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balance of NE
Siberian tundra

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The seasonal cycle of the greenhouse gas balance of a continental tundra site in the Indigirka lowlands, NE Siberia

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Abstract

Carbon dioxide and methane fluxes were measured at a tundra site near Chokur-dakh, in the lowlands of the Indigirka river in north-east Siberia. This site is one of the few stations on Russian tundra and it is different from most other tundra flux sta-tions in its continentality. A suite of methods was applied to determine the fluxes of *NEE*, *GPP*, R_{eco} and methane, including eddy covariance, chambers and leaf cuvettes. Net carbon dioxide fluxes were unusually high, compared with other tundra sites, with *NEE*$=-92$ g C m^{-2} yr^{-1}, which is composed of an $R_{eco}=+141$ g C m^{-2} yr^{-1} and *GPP*$=-232$ g C m^{-2} yr^{-1}. This large carbon dioxide sink may be explained by the continental climate, that is reflected in low winter soil temperatures ($-14°$C), reducing the respira-tion rates, and short, relatively warm summers, stimulating high photosynthesis rates. Interannual variability in *GPP* was dominated by the frequency of light limitation (R_g <200 W m^{-2}), whereas R_{eco} depends most directly on soil temperature and time in the growing season, which serves as a proxy of the combined effects of active layer depth, leaf area index, soil moisture and substrate availability. The methane flux, in units of global warming potential, was $+28$ g C-CO$_2$e m^{-2} yr^{-1}, so that the greenhouse gas balance was -64 g C-CO$_2$e m^{-2} yr^{-1}. Methane fluxes depended only slightly on soil temperature and were highly sensitive to hydrological conditions and vegetation composition.

1 Introduction

Tundra covers 8.7×10^6 km^2 globally or 6% of the global land area (USGS land use/land cover classification, Anderson, 1976; van der Molen et al., 2007, [1]). Despite this mod-est area the role of tundra in the global carbon cycle is important: Tundra soils contain

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[1]van der Molen, M. K., Dolman, A. J., Belelli, L., et al.: The carbon balance of Boreal Eurasia consolidated with eddy covariance measurements, in review, 2007.

an estimated 200 Pg C or about 30% of the global soil carbon pool (Post et al., 1982, Schlesinger, 1999, Hobbie et al., 2000); Global warming is expected (IPCC, 2007) and observed (Serreze et al., 2000, Chapin et al., 2005) to be amplified in polar regions and Northern wetlands (>60° N) emit ~8% of the global natural methane flux (Cao et al., 1998). Being a strong greenhouse gas, increasing methane concentrations are responsible for ~30% of the enhanced greenhouse effect (Forster et al., 2007). With an estimated carbon dioxide sink of 252 Tg C yr^{-1} for Russian tundra areas, their role in the global carbon cycle is also substantial (van der Molen et al., 2007[1]). The stability of the tundra carbon pools is thus relevant to the global climate. To constrain the fluxes of carbon dioxide and methane, several experiments have been conducted in the past decades (cf. Chapin et al., 2000; McGuire et al., 2003). Major conclusions of these experiments are that the fluxes, particularly those of methane, show a high spatial variability due to heterogeneity in topography, vegetation and hydrology, even at the smallest scales of polygons and floodplains (van Huissteden et al., 2005; Kwon et al., 2006). Decomposition of soil carbon is sensitive to temperature changes and enhanced thawing (Zimov et al., 2006; Wagner et al., 2007), but growth of tundra ecosystems is nitrogen limited, which would reduce the sensitivity to climate warming (Hobbie et al., 1998). The effect of climate warming in recent decades is already detected in the form of enhanced shrub growth and an advancement of the tree line (Jia et al., 2003, Lloyd et al., 2003; Esper et al., 2004).

The majority of the tundra field experiments were conducted on the North slope of Alaska and the Seward Peninsula in Alaska. The tundra areas in the Russian Federation, which comprise 45% of the total (van der Molen et al., 2007[1]), are much less studied: Heikinnen et al. (2002) report about an experiment in Vorkuta, European Russia, Wagner et al. (2003) about an experiment in the Lena Delta and Tsuzuyaki et al. (2001) and Corradi et al. (2005) about experiments in the floodplains of the Kolyma River near Cherskii. Christensen et al. (1995) performed methane flux measurements along a transect through Arctic Siberia via the Arctic Ocean. Other greenhouse gas flux studies were performed on tundra in Sweden (Christensen et al., 2004) and Greenland

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(Christensen et al., 2000; Friborg et al., 2000; Soegaard et al., 2000). However, climatic and environmental conditions vary considerably across the arctic, as a function of continentality (Sect. 2.1), distance to the northern treeline and the presence of mountain ranges (cf. Lynch et al., 2001).

Here we present the first combined observations of carbon dioxide and methane fluxes collected at a tundra site in the lowlands of the Indigirka river, near the village of Chokurdakh, in North East Siberia. At about 150 km from East Siberian Sea, the site is located roughly half way between the coast and the tree line. The nearest mountains are about 450 km to the South, making the latitudinal extend of the Indigirka lowlands some 600 km, in comparison with 350 km between the Brooks Range and the Arctic Ocean in Alaska. As such, climate is more continental and less influenced by topography at this site than most other Arctic measuring sites. As 45% of the tundra area lies in the Russian Federation, where it is the largest land cover class after forest, this study provides data necessary to understand the carbon dioxide and methane balances of Boreal Eurasia and to put existing tundra field experiment in a circumpolar perspective.

Carbon dioxide fluxes were measured from 2003 to 2006, continuously during the summer using eddy covariance, complemented with chamber and leaf cuvette measurements of ecosystem respiration and photosynthesis rates. Methane flux measurements were carried out during intensive field experiments in the summers of 2004, 2005 and 2006. The objective of this study is to present the seasonal cycle of the carbon dioxide and methane fluxes, the resulting global warming potential, as well as the annual sums. For a better interpretation of the measurements, we perform and validate a partitioning of the net carbon dioxide fluxes into the components of ecosystem respiration and photosynthesis. We also use and validate the ORCHIDEE photosynthesis model and the PEATLAND-VU methane flux model to scale up the measurements in time and space.

2 Site description and instrumentation

2.1 Site description

The research site (70°49′ 36.28″ N, 147°29′ 56.23″ E, 10 m a.m.s.l., Fig. 1) is located in the WNF Kytalyk (Crane) reserve, 28 km northwest of the village of Chokurdakh in the Republic of Sakha (Yakutia), Russian Federation. This puts the site roughly half way between the East Siberian Sea (150 km to the North) and the transition zone between taiga and tundra. The site is on the bottom of a thermokarst lake, that was drained by intersection of the Berelekh river, a tributary of the Indigirka river (Fig. 1).

