, we include a description of these attempts to correct calculated pCO<sub>2</sub> from the available parameter (pH, TA, DOC and measured pCO<sub>2</sub>).

Comment 2 - "SPECIFIC COMMENTS

P11708 L22: Some discussion of the effect of filtering TA samples would be helpful. The one study I am aware of which discusses the filtration effect on freshwater alkalinity is Williams et al. (2009), who did not observe a significant difference between filtered and unfiltered alkalinities. However, that study examined low-DOC waters, which may differ from the waters in this study. Technically you measured dissolved alkalinity (DA), not total alkalinity (TA). They are probably functionally equivalent, but it is a point worth mentioning. Chanson and Millero (2007) discuss the

57 filtration question with regard to open-ocean samples, and found no difference between filtered  
58 and unfiltered alkalinity. However, the particulate loads in your sampled rivers are probably much  
59 larger than open ocean waters.

60  
61 Reply 2 – All our TA measurements were made on filtered samples. We agree with the referee  
62 that filtering is crucial in freshwaters. However, we made no comparison between filtered and  
63 unfiltered samples, so we cannot improve any understanding on this aspect. However, we expect  
64 that in white water rivers, with high TSM values, there could be an effect of dissolution of CaCO<sub>3</sub>  
65 particles during titration that should affect the TA values. If we refer to definition of total alkalinity  
66 of Dickson (1981) it refers to the total sum of bases in one kilogram of sample for a pH > 4.5.  
67 Hence, TA should by definition relate solely to solutes and exclude particles.

68  
69 Dickson, AG. 1981. An exact definition of total alkalinity and a procedure for the estimation of  
70 alkalinity and total inorganic carbon from titration data. Deep-Sea Research Part a-  
71 Oceanographic Research Papers. 28:609-623.

72  
73 Change 2 – In the absence of comparative data, we avoided a long discussion on TA titration on  
74 unfiltered samples. Nevertheless, we mentioned in the revised MS that TA measurements should  
75 be done on filtered samples, otherwise some additional bias would appear in turbid samples.

76  
77 Comment 3 – “P11710 L14-28 This section provides some very valuable insights, and I especially  
78 appreciate the ranking of sites in Table 2 by pH and TA to illustrate the trends in pCO<sub>2</sub>  
79 overestimation, but I urge the authors to explore the interactions between pCO<sub>2</sub> over- estimation  
80 and pH and DOC further. I point this out because although the pairing of TA and pH to estimate  
81 pCO<sub>2</sub> is problematic, as well illustrated in this work, it is also potentially extremely valuable, as  
82 there is a wealth of TA and pH data worldwide stretching back many decades. In our Hunt et al.  
83 (2011) paper we observed a relatively robust linear relationship between pH and the percentage  
84 of TA comprised of NCA, especially at pH<7. Subsequent unpublished data has shown the same  
85 trend, albeit with a somewhat different slope of the linear regression. It seems that the authors  
86 could attempt a similar regression, which may offer a simple pH-based correction factor, either  
87 gross or site-specific, which could be applied to the calculated pCO<sub>2</sub> values. I wonder if such a  
88 correction factor could yield ‘good enough’ corrected calculated pCO<sub>2</sub> to make the use of TA and  
89 pH on broad regional or global scales possible. Additionally, we have seen promise in power  
90 relationships between DOC concentration and non-carbonate alkalinity concentration, which  
91 could potentially also be used to derive a correction factor for use with the TA-pH pairing.”

92  
93 Reply 3 – Refer also to reply 1. We fully agree it would be useful to provide such empirical  
94 correction factor. The referee refers to a method he used to estimate non carbonate alkalinity  
95 (NCA), based on the comparison of measured TA with that calculated from pH and DIC. Our  
96 dataset concerns pCO<sub>2</sub>, pH, TA, Temperature and DOC only. So we could calculate TA from the  
97 pH/pCO<sub>2</sub> pairing, which is relatively uncertain in comparison with calculation from the pH/DIC  
98 pairing, owing to the strong correlation between these two parameters (pCO<sub>2</sub> and pH, see also  
99 comment 5). As can be seen in Figure B, we found no correlation between NCA (or the proportion  
100 of NCA in TA) and pH or DOC.

101  
102 Change 3 – In table 2, we added a ranking of our dataset by DOC concentration. This ranking  
103 confirmed the general interpretation on the role of DOC on pCO<sub>2</sub> overestimation.

104  
105 Comment 4 – “Further, it seems that the authors have all the data to test a multivariate regression  
106 of pCO<sub>2</sub> overestimation against a combination of DOC and pH, which may be even more robust.  
107 Any insights the authors could provide to improve the use of TA and pH to calculate pCO<sub>2</sub> would  
108 be of great value.”

109  
110 Reply 4 – A multivariable regression against DOC and pH is inappropriate as these two variables  
111 are not independent, since DOC contributes to acidity and is thus correlated with pH. Indeed, any  
112 insight to improve the use of TA and pH to calculated would be of great value. However, we found

113 no evidence for the feasibility of a generally valid approach to correct for the bias, and until further  
114 studies provide more insights, our conclusion remains that pH/TA pairing should be abandoned in  
115 acidic waters.  
116  
117 Change 4 – No change related to this specific comment (multivariable analysis)  
118  
119 Comment 5 – P11711 L18-23: This is interesting! There is evidence in the coastal ocean  
120 literature as well that pairing pCO<sub>2</sub> and pH to derive TA and DIC produces problematic results  
121 (Cullison Gray et al. 2011). The authors of this coastal paper also conclude that the pH- pCO<sub>2</sub>  
122 pairing are particularly sensitive to measurement error, but did not address how precise pH and  
123 pCO<sub>2</sub> measurements would need to be in order to accurately retrieve DIC and TA. Mention of this  
124 topic is useful in the current manuscript, but perhaps deserves further exploration. It is clear that  
125 this manuscript focuses on the derivation of pCO<sub>2</sub> from pH and TA, but the authors have all the  
126 data needed to perform a very nice sensitivity analysis on the pH-pCO<sub>2</sub> pairing. While this may  
127 deviate from the central question of the manuscript, I would be very interested in the results of  
128 such an analysis.  
129  
130 Reply 5 – Indeed, this deviates from the central question of the MS. Nevertheless, we stress in  
131 our MS the difficulty to calculate TA from the pH/pCO<sub>2</sub> pairing, as these variables are generally  
132 well correlated.  
133  
134 Change 5 – we cite Cullison Gray et al. (2011) in our revised MS to strengthen our  
135 argumentation.  
136  
137 TECHNICAL CORRECTIONS  
138 Comment 6 - General: Would it make sense to present alkalinity in µmol L<sup>-1</sup> instead of mmol L<sup>-1</sup>?  
139 Reply/Change 6 – We now use µmol L<sup>-1</sup> throughout the manuscript.  
140  
141 Comment 7 – General: I suggest defining shorthand terms: perhaps pCO<sub>2</sub>calc (calculated pCO<sub>2</sub>)  
142 and pCO<sub>2</sub>obs (measured pCO<sub>2</sub>).  
143 Reply 7 – If these expressions are correct, we find it easier, more comfortable, and clearer to  
144 repeat “calculated pCO<sub>2</sub>” and “measured pCO<sub>2</sub>” throughout the ms.  
145  
146  
147 Comment 8 – P11702 L5: I would recommend avoiding ‘Nowadays’. For this line perhaps  
148 substitute ‘Currently’  
149 P11702 L6: add comma: ‘freshwaters, and. . .’ P11702 L20-23: awkward sentence  
150 P11703 L28: I would recommend avoiding ‘Nowadays’. For this line perhaps substitute ‘Presently’  
151 Change 8 – Modified in the revised MS, as suggested  
152  
153 Comment 9 – P11703 L29 and throughout: Change ‘water’ to ‘aquatic’  
154 Reply 9 – “water pCO<sub>2</sub>” is widely used by opposition to air pCO<sub>2</sub>, and may have a different  
155 meaning from “aquatic pCO<sub>2</sub>”  
156  
157 Comment 10 – P11704 L1-2: Some citations here comparing usage of direct and indirect pCO<sub>2</sub>  
158 observations may be needed  
159 Reply 10 – it is difficult to provide a reference list, as it will never be exhaustive or quantitative  
160 (how many papers using calculated versus measured ? how many measured versus calculated  
161 pCO<sub>2</sub> data in each paper ?, etc...). We have chosen two significant references at the global  
162 scale (Cole et al 1994, Raymond et al. 2013).  
163  
164 Comment 11 –  
165 P11704 L3: change to ‘carbonic acid (which are a function of temperature)  
166 P11704 L16: change to ‘adapted to the variability of pCO<sub>2</sub> found in freshwaters.’  
167 P11704 L17: change ‘and’ to ‘to’ P11704 L18: change ‘First works consisted in’ to ‘Earlier  
168 examples provided a’ P11704 L22: change to ‘showed a variability of +/-500’

169 Reply/ Change 11 – modified in the revised MS as suggested  
170  
171 Comment 12 – P11704 L26: Is 7% agreement excellent? This level might be better described as  
172 a 'good' agreement. I suspect pairing TA with a pCO<sub>2</sub> +/- 7%, or even +/- 5%, would yield  
173 extremely variable calculated pH DIC values, for example.  
174 Reply/ Change 12 – changed to "good agreement"  
175  
176 Comment 13 – P11704 L27: change to 'or underestimated, but'  
177 P11704 L27-29: In this case, I don't think 'we' refers to all the authors of this manuscript.  
178 Consider rewriting this sentence to avoid the 'we' pronoun: 'Concurrent measurements  
179 of...performed in 2003 in acidic, humic-rich...showed that pCO<sub>2</sub> calculation. . .'  
180 P11705 L7: change to 'organic acid anions'  
181 P11705 L15: 'exponentially' is a really quantitative term. Consider substituting 'dramatically'  
182 Reply/ Change 13 – modified in the revised MS as suggested  
183  
184  
185 Comment 14 – P11706 L4-12: A map of sampling locations, perhaps with unique markers  
186 corresponding to the various publications referenced, would be very helpful  
187 Reply 14 – A world map showing the 12 sampling sites is provided in the revised MS  
188  
189 Comment 15 – P11708 L12-13: awkward sentence  
190 P11708 L16: Separate into two sentences, suggest changing to 'In addition to the IR analysers  
191 generally used in this work, in the Sinnamary River pCO<sub>2</sub> was also. . .'  
192 P11708 L29: change to 'capped'  
193 P11709 L1: change to 'Shimadzu'  
194 P11709 L17-18: awkward sentence  
195 P11709 L24-25: awkward sentence  
196 P11710 L7: change to 'more than 10%'  
197 P11710 L12: change to 'The largest overestimation of pCO<sub>2</sub>calc occurred in the  
198 most. . .'  
199 P11710 L14 Should this refer to Fig 3b, instead of Fig. 2b?  
200 P11711 L3: the phrase 'calculated pCO<sub>2</sub> overestimation' is awkward  
201 P11711 L6: the rivers in Hunt et al. 2011 were located in New Hampshire, Massachusetts, and  
202 Maine (USA) and New Brunswick (Canada). Perhaps just say 'New England (USA) and New  
203 Brunswick (CA) rivers'  
204 Reply / change 15 – modified in the revised MS as suggested.  
205  
206 Comment 16 – P11711 L9-10: In our Hunt et al. 2011 paper we made some important  
207 assumptions about non-carbonate inorganic contributions to alkalinity, but did not test these  
208 assumptions experimentally. We did not have concurrent nutrient measurements, so our N, P and  
209 Si values were taken from other work in the Oyster River. For other rivers described in that paper  
210 we did not have any N,P or Si data to examine their effects upon alkalinity. From the pKa values  
211 of nitrate, phosphate and silicate it is probably safe to assume that they do not contribute to  
212 alkalinity at the observed river pH. We also assumed that aluminum species did not contribute,  
213 which may or not be true.  
214 Reply / change 16 – We changed the sentence "in addition, the contribution of inorganic species  
215 other than carbonate was negligible and most of the NCA could be attributed to organic acid  
216 anions" to "in addition, the contribution of inorganic species other than carbonate was assumed  
217 negligible and most of the NCA was attributed to organic acid anions".  
218  
219 Comment 17 – P11711 L16-18 Awkward sentence, suggest 'evidencing the predominant role of  
220 organic acids in DIC speciation and pH in such acidic systems. Because we did not directly  
221 measure DIC. . .'  
222 P11712 L3 Suggest change 'far from' to 'not only' P11712 L23 Suggest changing 'above' to  
223 'requiring'  
224 P11714 L1 'World' does not need to be capitalized P11714 L6 Suggest removing 'ones'.