2.1.1 Climatology

A basic climatology is presented in Table 1. With a mean January temperatures of −34.2°C, the Chokurdakh site is colder than at other field sites near Scoresbysund (North East Greenland), Kiruna (North Sweden) and Vorkuta (East European Russia) (−13 to −16°C), it is also colder than Barrow (Alaska, −25.4°C), and Tiksi, at the Lena Delta (−30°C). Temperatures below −40°C occur regularly. The mean July temperature of +10.4°C is warmer than in Barrow (4.6°C) and Scoresbysund (6.0°C), comparable to Tiksi and somewhat cooler than Kiruna and Vorkuta (12–14°C). Maximum temperatures at the site may reach over +25°C. The mean annual temperature is −10.5°C. Monthly mean temperatures are quite variable in the winter, and more constant in the summer. Annual mean precipitation amounts to 212 mm, of which about half falls as snow. Snow depths at the site are 60–80 cm and quite constant throughout the years, this is somewhat more than measured at the long-term weather station near the village. With wind speeds around 4 m s^{-1} the site is amongst the calmer sites, Barrow and Vorkuta in particular experience stronger winds (∼5 m s^{-1}). Wind directions are distributed fairly even, with a slight preference for North-easterly and South-westerly directions. On average, the summer months June, July, August receive about 260, 250 and 150 h of sun. The years 2004 and 2005 received considerably less and more than

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average (90 and 280 h, respectively).

During the operation years of the site, water level in the nearby river varied considerably, and consequently also on the measurement sites. In 2004, the river stage was relatively high after high snowmelt runoff. In 2005, the river stage was approximately 1.5–2 m lower, as a result of a dry winter and spring; moreover, the air temperatures were as high as 30°C during most of the field campaign. The river stage in 2006 was intermediate.

2.1.2   Geology and soils

Three major topographic levels occur around the measurement site (Fig. 2). The highest level is underlain by "Ice complex deposits" or "Yedoma", ice-rich silt deposits of Late Pleistocene, deposited as loess or fluvial silts (Schirrmeister et al., 2002; Gavrilov et al., 2003; Zimov et al., 2006). Near the site, the ice complex deposits occur in terrace-like 20–30 m high hills, probably representing a Pleistocene river terrace surface which has been eroded by thermokarst processes. Presence of cross-bedding in a riverbank exposure near the site indicates a fluvial origin of the sediments. The measurement site itself is located in a depression between two N-S trending ice complex remnants, constituting the second topographic level. This depression originated as a thermokarst lake of Holocene age, drained by fluvial erosion. The lowest topographic level is the present river plain, situated 2–3 m below the lake bottom. The river plain has a conspicuous fluvial relief with levees, back swamps and lakes. Active thermokarst features (slumps and thermo-erosional niches) are common along the river bank, and thermo-erosional expansion of creeks and sloughs is also common on the river plain.

The area is underlain by continuous permafrost. The active layer ranges from 25 cm in dry, peat-covered locations to 40 cm in wet locations. On the floodplain the active layer may be locally thicker. Both on the floodplain and the lake bottom a network of ice wedge polygons occurs, in general of the low-centered type. The ice wedge polygons on the lake bottom have a more mature appearance, with well developed ridges and

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wet, low centers. This suggests that the age of this surface is older than that of the river plain. Next to the polygons, also low, flat palsa-like features occur on the lake bottom, representing generally drier areas. The lake bottom is drained by a diffuse network of depressions, covered with a *Carex-Eriophorum* vegetation and underlain

5   by a generally thicker active layer.

The soils generally have a 10–15 cm organic top layer overlying silt. In case of wet sites, the organic layer consists of loose peaty material, composed either of sedge roots or *Sphagnum* peat, depending on the vegetation. Drier sites tend to have a thinner, more compact organic layer.

10   The Berelekh river meanders from west to east. In a band of about 30 km wide, centered at the river, many small lakes occur, with diameters of a few hundred meters. In a wider area (70°40′ N to 71°20′ N and 146°00′ E to 148°00′ E or about 75×75 km) fewer lakes occur, which are larger, with typical diameters of 10 km. Based on a detailed digital map, the fraction of water in the narrow band was estimated at 17% and at 22%

15   in the wider area. The relatively large fraction of water needs to be taken into account when interpreting the flux measurements made on land.

2.1.3 Vegetation composition

In the area in the immediate surroundings of the tower, the vegetation is a mosaic of: 1) drier sites (palsa, ridges along ice wedges) dominated by *Betula* nana, *Salix* sp.,

20   mosses and grasses; 2) isolated depressions (polygon centers, thawing ice wedges) dominated by submerged *Sphagnum*, *Potentilla palustris* and some *Carex*, 3) mires with *Sphagnum* hummocks and some *Salix* and 4) depressions with a dense *Carex-Eriophorum* vegetation.

For the methane flux measurements, a classification has been developed, linking

25   vegetation and geomorphology (Van Huissteden et al., 2005; Table 2, Fig. 2). At the highest level, this classification distinguishes between floodplain and river terrace/tundra. Next, dry and wet sites have been distinguished based on water table position, where "wet" is defined as largely water saturated soils, with a water table not

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lower than 5 cm below the surface. The lowest level of the classification is based on smaller morphological features and their vegetation. Van Huissteden et al. (2005) give a detailed description of all classes. The areal fraction of the terrain units has been determined by point counting at regular distances along transects, near the eddy correlation flux tower and on the floodplain area near the field station. In general, in the area to the south and west of the tower, the wet TW classes dominate while to the north and east of the tower the dry TD classes dominate (Fig. 3). On the floodplain, FW2 dominates.

2.2 Instrumentation

2.2.1 Eddy covariance and micrometeorology

At the site two masts are installed, the first mast contains the eddy covariance instrumentation, consisting of an ultrasonic anemometer (Gill Instruments, Lymington, UK, type R3-50) and an open path infra-red gas analyzer (Licor, Lincoln, NE, USA, type Li-7500) at an height of 4.7 m. Eddy covariance data were collected on a handheld computer (van der Molen et al., 2006) at a rates between 5 and 10 Hz, depending on the time between field visits and storage capacity. Eddy fluxes were computed on a half-hourly basis following the Euroflux methodology (Aubinet et al., 2000) with the addition of the angle of attack dependent calibration (van der Molen et al., 2004; Nakai et al., 2006). Storage flux was corrected for using the discrete approach, which did not significantly change the fluxes however. The second mast contains a shortwave radiometer (Kipp & Zn, Delft, the Netherlands, type albedometer, CM7b), up- and down facing longwave radiometers (The Eppley Laboratory, Newport, RI, USA, type PIR) and a net radiometer (Campbell Scientific, Logan, UT, USA, type Q7). 20 soil thermometers (made at the Vrije Universiteit Amsterdam) were dug into 2 profiles, each reaching 60 cm into the ground. The one profile was in a polygon depression, where the soil is more moist, the other profile was on the rim of a polygon with a relatively low water table. Soil moisture was not measured. Small scale variations in topography in relation

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to polygon mires causes a rather heterogeneous soil moisture field, however, during the entire growing season, the soil moisture conditions are wet. The instruments were usually installed in April each year and taken down for the winter in October. However, because the system was operated on solar power and batteries, and the area is inac-
cessible during the period of snow melt and ice breaking (May), power failures caused the system to shut down in the spring of 2004, 2005 and 2006, but each year's record starts at least within a few days after leave onset. Low solar radiation conditions cause some gaps in the record during the fall. An additional wind generator was installed in 2006 to help prevent power failures.