225 P11714 L5-10 This is a very long sentence. Suggest separating into two or three shorter  
226 sentences.  
227 P117714 L25 Suggest removing 'as such'  
228 Reply / change 17 – modified in the revised MS as suggested.  
229  
230 Comment 18 – REFERENCES  
231 Chanson, M. and F.J. Millero. 2007. Effect of filtration on the total alkalinity of open- ocean  
232 seawater. *Limnol. and Oceanogr.*: Methods 5: 292-295.  
233 Cullison Gray, S.E., DeGranpre, M.E., Moore, T.S., Martz, T.R., Friedrich, G.E. and K.S.  
234 Johnson. 2011. Applications of in situ pH measurements for inorganic carbon calculations. *Marine*  
235 *Chemistry* 125: 82-90.  
236 Williams, A.J., Andersen, C.B. and G.P. Lewis. 2009. Evaluating the effects of sample processing  
237 treatments on alkalinity measurements. *Journal of Hydrology* 377: 455- 464.  
238 Reply / change 18 We cite the Cullison Gray paper in order to strengthen the sensitivity to  
239 analytical errors in TA calculated from pH and pCO<sub>2</sub>.  
240  
241  
242 Anonymous Referee #2  
243 Comments to the manuscript by Abril et al. bg-2014-341, Technical note: “Large over- estimation  
244 of pCO<sub>2</sub> calculated from pH and alkalinity in acidic, organic-rich freshwaters”.  
245 Comment 1 – Overview :  
246 The study presents a technical note where directly determined pCO<sub>2</sub> (equilibrator or headspace  
247 techniques) and indirectly calculated pCO<sub>2</sub> (from total alkalinity, pH and temperature) were  
248 compared across a wide range of freshwater systems. The authors found a large deviation  
249 between the two methods where the indirect determination generally overestimate the pCO<sub>2</sub>.  
250 They further found that this overestimation was highest in low carbonate alkalinity, low pH and  
251 high DOC waters (> 100% in the 25% most organic-rich and acidic samples). They explain the  
252 large deviation by the influence of organic acids on the total alkalinity determination but also that  
253 the lower buffering capacity of the carbonate system at low pH increases the sensitivity of the  
254 calculated pCO<sub>2</sub>. The main message of the study is consequently that large scale studies of  
255 pCO<sub>2</sub> should not solely use indirect determination, this is especially true for low pH and organic-  
256 rich waters. Based on this the authors further suggest that recent large scale studies on CO<sub>2</sub>  
257 emissions from inland waters might have overestimated the flux due to being based on potentially  
258 biased data.  
259 The manuscript focus on an important methodological topic that is very suitable for publication in  
260 Biogeosciences I believe. The amount of published papers dealing with inland freshwater CO<sub>2</sub>  
261 emissions at various scales is rapidly increasing. Many of these are based on such indirect  
262 methods without or with little validation of the results. In the race of getting these global/regional  
263 estimates and due to the lack of directly determined pCO<sub>2</sub>, general chemistry data bases are  
264 often used but without considering the limitations of the data nor the methods used. Although the  
265 knowledge about weaknesses in the indirect determination of pCO<sub>2</sub> is not totally new, there is a  
266 need for bringing the discussion to light, something that I think this manuscript really does.  
267  
268 Reply 1 – We thank the referee for his/her very positive overall evaluation of our MS.  
269  
270 Comment 2 – General comments:  
271 With this background the manuscript is an important contribution to the research field. The  
272 authors present an impressive data set covering a wide geochemical and geographical range.  
273 Even not directly important for the comparison I especially appreciate the African contribution  
274 since this part of the world often is underrepresented in related studies. The manuscript is well  
275 written, includes solid data derived from standard and clear methodology and should be  
276 publishable after rather minor clarifications/additions. However, I find it a bit strange that the  
277 authors do not present ways to compensate for the influence of organic acids on the total  
278 alkalinity determination. This would be of great interest and use for future pCO<sub>2</sub> estimates based  
279 on standard water chemistry data. Methods to do that are there in the literature (see detailed  
280 comments below) and with the great data set that the authors present this would really lift the

281 story and the overall impact.

282

283 Reply 2 – Indeed, one crucial question raised by the three referees can be summarized as: “Is it  
284 possible to correct the bias in calculated pCO<sub>2</sub> data?”. In fact, we tested several methods to  
285 derive empirical relationships that could be used to correct the pCO<sub>2</sub> calculated from pH and TA.  
286 Unfortunately, we found no reliable consistent quantitative relationships to allow correcting for the  
287 bias in pCO<sub>2</sub> when values are derived from pH, DOC, and TA.

288

289 The first approach consisted in calculating organic alkalinity from pH and DOC using the models  
290 of Driscoll et al (1989) –which assumes a single apparent pK value for organic acids- and the  
291 model of Hruska et al. (2003) – a triprotic model which assumes three apparent pK values-.  
292 These two organic acid models applied to our data led to very similar organic alkalinity values  
293 (See attached figure 1A). The organic alkalinity was then subtracted from the TA and the pCO<sub>2</sub>  
294 was re-calculated from the measured pH and the TA value corrected from organic acids. pCO<sub>2</sub>  
295 values corrected that way were, however, still very different from those measured in the field (See  
296 figure 1B), being sometimes higher and sometimes lower than the measured values.

297

298 The second approach consisted in subtracting from the measured TA, the alkalinity calculated  
299 with the CO<sub>2</sub>sys program using as input parameters the measured pH and pCO<sub>2</sub>, in order to  
300 derive a non-carbonate alkalinity (NCA). Besides the fact that NCA derived that way was often  
301 negative (probably due to large sensitivity of calculation from the pCO<sub>2</sub>/pH pairing), it was neither  
302 (or poorly) correlated with DOC, nor with pH (See Figure 2). Consequently, we could not derive  
303 any empirical relationship useful to correct for the bias in pCO<sub>2</sub> calculation.

304

305 Change 2 – In the revised version of our MS, we include a description of these attempts to correct  
306 calculated pCO<sub>2</sub> from the available parameter (pH, TA, DOC and measured pCO<sub>2</sub>).

307

308

309 Comment 3 – Detailed comments:

310 Ln 5 p.11702. Replace “Nowadays” with “Still” for example.

311 Ln 15-17 p. 11703. I suggest to rephrase this sentence. In non-carbonate and organic- rich soils I  
312 would say that DOC commonly exceeds the DIC. Of course it is tricky to write about general  
313 global patterns. But I suggest to formulate it differently.

314 Reply / Change 3 – Modified in the revised MS, as suggested

315

316 Comment 4 – Ln 24-27 p.11703. It is not totally clear to me what variability the authors refer to  
317 when they mention that pCO<sub>2</sub> can vary more than 4 orders of magnitude. Spatially, temporally or  
318 maybe spatiotemporally? Could be clarified.

319 Reply / Change 4. In the revised MS, we specify that this refers to spatiotemporal variations.

320

321 Comment 5 – Ln 28 p. 11703. Again replace “Nowadays” with something more suitable Ln 3 p.  
322 11704. Should be “carbonic acid” not “carbon acid” Ln 16 p. 11705. Include “pCO<sub>2</sub>” before  
323 “values”

324 Reply / Change 5 – Modified in the revised MS, as suggested

325

326 Comment 6 – Ln 23- p. 11706-11707. The whole section about the pH determination sounds a bit  
327 vague (see below) and since pH is such a critical parameter in the calculations I think it needs to  
328 be improved. “Water temperature and pH were measured in the field with different probes  
329 depending on the origin of the dataset”. To measure pH is not easy, this is especially true for  
330 waters with low-ionic strength. In such waters specific electrodes are needed in order to receive a  
331 stable and accurate pH reading. It is not easy to judge how well this is considered or if it is a  
332 potential problem in the included systems from the existing pH description in the manuscript.  
333 From my own experience I often find field based pH meters more unreliable (even if well-  
334 calibrated) than lab based ones, however bringing the water to lab is associated with other  
335 uncertainty as the authors also mention.

336

337 Reply 6 – We totally agree that it is not easy to judge how well pH is measured in general. If this  
338 is true for our own dataset, it may be even more critical for pH data obtained by environmental  
339 agencies. Comments by the referee were based on his/her own experience, as in fact, no  
340 quantitative study that compares pH-meters, electrodes, field and lab measurements, storage  
341 impact, etc..., are available in the literature. In the discussion of the first version of our MS we  
342 wrote “Thus, one factor of variability throughout the dataset as well as in literature data is the  
343 accuracy of pH measurements – despite the care taken (e.g, calibrations with NBS buffers for  
344 each day of measurements), we cannot rule out that drift or malfunction of pH electrodes  
345 contribute to the observed variability, constituting an additional disadvantage compared to direct  
346 pCO<sub>2</sub> measurements with very stable gas analysers”.

347

348 Change 6 – in the revised MS, we stress more the difficulties of pH measuring, also in the method  
349 section of the paper, following the suggestions of the reviewer.

350

351 Comment 7 – Ln 6 p. 11707. I suggest to move the date to after “Rivers” Ln 16 p. 11707. What  
352 kind of IR gas analyser? Ln 21 p. 11708. Add “determination” after GC.

353 Reply/change 7 Modified in the revised MS as suggested.

354

355 Comment 8 – A general concern in the method section is the lack of info about the basic  
356 chemistry of the carbonate system. Below pH 5.4-5.6 there should in theory be no carbonate  
357 alkalinity. These limits (or similar) are also often used in studies where pCO<sub>2</sub> is calculated from  
358 alkalinity

359 Reply 8 – Alkalinity is operationally defined (in our case the GRAN function between pH 4 and 3,  
360 but also in the case of end-point titrations used in environmental agencies, titration can still be  
361 performed at pH of less than 5 (the pK of HCO<sub>3</sub><sup>-</sup>/CO<sub>2</sub> being around 4.5). In addition, as stated in  
362 our submitted MS, we used to aerate our sample before titration in order to remove a fraction of  
363 dissolved CO<sub>2</sub> before HCO<sub>3</sub><sup>-</sup> titration. This aeration increases the pH at the start of the titration  
364 and improves the quality (repeatability) of the alkalinity titration, even at pH around 5. Typical TA  
365 values in these cases are generally below 100 µmol L<sup>-1</sup>.

366 Change 8 – We provide more details on our TA protocol and mention these pH limits as  
367 suggested by the reviewer.

368

369 Comment 9 – Ln 2-4 p. 11710. The sentence that starts with “Our dataset. . .” Do not belong in  
370 the results as I see it. This is a reflection that is suitable for the discussion.

371 Reply / Change 9 – Sentence has been moved to discussion.

372

373 Comment 10 – Ln 1- p. 11711-. Here comes my main concern directly in the discussion. The  
374 influence of organic acids on the total alkalinity is not a new finding and there are ways to quantify  
375 and eventually compensate for the influence. In earlier acidification related research, for example  
376 Hruska et al. (2003) determined the organic acid concentration just by using pH and DOC as  
377 input parameters. This could then be removed from the total alkalinity (similar to the NCA derived  
378 from titration) in order to get a more reliable calculated pCO<sub>2</sub>. This procedure has also been used  
379 in earlier studies (see Humborg et al. 2010; Wallin et al. 2014). In addition, Wallin et al. (2014)  
380 presented a similar comparison as conducted in this manuscript where alkalinity based pCO<sub>2</sub> and  
381 pCO<sub>2</sub> derived from DIC measurements were compared. Despite considering the organic acids  
382 the comparison still showed rather large deviation in the low alkaline waters (highlighting the  
383 sensitivity of the carbonate system at low pH). Of course there might be problematic to use such  
384 a universal organic acid model over the wide range of included systems as in this study. But to  
385 not even mention this opening in the paper feels strange. I think this manuscript really quantify  
386 the problems but also has the potential to suggest solutions. I definitely agree though that an  
387 increased use of methods focusing on CO<sub>2</sub> determination are needed in the future with  
388 standardized protocol to follow.