2.2.2   Chamber measurements of the soil respiration rate of CO_2

Observations of the soil respiration rate were made using a portable gas analyzer (PP Systems, Hitchin, UK, type EGM-4) equipped with a closed chamber (type SRC-1) and a soil temperature probe (type STP-1). The chamber has dimensions of (height × diameter) = (15×10 cm). 25 aluminium rings on which the chamber fits precisely were installed in the field at various locations with representative vegetation cover, so that the respiration rates were measured each time at the same places without disturbing the soil. The increase in volume was corrected for. Observations were carried out during the summer field visits and preferably in a 24 h period at 3 hourly intervals. Some variation of soil respiration rate was observed between the various locations, but the average fluxes per location varied less than a factor 2. The differences were not consistent and variations due to temperature and weather were of similar magnitude. Because respiration rates at each of the 25 locations were normally distributed, as opposed to the methane fluxes (van Huissteden et al., 2005), we considered the mean soil respiration rate over each of the 25 locations as representative for the soil respiration in the footprint of the flux tower.

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2.2.3 Photosynthesis measurements

In 2003, measurements of photosynthesis rates were made using a portable LCA-4 (ADC Bioscientific, Herts, UK) Infrared Gas Analyser with leaf cuvettes. In 2004, a LI-6400 system (Li-Cor, Lincoln, NE, USA) was used. The photosynthetic activity was measured of leaves of *Betula*, *Salix*, *Eriophorum* plants. Of each species 2–3 leaves were sampled and the measurements were repeated 3–5 times. The entire measurement cycle was repeated at three hour intervals during 24 h periods, using the same leaves. After the last measurement cycle, the leaves were taken to determine the leaf area. No systematic difference in photosynthetic activity was observed between species, and the variation between leaves of the same species was of the same order of magnitude as the variation between species. For this reason the measurements taken in a three hour interval were averaged into a single value.

2.2.4 Methane flux measurements

Methane flux measurements were made during a number of consecutive days in the summers of 2004 (27–30 July), 2005 (20–27 July), 2006 (15–18 August). The methane flux measurements were made using static chambers, at 55 sites in 2004, 86 sites in 2005 and 60 sites in 2006, selected from the terrain classes in Table 2 for determination of spatial variation of the fluxes. The round static chambers were attached to a photo-acoustic gas monitor (model 1312, Innova AirTech Instruments, Ballerup, Denmark), capable of measuring CO_2, H_2O, N_2O and CH_4 concentrations. In 2006 a model 1412 Innova was used, equipped for H_2O and CH_4 measurements. The detection limit for CH_4 is 0.1 ppmv, resulting in a theoretical minimum detectable flux of 0.13 mg CH_4 m^{-2} hr^{-1} given the measurement setup. For further details on the measurement methodology we refer to van Huissteden et al. (2005). In total, 201 methane flux measurements have been made during these three field campaigns. Each flux measurement was accompanied by determination of the active layer depth and soil temperatures. At each site characteristics of the vegetation and soil profile, and the water table were recorded,

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using a hand auger. Each flux measurement was quality controlled following van Huissteden et al. (2005).

3 Validation of measurements and models

3.1 Energy balance closure

Figure 4 shows the energy balance closure as a method to test the quality of the eddy flux data. The linear least square regression through all data points shows a good energy balance closure of 99.8% with an offset of $14\,W\,m^{-2}$ and a standard deviation of $60\,W\,m^{-2}$. However, the binned data suggest an underestimation of the larger fluxes. The soil heat flux G was estimated here as a function of the change of temperature in the profile and typical soil characteristics as shown in Table 3.

3.2 CO_2 fluxes during calm conditions

Underestimation of ecosystem CO_2 fluxes under calm conditions is amongst the most prominent error sources of the eddy covariance method. Figure 5 shows some indications that this so-called u_*-problem may occur at the field site for $u_* < 0.2\,m\,s^{-1}$, although the underestimation is small compared to what is sometimes observed at other sites (cf. Dolman et al., 2004). Because calm, "night-time" conditions may also occur during the polar day at this high latitude site, we estimate NEE for $u_* < 0.2\,m\,s^{-1}$ as follows:

$$NEE = R_{eco,mod} - GPP_{mod} \quad (u_* < 0.2\,m\,s^{-1}) \tag{1}$$

where $R_{eco,mod}$ and GPP_{mod} are model estimates of the ecosystem respiration and photosynthesis rates (see Sects. 3.3 and 3.5).

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3.3 Parameterising the ecosystem respiration rate

The ecosystem respiration rate was estimated as the CO_2 flux measured with eddy covariance under conditions of low global radiation ($<20\,W\,m^{-2}$) and strong turbulence ($u_* > 0.2\,m\,s^{-1}$). Figure 6a shows that the ecosystem respiration rate increases with temperature, but that soil temperature does not explain all the variation. The timing in the growing season appears to explain a large part of the remaining variation, as indicated by the colours of the data points. A linear optimization of the model

$$R_{eco,mod} = R_0 Q_{10}^{(T/10)} \quad (180 < doy < 240) \tag{2}$$

where the base respiration R_0 ($\mu mol\,m^{-2}\,s^{-1}$) is written as a 3rd degree polynomial function of the day-of-year, is shown in Fig. 6b. The parameters of the model are given in Table 4. The rationale for using day-of-year as a proxy for R_0 is that active layer depth, biomass and substrate co-vary during the growing season and it is as yet impossible to distinguish between those. For dates outside the range 180 to 240, Eq. (2) is fixed to 180 or 240. The resulting $R_{eco,mod}$ has a slope of about 1.0 versus $R_{eco,obs}$ but with considerable scatter ($r^2=0.44$). The unexplained part of the variation may be due to heterogeneity of vegetation composition or ground water table. The respiration rate does not vary with wind direction.