389 Reply/changes 10 – see also reply 2 and the additional figure A

390 As in Wallin et al. 2014, we have used the tri-protic model for organic acids of Hruska et al.  
391 (2003) in order to estimate the alkalinity of organic acids. We subtracted this organic alkalinity to  
392 our measured TA and we re-calculated the pCO<sub>2</sub> from the pH and the corrected alkalinity. Water

393 pCO<sub>2</sub> corrected that way were not better correlated to the measured pCO<sub>2</sub>. To the contrary,  
394 some highly negative pCO<sub>2</sub> values appeared in the corrected dataset (Figure 1). In our revised  
395 MS, we describe these calculations in more details.  
396

397 Comment 11 – Ln 26-28 p. 11714. An evaluation of a pCO<sub>2</sub> method based on calculation from  
398 direct DIC determination was recently published this year (Åberg and Wallin, 2014). In this study  
399 the method was also compared to a direct headspace technique with good result.  
400 Reply/change 11 – we mention and refer to the results of the paper of Åberg and Wallin, 2014 in  
401 the revised MS, as suggested. However, we continue recommending in our MS field direct pCO<sub>2</sub>  
402 determination, as this is very precise, cheap and easy, and it does depend on the quality of pH  
403 measurements.  
404

405 Comment 12 – References :

406 Hruska, J., S. Köhler, H. Laudon, and K. Bishop (2003), Is a universal model of or- ganic acidity  
407 possible: Comparison of the acid/base properties of dissolved organic carbon in the boreal and  
408 temperate zones, *Environ. Sci. Technol.*, 37(9), 1726-1730, doi:10.1021/es0201552  
409 Humborg, C., C. M. Mörth, M. Sundbom, H. Borg, T. Blenckner, R. Giesler, and V. Ittekkot (2010),  
410 CO<sub>2</sub> supersaturation along the aquatic conduit in Swedish watersheds as constrained by  
411 terrestrial respiration, aquatic respiration and weathering, *Glob. Change Biol.*, 16(7), 1966-1978,  
412 doi:10.1111/j.1365-2486.2009.02092.x  
413 Wallin, M. B., S. Löfgren, M. Erlandsson, and K. Bishop (2014), Representative regional sampling  
414 of carbon dioxide and methane concentrations in hemiboreal headwa- ter streams reveal  
415 underestimates in less systematic approaches, *Glob. Biogeochem. Cycles*, 28(4), 465-479,  
416 doi:10.1002/2013gb004715  
417 Åberg, J., and M. B. Wallin (2014), Evaluating a fast headspace method for measuring DIC and  
418 subsequent calculation of pCO<sub>2</sub> in freshwater systems, *Inland Wat.*, 4(2), 157-166,  
419 doi:10.5268/IW-4.2.694  
420

421 Reply/change 12

422 We cited Wallin et al (2014), in the submitted MS

423 We added Hruska et al (2003), and Åberg, and Wallin (2014), and Humborg et al (2010) in the  
424 revised MS  
425

426 Anonymous Referee #3

427 Received and published: 27 August 2014

428 Comment 1 – I believe this is a well written and well structured manuscript on a topic of current  
429 and relevant interest to a wide community. The subject of methodological bias in calculated pCO<sub>2</sub>  
430 values is extremely important as more and more people are beginning to utilise large datasets to  
431 estimate GHG contributions from inland water systems. Whilst the concept of overestimation is  
432 not new, particularly in relation to DOC rich systems, this study presents a very nice dataset with  
433 which to test the hypotheses and explore links to other water characteristics. My comments  
434 below, which are few, are aimed primarily at improving what I believe is a very good manuscript in  
435 its current format.  
436

437 Reply 1 – We thank the referee for his/her very positive overall evaluation of our MS.  
438

439 Comment 2 – Introduction: whilst I realise there are practical limitations to the use of high  
440 temporal resolution sensors for in-situ CO<sub>2</sub> monitoring and therefore the number of studies are so  
441 far limited, it may be worth highlighting submerged NDIR sensors as an alternative direct method  
442 to measure pCO<sub>2</sub>.

443 Reply/change 2. We mention this membrane technique in our revised MS and cite the paper  
444 “Direct and continuous measurement of dissolved carbon dioxide in freshwater aquatic systems—  
445 method and applications by Johnson et al. (2010) in *ECOHYDROLOGY*.  
446

447 Comment 3 – Methods: I am generally happy with the methods section, I felt it was well written  
448 and sufficiently detailed using standard and reliable methods. Results section 3.2: I think more

449 could be made of the correlations that are described but not statistically modelled. Correlation  
450 statistics are not shown despite significant correlations being described.  
451 Reply/change 3 – correlation statistics (p values of linear regressions) are now provided in the  
452 revised MS, as suggested.

453  
454 Comment 4 – General Comment: Could simple empirical modelling be used to attempt a  
455 correction for the bias? Whilst this may not work across all systems it would be very interesting to  
456 try. The authors have shown an important overestimation in studies which utilise large datasets of  
457 water quality to indirectly calculate pCO<sub>2</sub> and subsequent evasion, they have not however  
458 attempted to produce a solution. If an additional calculation step, e.g. including a DOC correlation  
459 value, would allow better pCO<sub>2</sub> estimates from the available data that would be extremely useful  
460 and significantly improve the science. This may not be possible but I would like to see evidence it  
461 had been attempted.

462  
463 Reply 4 - Indeed, one crucial question raised by the three referees can be summarized as: "Is it  
464 possible to correct the bias in calculated pCO<sub>2</sub> data?". In fact, we tested several methods to  
465 derive empirical relationships that could be used to correct the pCO<sub>2</sub> calculated from pH and TA.  
466 Unfortunately, we found no reliable consistent quantitative relationships to allow correcting for the  
467 bias in pCO<sub>2</sub> when values are derived from pH, DOC, and TA.

468  
469 The first approach consisted in calculating organic alkalinity from pH and DOC using the models  
470 of Driscoll et al (1989) –which assumes a single apparent pK value for organic acids- and the  
471 model of Hruska et al. (2003) – a triprotic model which assumes three apparent pK values-.  
472 These two organic acid models applied to our data led to very similar organic alkalinity values  
473 (See attached figure 1A). The organic alkalinity was then subtracted from the TA and the pCO<sub>2</sub>  
474 was re-calculated from the measured pH and the TA value corrected from organic acids. pCO<sub>2</sub>  
475 values corrected that way were, however, still very different from those measured in the field (See  
476 figure 1B), being sometimes higher and sometimes lower than the measured values.

477  
478 The second approach consisted in subtracting from the measured TA, the alkalinity calculated  
479 with the CO<sub>2</sub>sys program using as input parameters the measured pH and pCO<sub>2</sub>, in order to  
480 derive a non-carbonate alkalinity (NCA). Besides the fact that NCA derived that way was often  
481 negative (probably due to large sensitivity of calculation from the pCO<sub>2</sub>/pH pairing), it was neither  
482 correlated with DOC, nor with pH (See Figure 2). Consequently, we could not derive any  
483 empirical relationship useful to correct for the bias in pCO<sub>2</sub> calculation.

484  
485 Change 4 – In the revised version of our MS, we include a description of these attempts to correct  
486 calculated pCO<sub>2</sub> from the available parameter (pH, TA, DOC and measured pCO<sub>2</sub>).

487  
488  
489  
490  
491 additional figure Captions

492  
493 Figure 1

494 A: comparison organic alkalinity calculated from pH and DOC using the models of Driscoll et al  
495 (1989) –which assumes a single apparent pK value for organic acids- and the triprotic model of  
496 Hruska et al. (2003) – which assumes three apparent pK values-. These two organic acid models  
497 applied to our data led to very similar organic alkalinity values (A).

498 B: this calculated organic alkalinity was then subtracted from the measured TA. Then the pCO<sub>2</sub>  
499 was re-calculated from the measured pH and the TA corrected from organic acids. pCO<sub>2</sub> values  
500 corrected that way were still very different from those measured in the field (B), although being  
501 sometime higher and sometimes lower than the measured values.

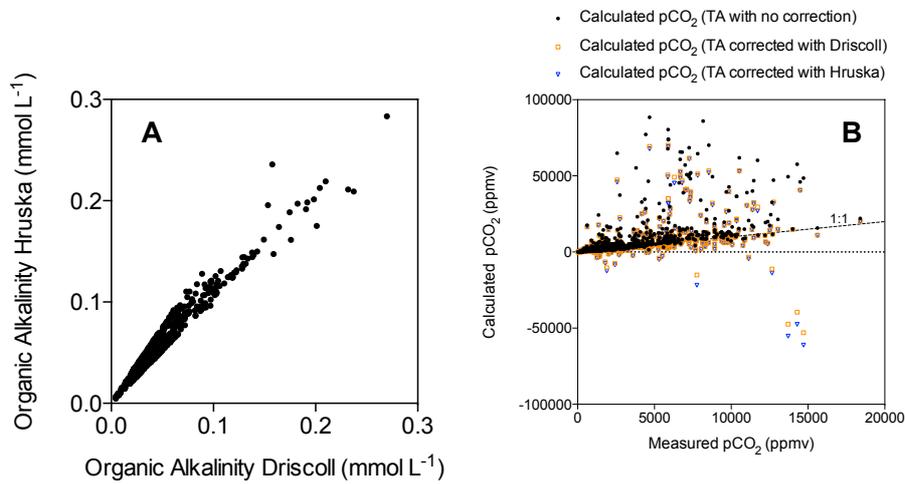
502  
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504 Figure 2

505 Non carbonate alkalinity (NCA), as the difference between the alkalinity calculated with the  
506 CO2sys program using pH and pCO<sub>2</sub> as input parameters, and the measured TA. NCA derived  
507 that way, expressed either in concentration (A,C) or as percentage of TA (B,D) was often  
508 negative and was neither (or weakly when expressed as % of TA) correlated with DOC, nor with  
509 pH. No quantitative empirical relationship could be deduced from these plots.  
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Figure 1

A: comparison organic alkalinity calculated from pH and DOC using the models of Driscoll et al (1989) –which assumes a single apparent pK value for organic acids- and the triprotic model of Hruska et al. (2003) – which assumes three apparent pK values-. These two organic acid models applied to our data led to very similar organic alkalinity values (A).  
B: this re-calculated organic alkalinity was then subtracted from the measured TA. Then the pCO<sub>2</sub> was re-calculated from the measured pH and the TA corrected from organic acids. pCO<sub>2</sub> values corrected that way were still very different from those measured in the field (B), although being sometime higher and sometimes lower than the measured values.

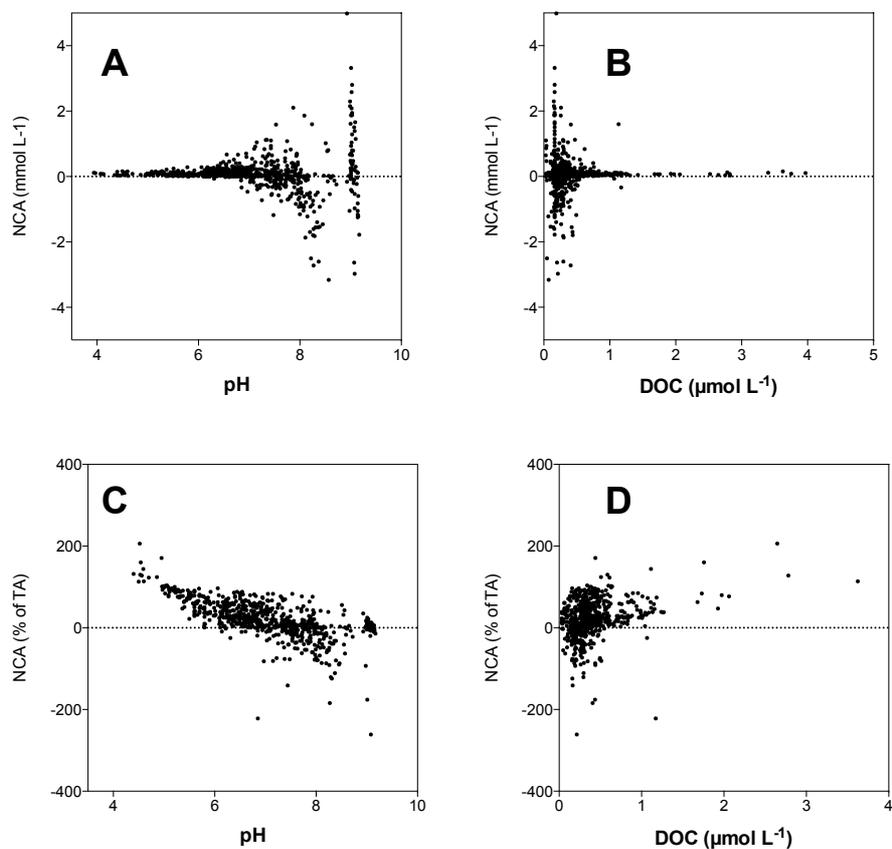


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Figure 2

Non carbonate alkalinity (NCA), as the difference between the alkalinity calculated with the CO2sys program using pH and pCO2 as input parameters, and the measured TA. NCA derived that way, expressed either in concentration (A,C) or as percentage of TA (B,D) was often negative and was neither (or weakly when expressed as % of TA) correlated with DOC, nor with pH. No quantitative empirical relationship could be deduced from these plots.