3.4 Validation of partitioning of NEE into R_{eco} and GPP

The partitioning of *NEE* as measured by the eddy covariance method into R_{eco} and *GPP* may be validated against chamber and leaf level photosynthesis measurements. Figure 7 shows a comparison of the diurnal cycles of respiration rates resulting from Eq. (2) with chamber measurements of the respiration rate for a few consecutive days in 2004 and 2005. Figure 7 shows agreement in the order of magnitude, but there is also considerable variation. Particularly on days 209 and 210 in 2004 (Fig. 7a), the chamber fluxes are larger than the eddy fluxes, but at these days relatively few reliable chamber

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measurements were made, due to malfunctioning of the battery. From day 211 to 214, when measurements were made at a higher temporal resolution, the agreement is closer. The x's indicate the variability of eddy covariance flux measurements at the corresponding times, for turbulent ($u_* > 0.2\,\mathrm{m\,s^{-1}}$) and dark ($R_g < 20\,\mathrm{W\,m^{-2}}$) conditions.

5   The photosynthesis rate *GPP* was estimated as $R_{eco} - NEE$. Diurnal cycles of photosynthesis rates were also measured independently (see Sect. 2). Figure 8 shows a good level of agreement between *GPP* from the partitioning of eddy fluxes on the one hand and leaf level measurements on the other hand on 14 July 2003 and 30 July 2004. Figure 8 also shows model simulations of *GPP*, using the ORCHIDEE model that

10 is described in Sect. 3.5. The close agreement in absolute values and in the shape of the diurnal cycles of R_{eco} and *GPP* gives confidence in the performance of the method of partitioning *NEE* into *GPP* and R_{eco}.

3.5   Modelling of photosynthesis and methane fluxes

The ORCHIDEE photosynthesis model (Krinner et al., 2005; Morales et al., 2005) was
15 used in combination with the PEATLAND-VU methane flux model (van Huissteden and van den Bos, 2003; van Huissteden, 2004). ORCHIDEE simulates *GPP* as a function of solar radiation, surface temperature, air humidity, air pressure, CO_2 concentration and surface conductance (Farquhar, 1980; Ball et al., 1987, Collatz et al., 1992). ORCHIDEE may also be used as a dynamical vegetation model, but in this application
20 only the photosynthesis module was used, without accounting for phenology. The C3 grassland plant functional type was used to simulate the tundra photosynthesis rate, after adapting the Vc,max (Vj,max) from 60 (120) to 35 (70). The Leaf Area Index was maintained at 1.0. The performance of the ORCHIDEE model for tundra is shown in Fig. 9, using mid summer data, when the vegetation was fully developed. Further
25 validation is provided in Fig. 8, where diurnal cycles of simulated *GPP* are compared with leaf cuvette measurements and with partitioned eddy covariance measurements. The simulated *GPP* are used for gap filling, when meteorological variables were available and eddy fluxes were not available, or when eddy fluxes were rejected due to low

turbulence conditions (Eq. 1).

PEATLAND-VU simulates CH_4 fluxes as the difference of production in the root zone and consumption by oxidation (cf. Walter, 2000). The production is a function of the rate of root exudation of labile organic compounds, which is generally assumed to depend on NPP, as well as a function of the availability of oxygen in the soil. For this purpose, we assume that NPP corresponds to 50% of *GPP* (Turner et al., 2006), or 125% of *NEE* (which are numerically identical in this instance). The NPP input for PEATLAND-VU is derived from the eddy covariance data, augmented with modelled values from ORCHIDEE whenever gaps in the data were present. Oxidation is a function of soil aeration, which varies with ground water table. The ground water table is modelled as a function of the balance between precipitation and evaporation, which are both observed, and snowmelt, using a modified version (Yurova et al., 2007) of the model of Granberg et al. (1999). The modelling of the water table for the site and the methane flux simulations using the modelled water table are described in a separate publication by Petrescu et al. (2007)[2]. With PEATLAND-VU, methane fluxes were simulated for three different vegetation types representative for the vegetation around the tower site: *Carex* vegetations (TW1 type), *Sphagnum* vegetations (TW3-4 type) and dry tundra vegetation (TD types). The NPP generated from the eddy covariance data or ORCHIDEE was used as input for PEATLAND-VU. It was assumed that the ombrotrophic *Sphagnum* vegetations and dry tundra vegetations had a lower primary production (20% lower than average NPP) than the TW1 type vegetation (20% higher than average NPP). On wet sites, oxidation is a function of vegetation species, with *Carex* species having aerenchyma, which enable efficient transport of methane to the atmosphere, with less oxidation. *Sphagnum* mosses have a symbiosis with methanotrophic bacteria (Raghoebarsing et al., 2005), resulting in a larger methane oxidation. The vegetation-related parameters (plant transport rate factor and oxidation factor) in

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[2]Petrescu, A. M. R., Jackowicz-Korczynski, M., Yurova, A., van Huissteden, J. C., Christensen, T. R., Crill, P. M., and Maximov, T. C.: Modelling methane emissions from arctic wetlands: effects of hydrological parameterization, in preparation, 2007.

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the CH_4 submodel of PEATLAND-VU have been set accordingly. For the dry sites, water table was set 20 cm lower than for the wet sites, in line with the field observations made during the methane flux measurements. After initial setting of these parameters, the model output was further optimized on both the methane production rate, a tuning parameter in the model (Walter, 2000), and the plant oxidation rate. The performance PEATLAND-VU model for methane fluxes is shown in Fig. 10, where simulated methane fluxes are plotted versus observed ones, for the three different simulated vegetation types and the three available measurement campaigns. Although the uncertainty ranges are quite wide, the actual values compare rather well. This model was used successfully also for methane fluxes in Stordalen, Abisko, Sweden (Petrescu et al., 2007[2]). The methane model was used to scale up the observed fluxes in time to the complete length of the growing season for the sites around the flux tower. The wet floodplain (FW) have been modelled by Petrescu et al. (2007)[2] but are not considered here because the data from the floodplain do not permit a consistent comparison between CO_2 and CH_4 fluxes.

3.6 Comparison of carbon dioxide and methane fluxes

As the main objective of this paper is to determine the net greenhouse gas budget for this site, it is convenient to express methane fluxes in units of global warming potential. Based on the IPCC 4th Assessment report (Forster et al., 2007), a mass of methane gas has 25 times more global warming potential (GWP) than the same mass of carbon dioxide gas, which has a GWP of 1 by definition. This factor of $25 \, g \, CO_2 \, (g \, CH_4)^{-1}$ results from integrating the radiative effects of a pulse emission (or removal) over a time horizon of 100 years. The GWP of methane decreases with the integration time, because atmospheric methane oxidises. In terms of GWP, a flux of $1 \, g \, CH_4 \, m^{-2} \, day^{-1}$ is equivalent to $25 \, g \, CO_2 \, m^{-2} \, day^{-1}$, or $25 \, g \, CO_2e \, m^{-2} \, day^{-1}$. However, it is common practice to express carbon dioxide fluxes in terms of the mass flux of the carbon atom only. Thus a flux of to $25 \, g \, CO_2 \, m^{-2} \, day^{-1}$ is written here as $25 \times (12/44) = 6.8 \, g \, C \, m^{-2}$

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day^{-1}. Table 5 gives an overview and comparison of these units and their numerical value. We use the units of carbon flux in g C-CO$_2$e m^{-2} day^{-1}. For carbon dioxide fluxes this is identical to g C m^{-2} day^{-1} numerically and in terms of global warming potential.

5   This methodology of assigning methane (and other ghg's) a global warming potential to compare its radiative effects of different greenhouse gasses has been commonly applied since it was adopted in the Kyoto protocol. However, it has a few shortcomings: first, the method strictly only applies for pulse emissions/removals, whereas natural landscapes are better characterized as continuous sources and sinks; second, it

10   only expresses the radiative effects over a fixed time horizon (100 years), whereas in practice the radiative effects evolve dynamically. Over short periods, methane emissions have strong radiative effects, but due to the chemical removal of methane from the atmosphere, the impact decreases over time. Carbon dioxide on the other hand, although being a less effective greenhouse, has a much longer residence time in the at-

15   mosphere. Consequently, the radiative effects accumulate and may eventually exceed those of methane. We adopt the methodology of Frolking et al. (2006) to compare the short-term and long-term effects of the carbon dioxide and methane fluxes from this tundra site.

4   Results

20   ### 4.1   Methane flux measurements

The methane fluxes show a large spatial and temporal variation. The fluxes measured on the river floodplain (FW classes) are considerably higher than those of the *Sphagnum*-rich sites on the river terrace (TW classes). Only sites in the TW1 class show fluxes that are comparable to those of FW sites (Van Huissteden et al., 2005), although

25   they are still lower (Fig. 11). Dry sites (TD and FD classes) generally show negative fluxes (uptake) and sometimes slightly positive fluxes. Positive fluxes decrease rapidly

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with lower water table (Van Huissteden et al., 2005).

Compared to van Huissteden et al. (2005), who reported about the 2004 campaign, the 2005 and 2006 field campaigns add to understanding of the temporal variation of the fluxes. Fluxes were highest in 2004, and lowest in 2005, despite the higher the air and soil temperatures. In 2006, the fluxes were intermediate, while the soil surface temperature was lower than in 2005. As such, high fluxes correspond well with high river water levels (Sect. 2.1.1). Statistical analysis shows that methane emission increases significantly with height of the water table and with active layer thickness. Methane emission decreases with surface temperature, and does not significantly vary with soil temperature at 10 cm depth. This surprising results are discussed in Sect. 5.1. The negative fluxes on dry sites do not show any significant correlation with the environmental variables above, albeit that negative fluxes only occur on sites with low water table. Sites where the water table was only a few cm below the surface may already show negative methane fluxes. At water table depths below –5 cm, fluxes higher than 5 mg CH_4 m^{-2} h^{-1} do not occur and negative fluxes dominate.

Upscaling of the methane fluxes has been performed by multiplying the fluxes with the areal fraction of the different terrain units. This gives an integrated methane flux for the terrace area around the flux tower and the investigated floodplain area (Table 6, Fig. 11). The integrated fluxes are small on the terrace compared to some of the site fluxes, due to the large relative area of dry sites (68%), which are mostly located to the east of the tower. The area to the west of the tower consists of a mosaic of dry and wet sites, associated with polygons. Particularly the polygon ridges contribute to the fraction of drier areas with negative fluxes, a phenomenon that is also known from other Siberian tundra sites (Wagner et al., 2003; Wille et al., 2007). In contrast, only 37% of the floodplain area consist of dry sites, and mosaic-like pattern are not as pronounced there. Consequently, the contribution of the floodplain sites to the integrated flux is more large considering their small relative area.

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4.2 Seasonal course of carbon dioxide and methane fluxes

Daily fluxes of *NEE*, its components *GPP* and R_{eco} and methane fluxes are presented in Fig. 12 for the entire period of record. The increase in *GPP* at the start of the growing season of 2003 is remarkably sharp. Uptake by photosynthesis is quite variable from day to day, whereas ecosystem respiration rates vary much slowlier throughout the year. As a consequence, *NEE* is also quite variable, particularly in a relative sense. The growing season lasts about 60 days, in July and August. The ecosystem respiration increases steadily from the start of the growing season until the second half of August, when it starts to decline. Methane fluxes make up a significant part of the greenhouse gas budget and are largest at the onset of the growing season, when wet conditions prevail due to snow melt. Considering that the methane flux in carbon flux units (Table 5) is only 12% of the flux indicated in Fig. 12, the methane fluxes play only a minor role in the carbon budget.

In order to quantify inter-annual variability, daily fluxes were averaged to weekly fluxes and shown in Fig. 13. It appears that interannual variability is small for ecosystem respiration, and larger for *GPP* and *NEE*, as well as for methane fluxes. The largest variability in *GPP* occurs at the start of the growing season, implying that the date of snow melt and the start and length of the growing season are important factors determining the carbon and greenhouse gas balances. The variability at the end of the growing season is smaller, because weather conditions are less important than the limitation due to the shorter day lengths. The lower panel of Fig. 13 again confirms that the greenhouse gas balance is primarily determined by the carbon dioxide component, with a smaller but significant role for methane fluxes. Photosynthesis rates are $3.5\,\mathrm{g\,C\,m^{-2}\,yr^{-1}}$ in the middle of the growing season, and consistently larger in 2003.

The annually cumulative fluxes are shown in Fig. 14. These fluxes result from averaging the weekly fluxes shown in Fig. 13 over all years and then integrating. Van der Molen et al. (2007)[1] show that this methods of calculating the mean annual flux over multiple years takes optimal advantage of the available data, particularly when non-

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overlapping gaps occur. It is clear that the net carbon flux *NEE* is the relatively small difference ($-92 \, \mathrm{g\,C\,m^{-2}\,yr^{-1}}$) between the large terms of *GPP* ($-232 \, \mathrm{g\,C\,m^{-2}\,yr^{-1}}$) and R_{eco} ($+141 \, \mathrm{g\,C\,m^{-2}\,yr^{-1}}$), which makes *NEE* sensitive to relatively small changes in either *GPP* or R_{eco}. The methane emissions are $28 \, \mathrm{g \, C\text{-}CO_2e \, m^{-2} \, yr^{-1}}$. This is equivalent to a methane emission of $4.1 \, \mathrm{g \, CH_4 \, m^{-2} \, yr^{-1}}$ and a carbon flux of $3.1 \, \mathrm{g\,C\,m^{-2}\,yr^{-1}}$. As a consequence, the greenhouse gas balance is negative, and the site is a net sink of $-64 \, \mathrm{g \, C\text{-}CO_2e \, m^{-2} \, yr^{-1}}$. for greenhouse gases.