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543 **REVISED MS: MARKED-UP VERSION**

544

545 **Technical Note: Large overestimation of pCO<sub>2</sub> calculated from pH and alkalinity in**  
546 **acidic, organic-rich freshwaters**

547

548 Gwenaël Abril<sup>1,2</sup>, Steven Bouillon<sup>3</sup>, François Darchambeau<sup>4</sup>, Cristian R. Teodoru<sup>3</sup>, Trent R.

549 Marwick<sup>3</sup>, Fredrick Tamoo<sup>3</sup>, Fred Ochieng Omengo<sup>3</sup>, Naomi Geeraert<sup>3</sup>, Loris

550 Deirmendjian<sup>1</sup>, Pierre Polensaere<sup>1</sup>, and Alberto V. Borges<sup>4</sup>

551

552

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555

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557

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567 **Abstract**

568 Inland waters have been recognized as a significant source of carbon dioxide (CO<sub>2</sub>) to the  
569 atmosphere at the global scale. Fluxes of CO<sub>2</sub> between aquatic systems and the atmosphere  
570 are calculated from the gas transfer velocity and the water-air gradient of the partial pressure  
571 of CO<sub>2</sub> (pCO<sub>2</sub>). Currently, direct measurements of water pCO<sub>2</sub> remain scarce in freshwaters,  
572 and most published pCO<sub>2</sub> data are calculated from temperature, pH and total alkalinity (TA).  
573 Here, we compare calculated (pH and TA) and measured (equilibrator and headspace) water  
574 pCO<sub>2</sub> in a large array of temperate and tropical freshwaters. The 761 data points cover a wide  
575 range of values for TA (0 to 14200  $\mu\text{mol L}^{-1}$ ), pH (3.94 to 9.17), measured pCO<sub>2</sub> (36 to 23000  
576 ppmv), and dissolved organic carbon (DOC) (29 to 3970  $\mu\text{mol L}^{-1}$ ). Calculated pCO<sub>2</sub> were  
577 >10% higher than measured pCO<sub>2</sub> in 60% of the samples (with a median overestimation of  
578 calculated pCO<sub>2</sub> compared to measured pCO<sub>2</sub> of 2560 ppmv) and were >100% higher in the  
579 25% most organic-rich and acidic samples (with a median overestimation of 9080 ppmv). We  
580 suggest these large overestimations of calculated pCO<sub>2</sub> with respect to measured pCO<sub>2</sub> are  
581 due to the combination of two cumulative effects: (1) a more significant contribution of  
582 organic acids anions to TA in waters with low carbonate alkalinity and high DOC  
583 concentrations; (2) a lower buffering capacity of the carbonate system at low pH, that  
584 increases the sensitivity of calculated pCO<sub>2</sub> to TA in acidic and organic-rich waters. No  
585 empirical relationship could be derived from our dataset in order to correct calculated pCO<sub>2</sub>  
586 for this bias. Owing to the widespread distribution of acidic, organic-rich freshwaters, we  
587 conclude that regional and global estimates of CO<sub>2</sub> outgassing from freshwaters based on pH  
588 and TA data only are most likely overestimated, although the magnitude of the overestimation  
589 needs further quantitative analysis. Direct measurements of pCO<sub>2</sub> are recommended in inland  
590 waters in general, and in particular in acidic, poorly buffered, freshwaters.

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## 600 1 Introduction

601 Inland waters (streams, rivers, lakes, reservoirs, wetlands) receive carbon from terrestrial  
602 landscapes, usually have a net heterotrophic metabolism, and emit significant amounts of CO<sub>2</sub>  
603 to the atmosphere (Kempe 1984; Cole et al. 1994; Raymond et al. 2013). This terrestrial-  
604 aquatic-atmosphere link in the global carbon cycle is controlled by complex biogeographical  
605 drivers that generate strong spatial and temporal variations in the chemical composition of  
606 freshwaters and the intensity of CO<sub>2</sub> outgassing at the water-air interface (e.g. Tamooh et al.  
607 | 2013; Dinsmore et al. 2013; Abril et al. 2014; [Borges et al. 2014](#)). Hence, large datasets are  
608 necessary in order to describe the environmental factors controlling these CO<sub>2</sub> emissions and  
609 to quantify global CO<sub>2</sub> fluxes from inland waters (Sobek et al. 2005; Barros et al. 2011;  
610 Raymond et al. 2013). Dissolved inorganic carbon (DIC) concentration and speciation in  
611 freshwaters greatly depend on the lithological nature of watersheds (Meybeck 1987). For  
612 instance, rivers draining watersheds rich in carbonate rocks have a high DIC concentration,  
613 | generally well above 1000 μmol L<sup>-1</sup>. Bicarbonate ions contribute to the majority of the total  
614 alkalinity (TA) in these waters, which have high conductivities and high pH. In these hard  
615 waters, dissolved CO<sub>2</sub> represents a minor fraction (5-15%) of the DIC compared to  
616 bicarbonates. In rivers draining organic-rich soils and non-carbonate rocks, DIC  
617 concentrations are lower (typically a few hundred μmol L<sup>-1</sup>) but dissolved organic carbon  
618 | (DOC) concentrations are higher, and commonly exceed the DIC concentrations. Organic  
619 acid anions significantly contribute to TA of these soft waters (Driscoll et al. 1989; Hemond  
620 1990), which have low conductivities and low pH. Dissolved CO<sub>2</sub> represents a large,  
621 | generally dominant, fraction of DIC in these acidic, organic-rich waters.  
622

623 Fluxes of CO<sub>2</sub> between aquatic systems and the atmosphere can be computed from the water-  
624 air gradient of the concentration of CO<sub>2</sub> and the gas transfer velocity (Liss and Slater 1974) at  
625 local (e.g. Raymond et al. 1997), regional (e.g. Teodoru et al. 2009), and global scales (e.g.

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628 Cole et al. 1994; Raymond et al. 2013). The partial pressure of CO<sub>2</sub> (pCO<sub>2</sub>) is relatively  
629 constant in the atmosphere compared to surface freshwaters pCO<sub>2</sub> that can vary by more than  
630 4 orders of magnitude [spatially and temporally](#) (Sobek et al. 2005; Abril et al. 2014).  
631 Consequently, water pCO<sub>2</sub> controls ~~the~~ intensity of the air-water flux, together with the gas  
632 transfer velocity. [Presently](#), both measured and calculated water pCO<sub>2</sub> data are used to  
633 compute CO<sub>2</sub> fluxes from freshwater systems, although calculated pCO<sub>2</sub> is overwhelmingly  
634 more abundant than directly measured pCO<sub>2</sub> ([e.g. Cole et al. 1994; Raymond et al. 2013](#)).  
635 pCO<sub>2</sub> can be calculated from the dissociation constants of carbonic acid ([which are a function](#)  
636 of temperature) and any of the following couples of measured variables: pH/TA, pH/DIC,  
637 DIC/TA (Park 1969). In a majority of cases, calculated pCO<sub>2</sub> is based on the measurements of  
638 pH/TA and water temperature. These three parameters are routinely measured by many  
639 environmental agencies, and constitute a very large database available for the scientific  
640 community. Calculation of pCO<sub>2</sub> from pH and TA has been initiated in World Rivers in the  
641 1970's (Kempe 1984) and relies on the dissociation constants of carbonic acid, and the  
642 solubility of CO<sub>2</sub>, all of which are temperature-dependent ([Harned and Scholes 1941; Harned](#)  
643 [and Davis 1943](#); Millero 1979; Stumm and Morgan 1996). Measured pCO<sub>2</sub> is based on water-  
644 air phase equilibration either on discrete samples (headspace technique, e.g. Weiss 1981) or  
645 continuously (equilibrator technique, e.g. Frankignoulle et al. 2001) using various systems  
646 and devices, followed by direct, generally infrared (IR), detection of CO<sub>2</sub> in the equilibrated  
647 gas. Commercial IR gas analysers are becoming cheaper and more accurate, [stable and](#)  
648 [compact](#), and provide a large range of linear response well adapted to [variability of pCO<sub>2</sub>](#)  
649 [found in freshwaters](#).  
650  
651 Only a limited number of studies have compared directly measured pCO<sub>2</sub> ~~to~~ computed pCO<sub>2</sub>.  
652 [Earlier examples provided a](#) comparison between pCO<sub>2</sub> measured by headspace equilibration  
653 coupled to gas chromatography (GC), and pCO<sub>2</sub> calculated from pH and DIC ([Kratz et al.](#)

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660 | [1997](#); Raymond et al. 1997). Reports by these authors in [Wisconsin lakes and the Hudson](#)  
661 | River show that the pCO<sub>2</sub> values were linearly correlated but showed a variability of ±500  
662 | ppmv around the 1:1 line, over a range of measured pCO<sub>2</sub> from 300 to 4000 ppmv. Later,  
663 | Frankignoulle and Borges (2001) reported the first comparison of pCO<sub>2</sub> calculated from pH  
664 | and TA and pCO<sub>2</sub> measured by equilibration coupled to an IR analyzer in an estuary in  
665 | Belgium. In this high TA (2500-4800 μmol L<sup>-1</sup>) and high pH (>7.4) system, they found a  
666 | good agreement between both approaches, calculated pCO<sub>2</sub> being either overestimated or  
667 | underestimated, but always by less than 7%. In 2003, concomitant measurements of pH, TA  
668 | and pCO<sub>2</sub> were performed in acidic, humic-rich ("black" type) waters of the Sinnamary River  
669 | in French Guiana (Abril et al. 2005; 2006). Calculation of pCO<sub>2</sub> from pH (~5) and TA (~200  
670 | μmol L<sup>-1</sup>) gave unrealistically high values compared to those measured directly with a  
671 | headspace technique (typically 30000 ppmv versus 5000 ppmv). Direct measurements of CO<sub>2</sub>  
672 | and CH<sub>4</sub> outgassing fluxes with floating chambers and the computation of the respective gas  
673 | transfer velocities of these two gases (Guérin et al. 2007) confirmed that pCO<sub>2</sub> values  
674 | calculated from pH and TA were overestimated compared to direct measurements in the  
675 | Sinnamary River. More recently, Hunt et al. (2011) and Wang et al. (2013) provided evidence  
676 | that organic acid anions in DOC may significantly contribute to TA in some rivers and  
677 | generate an overestimation of calculated pCO<sub>2</sub>. Butman and Raymond (2011) reported higher  
678 | calculated than measured pCO<sub>2</sub> in some US streams and rivers, but no information was  
679 | available on the potential role of organic acids on this overestimation. These authors  
680 | concluded that the low number of samples in their study reflected the need for more research  
681 | on this topic.

682 |  
683 | With the growing interest on pCO<sub>2</sub> determination in freshwaters globally, and given the  
684 | apparent simplicity and low cost of pH and TA measurements, the number of publications  
685 | that report calculated pCO<sub>2</sub> in freshwaters has increased dramatically in the last decade. Some

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702 | of these publications report extremely high and potentially biased  $p\text{CO}_2$  values in low-  
703 | alkalinity and high DOC systems. It has thus become necessary to pay attention to this issue  
704 | and investigate the occurrence of such potential bias and its magnitude in the different types  
705 | of freshwaters. Here, we present a large dataset of concomitant measurements of temperature,  
706 | pH, TA,  $p\text{CO}_2$ , and DOC in freshwaters. This is the first comprehensive dataset to investigate  
707 | the magnitude of the bias between calculated and measured  $p\text{CO}_2$ , as it covers the entire range  
708 | of variation of most parameters of the carbonate system in freshwaters. The objective of this  
709 | paper is to alert the scientific community to the occurrence of a bias in  $p\text{CO}_2$  calculation from  
710 | pH and TA in acidic, poorly buffered and organic-rich freshwaters, to briefly discuss its origin  
711 | in terms of water chemistry, and to provide the range of pH, TA and DOC values where  $p\text{CO}_2$   
712 | calculation should be abandoned and the range where it still gives relatively accurate results.