The uncertainty of annual totals of *NEE* collected with eddy covariance measurements was estimated at $41 \, \mathrm{g\,C\,m^{-2}\,yr^{-1}}$ by van der Molen et al. (2007)[1]. Based on the CH_4 flux measurements on the terrace, the coefficient of variation of the measured fluxes is 92%; in Fig. 10 the variation of the modelled values for 5-day periods is similar. Therefore we estimate the uncertainty of the methane flux measurements as $25.8 \, \mathrm{g \, C\text{-}CO_2e \, m^{-2} \, yr^{-1}}$. Assuming that uncertainty is normally distributed, the confidence level α that the site is a sink for carbon dioxide is $\alpha=0.94$, a source for methane ($\alpha=0.22$) and a sink for greenhouse gas gases ($\alpha=0.83$).

Using the methodology of Frolking et al. (2007) to determine the temporal evolution of radiative forcing of sustained carbon dioxide and methane fluxes, we find that on short time horizons (<13 years), the methane emission has stronger radiative impacts than the carbon dioxide sink. However, because the change in atmospheric methane concentration may be considered in equilibrium with the methane source, the radiative forcing due to methane emission has settled at $1.1 \times 10^{-14} \, \mathrm{W\,m^{-2}}$ per $\mathrm{m^2}$ of tundra source area, which is about $4.4 \times 10^{12} \, \mathrm{m^2}$ in the Russian Federation and about $8.7 \times 10^{12} \, \mathrm{m^2}$ worldwide (van der Molen et al., 2007[1]). Over time horizons longer than 13 years, the radiative effect of the sustained carbon dioxide removal from the atmosphere becomes dominant. Considering the age of tundra is older than that, this site may be considered a source of greenhouse gasses and acts to cool the climate.

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5 Discussion

5.1 Methane fluxes

A main feature of the CH_4 fluxes is the very high spatial variability which is related to vegetation and water table variability. The water table effect is directly related to anaerobic conditions in the soil and has been documented by many authors (Bartlett et al., 1992; Friborg et al., 2000; Heikinnen et al., 2002; Oberlander et al., 2002; Wagner et al., 2003; Kutzbach et al., 2004). Statistical analysis shows that the spatial heterogeneity of the terrain mainly affects water table variation, and soil temperature to a much smaller extent. The correlation patterns with soil temperature and active layer thickness further confirm the dominating effect of the water table. The negative correlation of soil surface temperature and the poor correlation of soil temperature at 10 cm depth with methane emission rates seems to contradict the often reported positive effect of temperature on methane fluxes of higher soil temperatures, related to microbial reaction rates (cf. Morrissey and Livingston, 1992; Christensen et al., 1995, 2003; Verville et al., 1998; Treat et al., 2007). We hypothesize that this reflects the evaporative cooling effect of the wet soil surface, and the generally thinner active layer on the terrace. Also the adaptation of the microbial population to low temperatures, causing high production rates even at near-zero temperatures, contributes to the low temperature sensitivity of the CH_4 fluxes (Wagner et al., 2003; Rivkina et al., 2000, 2007). Apparently, the sensitivity of CH_4 fluxes to temperature pertains rather to large scale variations between sites at different latitudes (Christensen et al., 2003) than within-site and short-term temporal variation in temperature. The positive correlation of methane flux with active layer depth may be a secondary effect. High water tables increase the methane flux but flooding also tends to increase active layer thickness (French, 1996, and references therein). Also, the active layer thickness co-varies with substrate availability throughout the season (sea above). The main driver of methane emission is water table, which determines soil temperature and active layer thickness as well, particularly on the river terrace.

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Water table also drives the temporal variability of the CH_4 fluxes. On a year-to-year time scale there appears to be no clear influence of temperature. This does not exclude that the temperature influence should operate on a seasonal time scale, but as yet our observations lack full seasonal coverage. However, Wagner et al. (2003) also report absence of any correlation of CH_4 fluxes with soil temperature from a site in the Lena delta. The relation of CH_4 flux to water table depth is approximately exponential, the fluxes decrease very rapidly with lower water table. The river water stage appears particularly important for parts of the floodplain.

The implication for climate change effects on CH_4 fluxes from tundra landscapes is that an increase of precipitation and river water discharge will have a stronger influence than the increase of temperature. In particular changes in river regime will have a comparatively large influence, since the CH_4 production on the floodplain is comparatively large (Van Huissteden et al., 2005).

5.2 Carbon dioxide fluxes

The small-scale heterogeneity that is so prominent in the methane fluxes, is much less pronounced for carbon dioxide fluxes. Both the ecosystem respiration rates and photosynthesis rates, measured with chambers and leaf cuvettes, were variable in time and between sites, but the amount of variation is in the order of a factor of two, and not orders of magnitude, as for methane. The variation in *GPP* and R_{eco} could not be well explained by vegetation type, water table depth or active layer thickness. Moreover, soil temperature appears to determine the ecosystem respiration rates only to a limited extend (Fig. 6), as was also observed for methane. We hypothesize that small scale variations in hydrology, soil temperature, soil composition, organic matter content, active layer depth and soil moisture/water table depth are interrelated in such a complex way that the current measurements are insufficient to untangle their individual influences.

Photosynthesis appears mainly limited by radiation (Fig. 4) and much less by temperature or vapour pressure deficit. Based on the sunhour anomalies (Sect. 2.1.1),

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where 2004 had significantly less and 2005 significantly more sunhours than climatologically normal, whereas 2003 and 2006 are close to normal, the photosynthesis rates would be expected to change accordingly. This is however, not the case. Instead, photosynthesis rates are largest in 2003 and relatively small in 2006. The explanation of this apparent inconsistency is that the sunhour anomalies do not correlate well with the relative frequency of global radiation levels below the threshold of $200\,W\,m^{-2}$, when photosynthesis becomes severely radiation limited (Fig. 9). Instead, the relative frequency of radiation limitation occurs more than average in 2003 and below average in 2006. In 2004 and 2005, this frequency is close to the mean. This implies that photosynthesis rates indeed mainly depend on the occurrence of radiation limitation. During daytime, severe cloudiness is required to reduce global radiation levels below the threshold of $200\,W\,m^{-2}$. In this perspective, the absence of large mountain ranges in northern zone of Siberian tundra may be a relevant difference with Alaska, considering the relationship between topography and frontogenesis that was shown for Alaska (Lynch et al., 2001). Photosynthesis rates are not often limited by temperature, except for temperatures below 4°C. Figure 4b shows that in the temperature range between 10 and 20°C, not much gain in maximum photosynthesis capacity may be expected. Similarly, vapour pressure deficit does not often limit photosynthesis and high vapour pressure deficits are actually quite rare.