713

## 714 | **2 Material and Methods**

### 715 | 2.1 Sample collection

716 | Our dataset consists of 761 concomitant measurements of temperature, pH, TA, water  $p\text{CO}_2$ ,  
717 | and DOC in 12 contrasting tropical and temperate systems in Europe, Amazonia and Africa  
718 | ([Fig. 1](#); Table 1). These samples were obtained in the Central Amazon River and floodplains  
719 | system in Brazil, the Athi-Galana-Sabaki River in Kenya, the Tana River (Kenya), small  
720 | rivers draining the Shimba Hills in southeastern Kenya, the Congo River and tributaries in the  
721 | Democratic Republic of [the Congo \(DRC\)](#), Lake Kivu in Rwanda and DRC, the Leyre River  
722 | and tributaries in France, the Loire River in France, the Meuse River in Belgium, the Rianila  
723 | and Betsiboka Rivers in Madagascar, the Sinnamary River downstream of the Petit Saut  
724 | Reservoir in French Guiana, and the Zambezi River in Zambia and Mozambique ([Fig. 1](#)).  
725 | Details on some of the sampling sites can be found in Abril et al. (2005; 2014), Borges et al.  
726 | (2012 ; 2014), Marwick et al. (2014a,b), Polsenaere et al. (2013), Tamooh et al. (2013),  
727 | [Teodoru et al. \(2014\)](#). These watersheds span a range of climates and are occupied by

728 different types of land cover, which include tropical rainforest (Amazon, Congo, Rianila), dry  
729 savannah (Tana, Athi-Galana-Sabaki, Betsiboka, [Zambezi](#)), temperate pine forest growing on  
730 podzols (Leyre), mixed temperate forest, grassland and cropland (Meuse) and cropland  
731 (Loire). Lithology is also extremely contrasted as it includes for instance carbonate-rocks  
732 dominated watershed as for the Meuse, sandstone dominated silicates (Leyre), and  
733 precambrian crystalline magmatic and metamorphic rocks with a small proportion of  
734 carbonate and evaporite rocks for the Congo river.

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## 736 2.2 Field and laboratory measurements

737 Although pH measurements might seem almost trivial, highly accurate and precise pH data  
738 are in fact not easy to obtain, especially in low-ionic strength waters, where electrode readings  
739 are generally less stable. Even though pH measurements in the laboratory might be more  
740 accurate, it is crucial to measure pH *in situ* or immediately after sampling, as pH  
741 determination several hours or days after sampling will be affected by CO<sub>2</sub> degassing and/or  
742 microbial respiration (Frankignoulle and Borges 2001). In this work, water temperature and  
743 pH were measured in the field with different probes depending on the origin of the dataset.  
744 However, all the pH data were obtained with glass electrodes and rely on daily calibration  
745 with two-point United States National Bureau of Standards (NBS) standards (4 and 7).  
746 Measurements were performed directly in the surface water, or in collected water  
747 immediately after sampling.

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748  
749 Several techniques were used to measure water pCO<sub>2</sub>. Water-gas equilibration was performed  
750 with a marbles type equilibrator (Frankignoulle et al. 2001) for the Amazon, Loire, Leyre,  
751 Sinnamary and Congo [Rivers](#) (December 2013) as well for Lake Kivu, or with a Liqui-Cel  
752 MiniModule membrane contactor equilibrator (see Teodoru et al. 2009; [2014](#)) for the  
753 Zambezi and some sites within the Congo basin (December 2012): water was pumped either

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768 continuously from a ship, or on an ad hoc basis from the bank of the rivers after waiting ~15  
769 minutes for complete equilibration; air was continuously pumped from the equilibrator to the  
770 gas analyser (see *e.g.* Abril et al. 2014 for a more detailed description of the system). A  
771 syringe-headspace technique (Kratz et al. 1997; Teodoru et al. 2009) was used **in the field** in  
772 all African rivers and in the Meuse River: 30 mL volume of atmospheric air was equilibrated  
773 with 30 mL volume of river water by vigorously shaking during 5-10 min in four replicate  
774 gas-tight syringes. The four replicates 30 mL of equilibrated gas and a sample of atmospheric  
775 air were injected in an IR gas analyser (**Li-Cor® models 820 or 840, or PP systems® model**  
776 **EGM-4**); the first gas injection served as a purge for the **air** circuit and cell and the three other  
777 injections were used as triplicate pCO<sub>2</sub> determination (average repeatability of ±1%). The  
778 pCO<sub>2</sub> in the river water was deduced from that measured in the headspace accounting for the  
779 initial pCO<sub>2</sub> in the air used for equilibration, water temperature in the river and in the water at  
780 equilibrium in the syringe, and based on Henry's law. Comparison between syringe-  
781 headspace and marbles or membrane equilibrator was made during two cruises on the Congo  
782 River and three cruises in the Zambezi basin and gave very consistent results, deviation from  
783 the 1:1 line being always less than 15% (see **Fig. 2**). **This highlights the consistency of the**  
784 **present data-set of direct pCO<sub>2</sub> measurements although different techniques were used.** A  
785 serum bottle-headspace technique (Hope et al. 1995) was also used on the Sinnamary River;  
786 surface water was sampled in 120 mL serum bottles that were **poisoned with HgCl<sub>2</sub> and**  
787 **sealed excluding air bubbles.** Back in the laboratory, a **40 mL** headspace was created with  
788 pure N<sub>2</sub> (Abril et al. 2005). The CO<sub>2</sub> concentration of equilibrated gas in the headspace was  
789 analysed by injecting small volumes (0.5 mL) of gas in a GC calibrated with certified gas  
790 mixtures.  
791  
792 Immediately after water-gas phase equilibration, CO<sub>2</sub> was detected and quantified in most  
793 samples with non-dispersive **IR** gas analysers (Frankignoulle et al. 2001; Abril et al. 2014).

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801 The gas analysers were calibrated before each field cruise, with air circulating through soda  
802 lime or pure N<sub>2</sub> for zero and with a certified gas standard for the span. Depending on the  
803 cruises and expected pCO<sub>2</sub> ranges, we used gas standard concentration of 1000 - 2000 ppmv,  
804 or a set of calibration gases at 400, 800, 4000 and 8000 ppmv. Stability of the instrument was  
805 checked after the cruise, and deviation of the signal was always less than 5%. These  
806 instruments offer a large range of linear response, depending on manufacturers and model: 0-  
807 20000 ppmv or 0-60000 ppmv. The linearity of a Li-COR® Li-820 gas analyser was verified  
808 by connecting it to a closed circuit of gas equipped with a rubber septum to allow injection of  
809 pure CO<sub>2</sub> with a syringe. Linearity was check by injecting increasing volumes of CO<sub>2</sub> in order  
810 to cover the whole range of measurement and was excellent between zero and ~20000 ppmv.  
811 In addition to the JR analysers generally used in this work, in the Sinnamary River, pCO<sub>2</sub> was  
812 also measured with an INNOVA® 1312 optical filter IR photoacoustic gas analyser (range 0  
813 - 25000 ppmv) connected to an equilibrator and with a Hewlett Packard® 5890 GC equipped  
814 with a thermal conductivity detector (TCD); both analysers were calibrated with a gas mixture  
815 of 5000 ppmv of CO<sub>2</sub>. Both methods gave results consistent at ± 15% in the 0 – 13000 ppmv  
816 range (Abril et al. 2006). Sinnamary data reported here are from headspace and GC  
817 determination.  
818  
819 TA was analysed by automated electro-titration on 50 ml filtered samples with 0.1N HCl as  
820 titrant. Equivalence point was determined with a Gran method from pH between 4 and 3  
821 (Gran 1952). Precision based on replicate analyses was better than ±5 µmol L<sup>-1</sup>. TA  
822 measurements should be done on filtered samples; otherwise some overestimation would  
823 occur in turbid samples, which may content significant amount of acid-neutralizing particles  
824 (e.g. calcium carbonate). In contrast to TA measurements based on titration to an endpoint of  
825 5.6 (e.g. Wallin et al. 2014), the Gran titration method allows the determination of TA values  
826 in samples with situ pH down to ~4.5, i.e. very close to the dissociation constant of HCO<sub>3</sub><sup>-</sup>

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836 /H<sub>2</sub>CO<sub>3</sub>. In most acidic samples with low TA, reproducibility was improved by slightly  
837 increasing the pH by up to 0.2 units by vigorously stirring during ~15min in order to degas as  
838 much CO<sub>2</sub> as possible before starting the titration. DOC was measured on samples filtered  
839 through pre-combusted (490°C) glass fibre filter with a porosity of 0.7 µm and stored  
840 acidified with ultrapure H<sub>3</sub>PO<sub>4</sub> in borosilicate vials capped with polytetrafluoroethylene  
841 stoppers. Analysis was performed with a Shimadzu TOC5000 analyser based on high  
842 temperature catalytic oxidation, after removal of dissolved CO<sub>2</sub> for samples from Amazon,  
843 Loire, Leyre, and Sinnamary Rivers. DOC concentrations were measured with a customized  
844 wet oxidation TOC analyzer (Thermo HiperTOC, or IO Analytical Aurora 1030W) coupled to  
845 a Delta+XL or Delta V IRMS.

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846  
847 2.3 pCO<sub>2</sub> calculation from pH and TA  
848 We calculated pCO<sub>2</sub> from TA, pH and temperature measurements using carbonic acid  
849 dissociation constants of Millero (1979) (based on those of Harned and Scholes (1941) and  
850 Harned and Davis (1943)) and the CO<sub>2</sub> solubility from Weiss (1974) as implemented in the  
851 CO2SYS program. Hunt et al. (2011) reported discrepancy lower than 2% for pCO<sub>2</sub> computed  
852 this way with those obtained with the PHREEQC program (Parkhurst and Appelo, 1999).  
853 Differences in software or dissociation constants cannot account for the large bias in  
854 calculated pCO<sub>2</sub> compared to measured pCO<sub>2</sub> we report in this paper.

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### 856 3 Results

#### 857 3.1 Data ranges and patterns in the entire dataset

858 Measured pCO<sub>2</sub> varied between 36 ppmv in a floodplain of the Amazon River and 23000

859 ppmv in a first order stream of the Leyre River (Table 1). Minimum values of pH and TA

860 occurred in the Congo River (pH = 3.94 and TA = 0) and maximum values in Lake Kivu (pH

861 = 9.16 and TA = 14200 µmol L<sup>-1</sup>). Highest DOC concentrations (>3000 µmol L<sup>-1</sup>) were

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878 observed in small streams in the Congo basin and in first order streams draining podzolized  
 879 soils in the Leyre basin. Lowest DOC concentrations ( $<40 \mu\text{mol L}^{-1}$ ) occurred in some  
 880 tributaries of the Athi-Galana-Sabaki, in the Rianila and Betsiboka Rivers, and in the Shimba  
 881 Hills streams. When considering the whole dataset, measured  $\text{pCO}_2$  and DOC were negatively  
 882 correlated with  $\text{pH}$ , whereas TA was positively correlated with  $\text{pH}$  (Fig. 3,  $p < 0.0001$  for the  
 883 three variables). This illustrates the large contrast in acid-base properties between acidic,  
 884 organic-rich and poorly buffered samples on the one hand, and basic, carbonate-buffered  
 885 samples on the other hand.