5.3 Winter fluxes

Because our measurement setup depends on solar energy, and because of the harsh climate and the inaccessibility of the area, we were unable to measure carbon dioxide and methane fluxes in the winter period. Where previously it was thought that carbon dioxide and methane emissions from frozen soils are negligible, evidence is accumulating that they may actually make up a considerable part of the annual balances. Wintertime carbon dioxide emissions may be $1.3–10.9\,g\,C\,m^{-2}\,winter^{-1}$ in Alaska (Fahnestock et al., (1998), $8.1\,g\,C\,m^{-2}\,winter^{-1}$ in Greenland (Soegaard et al., 2000), $4–6\,g\,C\,m^{-2}\,winter^{-1}$ in Vorkuta, European Russia (Heikinnen et al., 2002). These numbers often

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amount to about 20% of the annual sum (Chapin et al., 2000). It should be mentioned though that due to the more continental climate in Chokurdakh, the soil temperatures of $\pm-14°C$ in the springs of 2004, 2005 and 2006 and around $\pm-10°C$ in 2007 (probably as a result of the deeper snow) are much colder than observed in Alaska (-5.6 to $-3.6°C$, Fahnestock et al., 1998). Winter time methane emissions may be between 0.2 and 0.8 g C-CH$_4$ m^{-2} winter^{-1} with a peak emission of 7.8 g C-CH$_4$ g C m^{-2} winter^{-1} on wet flarks. Zimov et al. (1997) present a flux of 1.13 g C-CH$_4$ m^{-2} winter^{-1} from Siberian lakes. Thus winter fluxes of carbon dioxide may be relatively small compared to the summer time *NEE*, but winter time methane fluxes may contribute up to an extra 25% of the summer fluxes. The relative importance of methane emissions may be explained by the anaerobic conditions that may prevail in frozen, snow covered soils (Corradi et al., 2005). Also, microbial production of CH$_4$ has been shown to continue at subzero temperatures in arctic soils (Rivkina et al., 2000, 2007; Panikov and Sizova, 2006; Wagner et al., 2007). Assuming a winter carbon dioxide flux of 5 g C m^{-2} yr^{-1} and a winter methane flux of 1 g C m^{-2} yr^{-1}=9 g C-CO$_2$e m^{-2} yr^{-1}, would change the *NEE* to -87 g C m^{-2} yr^{-1}, the methane flux to 4.1 g C m^{-2} yr^{-1} with a GWP of 37 g C-CO$_2$e m^{-2} yr^{-1}, and the GHG balance to -50 g C-CO$_2$e m^{-2} yr^{-1}.

5.4 Short-term and long-term sensitivity

On the short-term, the carbon dioxide balance of this tundra ecosystem may be influenced primarily photosynthesis rates as a function of cloud-radiation interactions and by respiration rates via temperature. Another short-term change may be through changes in the length of the growing season. All other things remaining equal, Fig. 14 suggests that longer growing seasons are in favour of a stronger carbon dioxide sink if thawing starts earlier, but in favour of a reduced sink if the end of the growing season is postponed. This difference is because respiration is limited by temperature at the start, whereas photosynthesis is limited by sunlight at the end of the season. Apart from light limitation, phenology also determines photosynthesis rates, as shown in Fig. 15, where

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the ratio of the actual (observed) *GPP* over the potential (simulated) *GPP* is shown. As the model was calibrated for the mid summer period and plant properties were kept constant, the simulated *GPP* represents the maximum photosynthesis rate under the given weather conditions. Short-term variations due to weather are thus accounted for in both terms, so that consistent variations in the ratio are caused by phenology, such as increasing leaf area and photosynthetic capacity. Methane fluxes are most directly affected by changes in the hydrological cycle (Moore et al., 1993; Walter et al., 1996).

On the longer-term, changes in climate may impact the carbon dioxide balance through changes in vegetation composition and permafrost conditions, whereas methane fluxes depend on vegetation composition as well and on hydrology. Zimov et al. (2006) show that the carbon rich tundra soils are sensitive to enhanced thawing and that large amounts of carbon may be released initially. In their current form, we should be very careful to extend our observations to the longer-term. This may be done using dynamic vegetation models in combination with long-term satellite records. The terms in the carbon and greenhouse gas balances appear rather robust and the balance is not so tight that a shift from sink to source should be expected in the near future for this northern continuous permafrost site. The mean annual temperature of −10.5°C is cold enough to sustain permafrost conditions, even in a warming climate. However, a temperature increase of a few °C may be enough to accelerate permafrost dynamics and to increase the rate at which thermokarst lake thaw and refreeze, thereby releasing large amounts of carbon dioxide and methane (Zimov et al., 1997; Stokstad, 2004; Zimov et al., 2006). At sites with discontinuous permafrost, permafrost melting has shown to cause a large increase of methane emissions by a change from *Sphagnum*-dominated vegetation on hummocks to *Carex*-dominated vegetation in hollows (Turetsky et al., 2002; Christensen et al., 2004). Locally, similar changes have been observed at our site, where drier ridges or palsas (TD type) with a thin active layer partly melt at their edges and are replaced by wet *Carex* vegetation (TW1 type) with a thicker active layer. This occurs over very small areas however, at the site the permafrost appears to be stable.

For the CH_4 fluxes, our data suggest that the long-term sensitivity depends on hydrological changes. The floodplain fluxes dominate the CH_4 emission of the area and prove to be strongly related to river discharge. Also the emission from the TW1 *Carex* vegetations outside the floodplain depend on water table and thus precipitation sur-
plus. By contrast, effects of soil temperature differences are absent. Climate change will increase CH_4 fluxes by increase of precipitation, not by increase of temperature, unless the temperature increase will cause instability of the permafrost.

On the long-term changes in carbon cycling may be caused for an important part by changes in vegetation composition. In this respect, such changes have already been
observed to occur in Alaska (Jia et al., 2003; Lloyd et al., 2003; McGuire et al., 2003; Wilmking et al., 2006) as well as for western Siberia (Esper and Schweingruber, 2004), but similar studies in NE Siberia are lacking.

5.5 Surface water percentage and up scaling

In Sect. 2.1.2 it was mentioned that near 20% of the area surrounding the site con-
sists of surface water. This may have implications for up scaling the greenhouse gas balance of the site for two reasons: first the direct effect of reducing the land area by taking the percentage surface water into account: the carbon dioxide exchange between lakes and the atmosphere is probably smaller than between land and the atmosphere, whereas methane emissions from lakes may be larger than from dry land
(Bartlett et al., 1992). Consequently, on a larger scale, the greenhouse gas balance of the larger area is probably more neutral than presented in Fig. 14. Flux measurements from lakes are not yet available. Second, lakes and rivers are often bordered with floodplains and marshlands with high CH_4 emission rates. In the ORCHIDEE-PEATLAND-VU simulations we have considered only the vegetation types surrounding
the flux tower. However, the methane emissions from floodplain areas are much larger than from drier parts, so that the up scaling is quite sensitive to the amount of floodplain and lake surface in the area. At present, our knowledge about the distribution of methane fluxes over different parts of the floodplain is limited and, perhaps more

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importantly, information about which fraction of the larger area is covered with flood-plain is entirely lacking. The latter is a function of the length of the river, the relief of the floodplains and the river level. For these reasons it is as yet unfeasible to give a realistic estimate of the methane fluxes of the larger area.