### 887 3.2 Comparison between measured and calculated $\text{pCO}_2$

888 Calculated  $\text{pCO}_2$  was more than 10% lower than measured  $\text{pCO}_2$  in 16% of the samples; both  
 889 methods were consistent at  $\pm 10\%$  in 24% of the samples; calculated  $\text{pCO}_2$  was more than  
 890 10% higher than measured  $\text{pCO}_2$  in 60% of the samples and more than 100% higher in 26%  
 891 of the samples. Absolute values, as expressed in ppmv, were largely shifted toward  
 892 overestimation, calculated *versus* measured  $\text{pCO}_2$  data being well above the 1:1 line, and  
 893 calculated minus measured  $\text{pCO}_2$  values ranging between -6180 ppmv and +882022 ppmv  
 894 (Fig. 4). The largest overestimation of calculated  $\text{pCO}_2$  occurred in the most acidic samples,  
 895 whereas underestimations of calculated  $\text{pCO}_2$  occurred in neutral or slightly basic samples  
 896 (Fig. 4b). Ranking the data according to the  $\text{pH}$ , TA and DOC reveal that overestimation of  
 897 calculated  $\text{pCO}_2$  compared to measured  $\text{pCO}_2$  increased in acidic, poorly buffered waters in  
 898 parallel with an increase in the DOC concentration (Table 2). Discrepancies between  
 899 calculated and measured  $\text{pCO}_2$  were very different from one system to another, depending on  
 900 the chemical status of the waters. On average at each sampled site, the relative overestimation  
 901 of calculated  $\text{pCO}_2$  decreased with  $\text{pH}$  and TA and increased with DOC (Fig. 5).

902 Overestimation of calculated  $\text{pCO}_2$  was on average  $<10\%$  in the Kivu Lake, and the Meuse,  
 903 Loire, Shimba Hills and Tana Rivers, which all have neutral or basic  $\text{pH}$ ,  $\text{TA} > 1000 \mu\text{mol L}^{-1}$

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**Déplacé vers le bas [1]:** Our dataset probably covers the full range of conditions of carbon speciation that can be encountered in continental surface waters.
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925 | and low to moderate DOC concentrations ( $< 400 \mu\text{mol L}^{-1}$ ) (Fig. 5). In contrast, calculated  
926 |  $\text{pCO}_2$  was overestimated by  $>200\%$  on average in the Congo, Leyre, Sinnamary and Amazon  
927 | Rivers, which have acidic pH, TA  $< 500 \mu\text{mol L}^{-1}$  and highest DOC concentration, reaching  
928 |  $1000 \mu\text{mol L}^{-1}$  on average in the Congo. The cases of Athi-Galana-Sabaki, Rianila, Betsiboka,  
929 | and Zambezi Rivers were intermediate in pH, TA and DOC, and with average  
930 | overestimations of calculated  $\text{pCO}_2$  of 50-90% (Fig. 5).

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## 932 | 4 Discussion

### 933 | 4.1 Origin of overestimation of calculated $\text{pCO}_2$

934 | Our dataset (Fig. 3; Table 1) probably covers the full range of conditions of carbon speciation  
935 | that can be encountered in continental surface waters. A  $\text{pCO}_2$  overestimation negatively

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936 | correlated with pH ( $p = 0.001$ ) and TA ( $p = 0.005$ ) and positively correlated with DOC ( $p <$   
937 |  $0.001$ ) (Fig. 5) is consistent with the observations of Cai et al. (1998) in the freshwater end-

938 | members of some estuaries in Georgia, USA, and of Hunt et al. (2011) in rivers in New

939 | England (USA) and New Brunswick (CA). These authors performed NaOH back-titration in

940 | order to measure non-carbonate alkalinity (NCA). They found that NCA accounted for a large

941 | fraction (in some cases the majority) of TA; in addition, the contribution of inorganic species

942 | other than carbonate was assumed negligible and most of the NCA was attributed to organic

943 | acid anions. Hunt et al. (2011) also showed that in the absence of direct titration of NCA,

944 | which is labour-intensive and whose precision may be poor, this parameter could be

945 | calculated as the difference between the measured TA and the alkalinity calculated from

946 | measurements of pH and DIC and the dissociation constants of carbonic acid. Using the latter

947 | approach, Wang et al. (2013) obtained a positive correlation between NCA and DOC

948 | concentrations in the Congo River, evidencing the predominant role of organic acids in DIC

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949 | speciation and pH in such acidic system. Because we did not directly measure DIC in this

950 | study, we could not calculate NCA with the same procedure as these studies. We attempted to

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968 calculate TA from our measured pH and pCO<sub>2</sub> with the CO2SYS program. However, TA  
 969 values calculated this way were inconsistent with other measured variables (with sometimes  
 970 negative values). Indeed, because pH and pCO<sub>2</sub> are too interdependent in the carbonate  
 971 system, very small analytical errors on these variables lead to large uncertainties in the  
 972 calculated TA (Cullison Gray et al. 2011). A second attempt to correct our TA data from  
 973 NCA consisted in calculating organic alkalinity using pH and DOC as input parameters. We  
 974 compared the model of Driscoll et al. (1989), which assumes a single pK value for all organic  
 975 acids, and the triprotic model of Hruska et al. (2003), which assumes three apparent pK values  
 976 for organic acids. These two models applied to our pH and DOC gave very similar organic  
 977 alkalinity values, which could be subtracted from the measured TA. In the most acidic  
 978 samples (e.g. some sites from the Congo basin), modelled organic alkalinities were larger than  
 979 measured TA and the difference was thus negative. Nevertheless, we then recalculated pCO<sub>2</sub>  
 980 from the measured pH and the TA corrected from organic alkalinity. Calculated pCO<sub>2</sub>  
 981 corrected with that method were, however, still very different from those measured in the  
 982 field, being sometimes higher and sometimes lower than the measured pCO<sub>2</sub>, without any  
 983 meaningful pattern (indeed, corrected pCO<sub>2</sub> was negatively correlated ( $p < 0.001$ ) with  
 984 measured pCO<sub>2</sub>). Consequently, we were unable to derive any empirical relationship to  
 985 correct for the bias in pCO<sub>2</sub> calculation from pH and TA. Nevertheless, the negative  
 986 correlation between pH and DOC and positive correlation between pH and TA (Fig. 3)  
 987 confirm a strong control of organic acids on pH and DIC speciation across the entire dataset.  
 988  
 989 As discussed by Hunt et al. (2011), a significant contribution of organic acids to TA leads to  
 990 an overestimation of calculated pCO<sub>2</sub> with the CO2SYS program, or with any program that  
 991 accounts only for the inorganic species that contribute to TA. It is thus obvious that the  
 992 observed increase in pCO<sub>2</sub> overestimation when pH decreases (Fig. 4b & 5; Table 2) is due to  
 993 an increasing contribution of organic acid anions to TA. However, this effect is not the only

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1008 driver of the observed overestimation of pCO<sub>2</sub>, which is also due to a decrease in the  
1009 buffering capacity of the carbonate system at acidic pH. To investigate the magnitude of this  
1010 second effect, we calculated the factor dpCO<sub>2</sub>/dTA (in ppmv mol<sup>-1</sup>), which describes the  
1011 change in calculated pCO<sub>2</sub> induced by a change in TA. This factor, which is the opposite of a  
1012 buffer factor as it reflects the sensitivity of pCO<sub>2</sub> calculation to the TA, increases  
1013 exponentially when pH decreases (Fig. 6a), i.e. it is proportional to the H<sup>+</sup> concentration. To  
1014 go further in this theoretical analysis, we computed the difference between the pCO<sub>2</sub>  
1015 calculated at a given TA value and the one calculated at a slightly higher TA value (TA+ X  
1016 μmol L<sup>-1</sup>). These calculations reveal an extreme sensitivity of calculated pCO<sub>2</sub> to TA at acidic  
1017 pH (Figure 5b). For instance, increasing TA by 5 μmol L<sup>-1</sup> (a value close to the precision of  
1018 TA titrations) increases the calculated pCO<sub>2</sub> by 31 ppmv at pH 7, by 307 ppmv at pH 6 and by  
1019 3070 at pH 5. Increasing TA by 100 μmol L<sup>-1</sup> (a typical value of NCA found in freshwaters,  
1020 Driscoll et al. 1994; Cai et al. 1998; Hunt et al. 2011), increases the calculated pCO<sub>2</sub> by 615  
1021 ppmv at pH 7, by 6156 ppmv at pH 6 and by 61560 ppmv at pH 5. Note that this increase in  
1022 calculated pCO<sub>2</sub> is independent of the chosen initial TA value. The difference between  
1023 calculated and measured pCO<sub>2</sub> from our data-set, shows that a NCA contribution around 100  
1024 μmol L<sup>-1</sup> is sufficient to explain the overestimation of calculated pCO<sub>2</sub> of most samples at  
1025 pH<6, whereas a NCA contribution higher than 500 μmol L<sup>-1</sup> would be necessary for several  
1026 samples at circumneutral and slightly basic pH (Fig. 5b). Samples requiring this high NCA  
1027 contribution are from the Athi-Galana-Sabaki and Zambezi watersheds, and correspond to TA  
1028 values well above 1000 μmol L<sup>-1</sup>. A NCA value of 500 μmol L<sup>-1</sup> in these samples is thus  
1029 plausible.  
1030  
1031 We have no definitive explanation for lower calculated than measured pCO<sub>2</sub>, which is  
1032 observed mainly at neutral to slightly basic pH, e.g., in the Zambezi River (Fig. 4). In most of  
1033 these samples, owing to the relatively high TA value, an overestimation of pH of less than 0.2

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1043 units is sufficient to account for the low calculated pCO<sub>2</sub> compared to measured values. In  
1044 general, it is not easy to judge how accurate are pH measurements, and even less when data  
1045 come from environmental agencies. Thus, one factor of variability throughout the dataset as  
1046 well as in literature data is the accuracy of pH measurements – despite the care taken (e.g,  
1047 calibrations with NBS buffers for each day of measurements), we cannot rule out that drift or  
1048 malfunction of pH electrodes contribute to the observed variability, constituting an additional  
1049 disadvantage compared to direct pCO<sub>2</sub> measurements with very stable gas analysers.

1050

#### 1051 4.2 Impact on estimates of CO<sub>2</sub> emissions from freshwaters

1052 According to our analysis, overestimation of calculated pCO<sub>2</sub> is largest in acidic, poorly  
1053 buffered and organic-rich waters. Consequently, the overestimation of regional and global  
1054 CO<sub>2</sub> emissions computed from calculated pCO<sub>2</sub> depends on the relative contribution of these  
1055 types of waters worldwide. In their analysis, Raymond et al. (2013) have discarded all  
1056 calculated pCO<sub>2</sub> values with a pH value of less than 5.4, as well as all pCO<sub>2</sub> values above  
1057 100000 ppmv. These criteria would exclude only 8% of samples from our dataset. Indeed,  
1058 from our analysis, it appears that overestimation of calculated pCO<sub>2</sub> occurs at pH much higher  
1059 than 5.4 (Fig. 4,5&6; Table 2). Both techniques were consistent at ±10% on average in only  
1060 five of the twelve studied systems, which combine a circumneutral to basic pH with a TA  
1061 concentration well above 1000 μmol L<sup>-1</sup> (Fig. 5). Although it would not be sufficient for the  
1062 cases of the Zambezi and Athi-Galana-Sabaki Rivers where overestimation is still significant,  
1063 a TA value above 1000 μmol L<sup>-1</sup> appears as a more robust criteria, rather than a pH threshold,  
1064 to separate calculated pCO<sub>2</sub> affected by bias from those consistent with measured pCO<sub>2</sub>  
1065 (Table 2). In fact, pCO<sub>2</sub> calculation from pH and TA in freshwaters historically relies on  
1066 theoretical background and validation data in high alkalinity waters (Neal et al. 1998),  
1067 including karstic waters (Kempe 1975). At the global scale, high TA typically occurs in rivers  
1068 draining watersheds with a significant proportion of carbonate rocks, typically >30% of their

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1073 surface area if the criteria of TA > 1000  $\mu\text{mol L}^{-1}$  is chosen and the normalized weathering  
1074 rates of Meybeck (1987) are applied. According to Meybeck (1987), the average and  
1075 discharge-weighted TA is around 900  $\mu\text{mol L}^{-1}$  for world rivers and around 600  $\mu\text{mol L}^{-1}$  for  
1076 tropical rivers. [Among the 25 largest rivers in the world, 15 have a TA > 1000  \$\mu\text{mol L}^{-1}\$](#)   
1077 [according to Cai et al. \(2008\)](#). The two largest rivers in the world in terms of discharge, the  
1078 Amazon and the Congo, are also well below this limit of 1000  $\mu\text{mol L}^{-1}$  and have large  
1079 overestimation in calculated pCO<sub>2</sub> (on average 200% and 360%, respectively). Very low TA  
1080 and pH and high DOC values have also been reported in boreal streams and rivers (Humborg  
1081 et al. 2010; Dinsmore et al. 2012; Wallin et al. 2014).

1083 In lakes, the highest pCO<sub>2</sub> values in the literature [come from tropical black water lakes and](#)  
1084 [were](#) also calculated rather than directly measured (Sobek et al. 2005). [Calculated pCO<sub>2</sub> was](#)  
1085 [65250 ppmv in Lago Tupé in the Brazilian Amazon, a Ria Lake connected to the Rio Negro,](#)  
1086 [where, according to our own dataset, pH is below 5 and TA is around 70  \$\mu\text{mol L}^{-1}\$ . It was,](#)  
1087 [18950 ppmv in Kambanain Lake in Papua New Guinea corresponding to a pH value of 6.1](#)  
1088 [and a TA value of 350  \$\mu\text{mol L}^{-1}\$  \(Vyverman 1994\).](#) This suggests a widespread overestimation  
1089 of calculated pCO<sub>2</sub> that significantly impacts the estimation of global CO<sub>2</sub> emissions from  
1090 inland waters. However, a precise analysis based on exact quantitative information on the  
1091 relative contribution of acidic and high and low alkalinity waters to the total surface area of  
1092 inland waters is necessary in order to evaluate the exact magnitude of the overestimation.

## 1093

### 1094 5 Conclusion

1095 From our analysis, it appears that the validity of calculating pCO<sub>2</sub> from pH, TA, and  
1096 temperature is most robust in freshwaters with circumneutral to basic pH and with TA  
1097 exceeding 1000  $\mu\text{mol L}^{-1}$ . At lower TA and pH, however, calculated pCO<sub>2</sub> (and hence, CO<sub>2</sub>  
1098 degassing rates) are overestimated by 50 to 300% relative to direct, *in situ* pCO<sub>2</sub>

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1111 measurements. Since a large majority of freshwater systems globally have characteristics  
1112 outside of the range of applicability of pCO<sub>2</sub> calculation, it appears reasonable to assume that  
1113 recent estimates of global CO<sub>2</sub> emission from lakes and rivers, which are based exclusively on  
1114 calculated pCO<sub>2</sub> data, are overestimated. We propose that while TA and pH measurements  
1115 remain useful to describe the aquatic chemistry, data on pCO<sub>2</sub> should in the future rely on  
1116 direct measurements of pCO<sub>2</sub>. Even if some studies report relatively robust calculation of  
1117 pCO<sub>2</sub> from pH and DIC measurements (Raymond et al. 1997; Kratz et al. 1997; Aberg and  
1118 Wallin 2014), direct pCO<sub>2</sub> in the field are stable, precise, and straightforward and do not  
1119 depend on the quality of pH measurements, which are often uncertain. Further, high-quality  
1120 DIC measurements are very time consuming, fairly complicated to set-up and do not allow to  
1121 carry out continuous measurements in a simple and straightforward fashion. Although there  
1122 are some practical limitations to their use in the field, submerged IR sensors, which allow  
1123 high temporal resolution, are also promising (Johnson et al. 2010). The long-term instrument  
1124 stability and accuracy based on newly developed Off-Axis Integrated Cavity Output  
1125 Spectroscopy and Cavity Ring-Down Spectroscopy technologies seems to improve in  
1126 comparison to traditional IR instruments, although the latter are more affordable, compact and  
1127 with lower power requirements. Joint international efforts are necessary to define most  
1128 appropriate protocols for the measurements of dissolved inorganic carbon parameters in  
1129 freshwaters.

1130  
1131

### 1132 *Acknowledgements*

1133 The dataset used in this study was collected in the framework of projects funded by the  
1134 Cluster of Excellence COTE at the Université de Bordeaux (ANR-10-LABX-45, CNP-Leyre  
1135 project), the European Research Council (ERC-StG 240002, AFRIVAL, African river basins:  
1136 Catchment-scale carbon fluxes and transformations, <http://ees.kuleuven.be/project/afriaval/>),

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**Supprimé:** pCO<sub>2</sub>, DIC and pH

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1146 | the Fonds National de la Recherche Scientifique (FNRS, CAKI, 2.4.598.07, [TransCongo](#),  
1147 | 14711103), the French national agency for research (ANR 08-BLAN-0221, CARBAMA  
1148 | project <http://carbama.epoc.u-bordeaux1.fr/>), the Research Foundation Flanders (FWO-  
1149 | Vlaanderen), the Belgian Federal Science Policy (BELSPO-SSD projects [COBAFISH](#) and  
1150 | [EAGLES](#)), the Research Council of the KU Leuven, and the Institut de Radioprotection et  
1151 | Sureté Nucléaire, France (FLORE project). We thank the Hydreco Laboratory in French  
1152 | Guiana, and Patrick Albéric (ISTO Orléans) who analysed some of the data reported here,  
1153 | Aurore Beulen (ULg) for collection of Meuse dataset, Marc-Vincent Commarieu ([ULg](#)) for  
1154 | analytical support, [two anonymous reviewers and C.W. Hunt \(reviewer\) for constructive](#)  
1155 | [comments on the previous version of the paper](#). AVB is a senior research [associate](#) at the  
1156 | FNRS.

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**Commentaire [1]:**

The volume and pages should be available soon (paper is submitted and BGD in production), something to look out for the proofing stage.

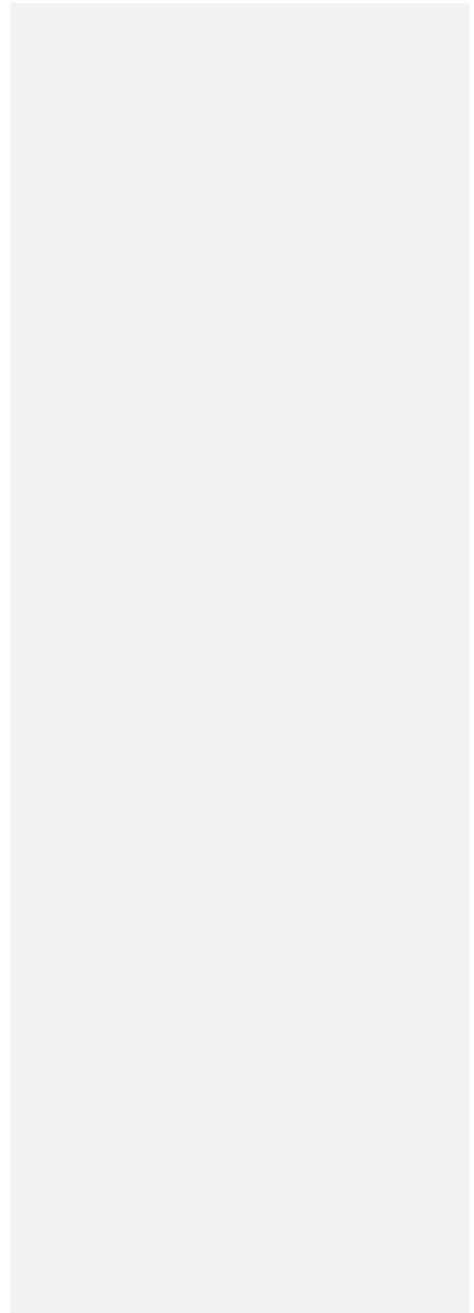
Table 1. Summary of the presented dataset. Average, minimum and maximum values of temperature, dissolved organic carbon (DOC), pH (measured on the NBS scale), total alkalinity (TA) and measured partial pressure of CO<sub>2</sub> (pCO<sub>2</sub>) in the different freshwater ecosystems.

Country	Watersheds	Temperature (°C)			DOC ( $\mu\text{mol L}^{-1}$ )			pH (NBS scale)			TA ( $\mu\text{mol L}^{-1}$ )			Measured pCO <sub>2</sub> (ppmv)			N
		Av.	Min.	Max.	Av.	Min.	Max.	Av.	Min.	Max.	Av.	Min.	Max.	Av.	Min.	Max.	
Brazil	Amazon	30.3	27.4	34.3	352	118	633	6.60	4.53	7.60	385	30	1092	4204	36	18400	155
Kenya	Athi-Galana-Sabaki	25.9	19.8	36.0	307	29	1,133	7.69	6.49	8.57	2290	407	5042	2811	608	10405	44
DRC	Congo	26.3	22.6	28.2	1002	149	3968	6.01	3.94	7.22	212	0	576	6093	1582	15571	97
DRC / Rwanda	Lake Kivu	24.0	23.0	24.7	162	142	201	9.05	8.99	9.17	13037	12802	13338	660	537	772	53
France	Leyre	12.5	7.9	19.2	588	142	3625	6.20	4.40	7.41	280	38	1082	4429	901	23047	92
France	Loire	15.5	8.8	19.3	195	167	233	8.70	8.07	9.14	1768	1579	1886	284	65	717	18
Belgium	Meuse	18.1	13.3	25.9	229	102	404	7.89	6.95	8.59	2769	360	7141	2292	176	10033	50
Madagascar	Rianila and Betsiboka	25.4	20.2	29.5	138	33	361	6.84	5.83	7.62	233	76	961	1701	508	3847	36
Kenya	Shimba Hills	25.1	21.9	31.8	214	36	548	7.37	6.22	8.93	1989	227	14244	2751	546	9497	9
French Guiana	Sinnamary	27.1	24.1	28.7	419	213	596	5.50	5.08	6.30	143	66	290	7770	1358	15622	49
Kenya	Tana	26.6	25.0	27.9	321	193	651	7.65	7.32	8.02	1619	1338	2009	2700	845	6014	51
Zambia / Mozambique	Zambezi	26.9	18.8	31.8	252	103	492	7.59	5.06	9.08	1245	52	3134	2695	151	14004	107
Entire dataset		24.6	7.9	36.0	408	29	3968	7.00	3.94	9.17	1731	0	14244	3707	36	23047	761

Table 2. Median and average values of [dissolved organic carbon \(DOC\)](#), [pH \(measured on the NBS scale\)](#), [total alkalinity \(TA\)](#), and calculated *minus* measured [partial pressure of CO<sub>2</sub> \(pCO<sub>2</sub>\)](#) in the dataset.

	N	% of samples	cal – meas pCO <sub>2</sub> (ppmv)		cal – meas pCO <sub>2</sub> (% of meas pCO <sub>2</sub> )		pH		TA (μmol L <sup>-1</sup> )		DOC (μmol L <sup>-1</sup> )	
			Med.	Av.	Med.	Av.	Med.	Av.	Med.	Av.	Med.	Av.
All samples	761	100%	+611	+10692	+23%	+194%	6.94	7.00	467	1731	315	408
Ranked by calculated-measured pCO <sub>2</sub> as % of measured pCO <sub>2</sub>												
< -10%	122	16%	-540	-890	-34%	-36%	7.89	7.85	1269	1766	259	275
± 10%	174	23%	+15	+50	+2%	+1%	7.67	7.78	1576	3735	228	273
> +10%	465	61%	+2430	+17710	+72%	+327%	6.52	6.49	308	972	360	497
> +50%	280	37%	+ 5490	+ 28660	+162%	+526%	6.18	6.14	192	460	375	567
> +100%	199	26%	+ 9080	+39120	+270%	+710%	5.89	5.96	166	364	389	602
Ranked by pH												
pH > 7	368	48%	+1	+82	+1%	+15%	7.82	7.92	1572	3284	231	255
pH < 7	393	52%	+3280	+20630	+71%	+362%	6.30	6.13	232	277	413	558
pH 6-7	256	34%	+1580	+2710	+40%	+96%	6.58	6.55	334	370	350	427
pH < 6	136	18%	+18410	+ 54486	+308%	+864%	5.50	5.35	93	101	487	828
pH < 5	25	3%	+115580	+209910	+1645%	+3180%	4.53	4.53	41	45	1427	1,843
Ranked by TA												
TA > 2000 μmol L <sup>-1</sup>	110	14%	+20	+340	+2%	+12%	8.58	8.47	7023	8326	163	202
TA 1000-2000 μmol L <sup>-1</sup>	157	21%	-8	-163	-2%	-9%	7.81	7.83	1566	1534	271	295
TA 500-1000 μmol L <sup>-1</sup>	99	13%	+1307	+1900	+28%	+72%	6.97	7.11	651	697	304	318
TA < 500 μmol L <sup>-1</sup>	395	52%	+2070	+20090	+64%	+350%	6.30	6.24	222	232	400	538
TA < 100 μmol L <sup>-1</sup>	82	11%	+6840	+60560	+230%	+1040%	5.50	5.35	59	56	603	988
Ranked by DOC												
DOC < 200 μmol L <sup>-1</sup>	<u>179</u>	<u>24%</u>	<u>+40</u>	<u>+776</u>	<u>+5%</u>	<u>+62%</u>	<u>7.89</u>	<u>7.92</u>	<u>1579</u>	<u>4807</u>	<u>163</u>	<u>149</u>
DOC 200-300 μmol L <sup>-1</sup>	<u>167</u>	<u>22%</u>	<u>+102</u>	<u>+2755</u>	<u>+5%</u>	<u>+69%</u>	<u>7.56</u>	<u>7.37</u>	<u>1132</u>	<u>1259</u>	<u>258</u>	<u>252</u>
DOC 300-400 μmol L <sup>-1</sup>	<u>165</u>	<u>22%</u>	<u>+887</u>	<u>+4473</u>	<u>+25%</u>	<u>+101%</u>	<u>6.90</u>	<u>6.93</u>	<u>499</u>	<u>866</u>	<u>341</u>	<u>344</u>
DOC > 400 μmol L <sup>-1</sup>	<u>250</u>	<u>33%</u>	<u>+3070</u>	<u>+27197</u>	<u>+59%</u>	<u>+434%</u>	<u>6.15</u>	<u>6.14</u>	<u>200</u>	<u>415</u>	<u>555</u>	<u>765</u>
DOC > 800 μmol L <sup>-1</sup>	<u>79</u>	<u>10%</u>	<u>+4995</u>	<u>+62784</u>	<u>+92%</u>	<u>+886%</u>	<u>5.80</u>	<u>5.62</u>	<u>94</u>	<u>180</u>	<u>1099</u>	<u>1438</u>

|



## Figure captions

Figure 1. – [Location of the sampling sites in Africa, Amazonia, and Europe.](#)

Figure 2. – Comparison of results of different water-air equilibration designs for direct pCO<sub>2</sub> measurements; pCO<sub>2</sub> measured with a marbles equilibrator (Congo) and with a membrane equilibrator (Congo and Zambezi) are plotted against pCO<sub>2</sub> measured with a syringe headspace technique. Detection was made with an IR gas analyser.

Figure 3. – Plot of carbon variables *versus* pH in the studied freshwater systems. Top panels are shown with a linear scale and bottom panels with a logarithmic scale; (a,b): measured pCO<sub>2</sub>; (c,d) Total alkalinity; (e,f) Dissolved organic carbon. Zero TA values are plotted as 0.001 in order to be visible on the log pCO<sub>2</sub> scale. Rianila and Bestiboka are plotted together although they belong to different watersheds in Madagascar.

Figure 4. – Comparison between measured and calculated pCO<sub>2</sub> for the whole dataset; (a) calculated *versus* measured pCO<sub>2</sub>, the line shows when measured pCO<sub>2</sub> equals calculated pCO<sub>2</sub>; (b) the difference between calculated and measured pCO<sub>2</sub> as a function of pH; same symbols as in Figure 3.

Figure 5. – Average percentages of pCO<sub>2</sub> overestimation (calculated as 100 x (calculated pCO<sub>2</sub>-measured pCO<sub>2</sub>)/measured pCO<sub>2</sub>), as a function of (a) pH, (b) TA and (c) DOC, for the 12 studied sites. Error bars indicate the standard deviation from the mean for each freshwater systems.

Figure 6. – Sensitivity of pCO<sub>2</sub> overestimation to pH; (a) theoretical factor dpCO<sub>2</sub> / dTA, which describes the sensitivity of calculated pCO<sub>2</sub> to the TA value; (b) the solid lines show the increase in calculated pCO<sub>2</sub> induced by various increases in TA, as functions of pH; these lines mimic the overestimation of calculated pCO<sub>2</sub> generated by increasing contributions of organic alkalinity to the TA; field data (as calculated – measured pCO<sub>2</sub>) have been plotted for

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comparison; same symbols as in Figure 3. Note that negative values do not appear in the logarithmic scale.

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Figure 1. –

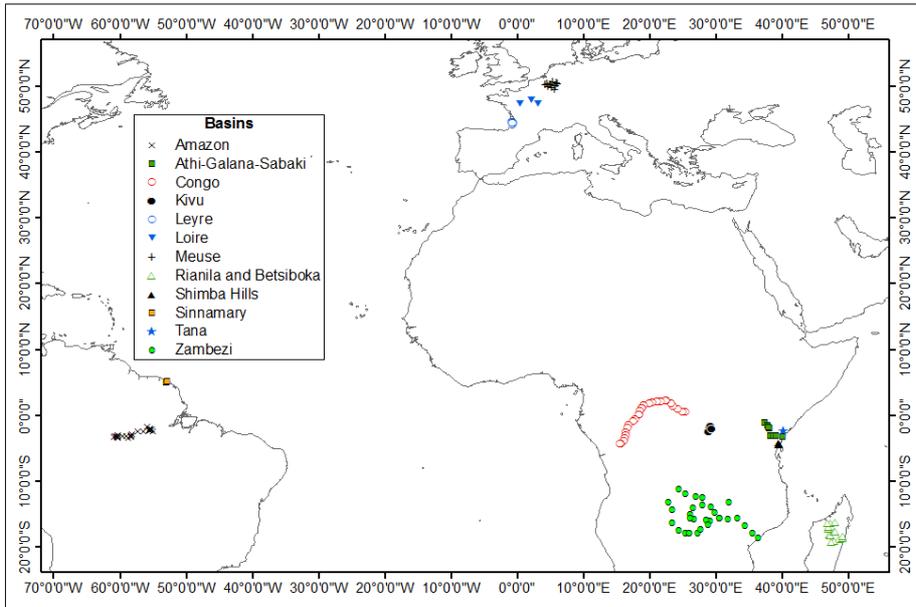
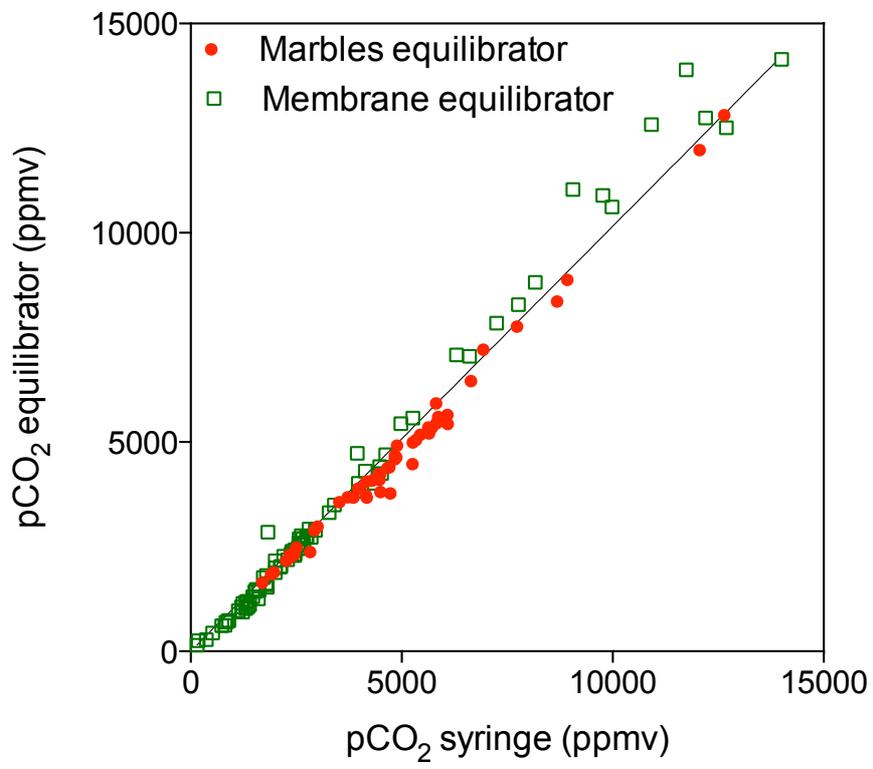


Figure 2. -



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Figure 3. -

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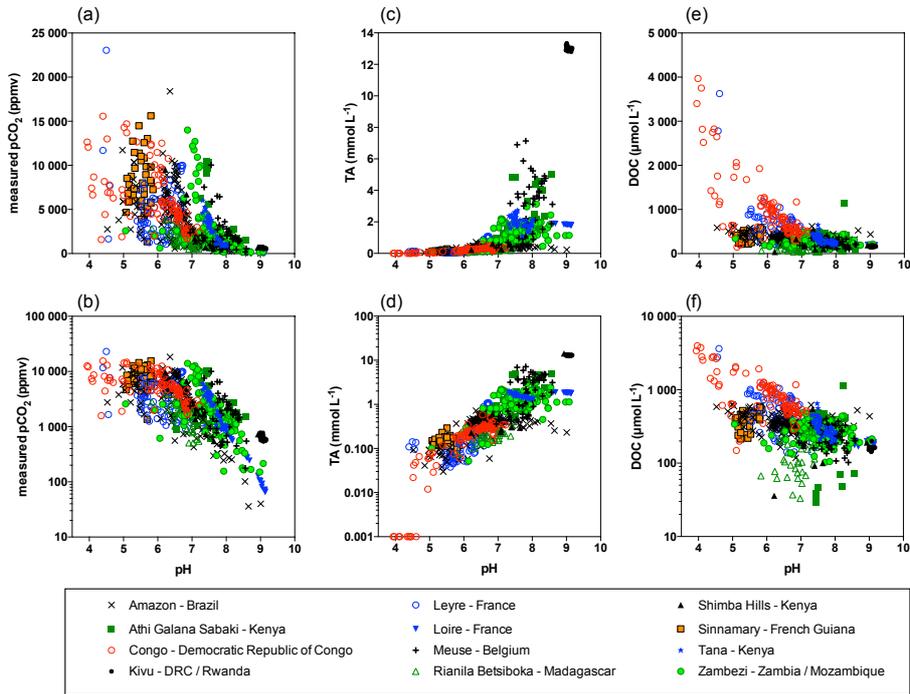


Figure 4. -

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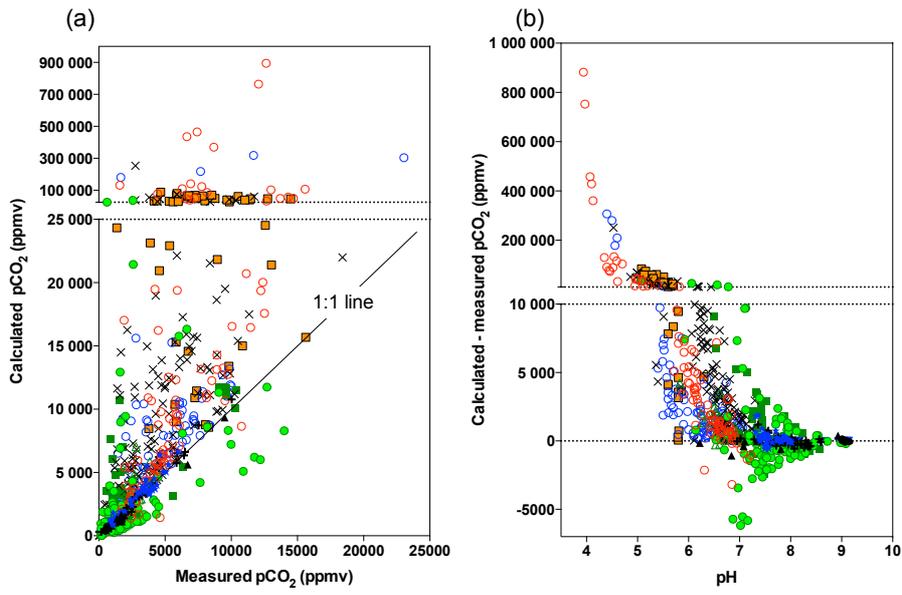


Figure 5. -

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