5.6 Comparison with other sites

A comparison of the carbon dioxide and methane fluxes observed in Chokurdakh with those observed at other arctic tundra sites is given in Table 7. The mean daily carbon dioxide flux is within the range observed for different vegetation types in Greenland. On an annual scale, the carbon dioxide sink is quite a bit larger than observed at other sites. Some sites in Alaska and Greenland even act as sources of carbon dioxide. The smaller *NEE* at other sites may be explained by the smaller *GPP* at Vorkuta (Heikinnen et al., 2002) and the larger R_{eco} at Toolik Lake (Oberbauer et al., 1998). The smaller *NEE* at Greenland (Soegaard et al., 2000) may be explained by the shorter growing season there. Possibly, a general explanation of the large *NEE* is that the site expe-riences a more continental climate than other sites, so that ecosystem respiration is limited by the cold soils, with temperatures lagging behind the air temperatures, but with warm summers, which stimulates photosynthesis. Daily methane fluxes are vari-able between the sites, due to vegetation, and change substantially during the season. Nevertheless, the daily methane fluxes are well in the range found in the literature, be it often on the larger side. Annual fluxes are also quite comparable between sites, with the exception of the fluxes measured in the floodplains of the Kolyma river (Corradi et al., 2005), which are much higher, probably as a result of the high water table and the high nutrient availability.

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6 Conclusions

At an arctic tundra site in North East Siberia, near the village of Chokurdakh in the lowlands of the Indigirka river, we observed an mean annual carbon dioxide flux of $-92 \, \mathrm{g \, C \, m^{-2} \, yr^{-1}}$, which is the net result of $232 \, \mathrm{g \, C \, m^{-2} \, yr^{-1}}$ of uptake by photosynthesis and a release of $141 \, \mathrm{g \, C \, m^{-2} \, yr^{-1}}$ as ecosystem respiration. The mean annual methane emission amounts to $28 \, \mathrm{g \, C\text{-}CO_2 e \, m^{-2} \, yr^{-1}}$ (=$4.1 \, \mathrm{g \, CH_4 \, m^{-2} \, yr^{-1}}$), so that the net greenhouse gas balance becomes $-64 \, \mathrm{g \, C\text{-}CO_2 e \, m^{-2} \, yr^{-1}}$. Because the emitted methane is removed from the atmosphere by oxidation, the radiative effect of the sustained carbon dioxide sink dominate over time horizons longer than 13 years, which, considering the old age of the site, means that the site acts to cool the global climate. The greenhouse gas balance would probably be more neutral if the percentage of lakes and floodplains would be taken into account. The net carbon dioxide flux is large compared to other arctic tundra sites, probably as a result of the more continental climate. On the short-term, photosynthesis appears to depend most on the frequency of radiation limitation due to severe cloudiness, ecosystem respiration rates depend on temperature, but also on water level, active layer depth and time in the growing season. Methane fluxes are highly variable on small spatial scales. This heterogeneity is primarily related to depth of the water table and on the occurrence of vegetation types with aerenchyma, to transport methane from the soil to the atmosphere. Further variation may be explained by the exudation of labile organic compounds by plant roots, which is related to photosynthesis rates, soil temperature, and active layer depth. The methane fluxes are insensitive to soil temperature but depend strongly on changes in hydrologic conditions. Potential positive feedbacks between climate change and arctic methane fluxes is governed by precipitation increase rather than warming for this continuous permafrost area.

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Table 1. Basic climatology of as observed at the weather station near the village of Chokurdakh (WMO station 21946 Chokurdakh) between 1999–2006. Snow depths at the field station are usually 60–80 cm. Precipitation is reported for 1994–1999.

		Jan	Feb	March	April	May	June	July	Aug	Sep	Oct	Nov	Dec	mean total
temperature	°C	−34.2	−34	−26.6	−17.9	−5.1	6.9	10.4	7.8	1.3	−10.8	−23.2	−30.5	−13
precipitation	mm	9.7	9.7	9.4	6.0	13.5	26.5	22.6	35.2	28.1	21.8	18.6	10.5	*211.7*
wind speed	m s^{-1}	3.3	3.6	3.3	3.7	4.4	4.8	4.6	4.0	3.9	3.6	3.4	3.9	3.9
sun hours	hr day^{-1}	0.0	1.4	4.7	8.0	7.8	9.2	8.0	4.4	2.1	1.4	0.3	0.0	3.9
frost days	days	31	28	31	30	28.8	7.8	0.3	2.6	17.2	30.9	30	31	*268.5*
snow days	days	30.6	27.9	30.3	26.9	27.8	5.3	0	0	8.1	26.3	29.5	30.2	*242.9*
show depth	cm	26.5	31.6	37.1	38.3	30.8	1.5	0	0	2.7	10.1	20.6	23.8	18.6

Table 2. Site classification based on geomorphology, water table position and vegetation. Areal fraction with respect to total F (Floodplain) or T (terrace) area.

Code	Geomorphology			Soil	Areal fraction	water table (cm)	active layer (cm)	Vegetation, periglacial surface features
FD1	Floodplain	Dry	Levee	mineral	0.15	20	52-54	Tall *Salix* shrub with moss and lichen understorey
FD2		Dry	Levee-backswamp transition	mineral	0.22	May-20	25-50	Sedges and grasses, sometimes also willow and mosses; actively growing ice wedge polygons
FW1		Wet	Low backswamp	organic-mineral	0.24	0	40-43	Reddish, short grass vegetation (*Arctophila fulva*) with some moss and *Caltha palustris*
FW2		Wet	Backswamp	organic-mineral	0.39	0	30-55	Tall sedges, rushes, *Eriophorum*
TD1	Terrace	Dry	Palsa, polygon ridge	organic on mineral	0.32	-	15-25	*Betula nana*, grasses, mosses and lichens. Cracks, small polygons, mud boils.
TD2		Dry	Tussock tundra	organic on mineral	0.15	-	15-25	*Eriphorum* tussocks, mosses, lichens, species-rich vegetation
TD3		Dry	Dry grass flat	organic on mineral	0.05	-	12-20	Short grasses
TD4		Dry	Dry Shagnum flat	organic	0.15	-	20	Sphagnum peat overlying permafrost ice at shallow depth. *Sphagnum* has a dry appearance.
TW1		Wet	Depression, diffuse drainage	organic on mineral	0.13	0-5	26-48	Tall sedges, *Eriophorum*
TW2		Wet	Thawing ice wedge depression	organic, ice	<0.01	0	28-50	Standing water, some *Sphagnum* and *Carex*
TW3		Wet	Low polygon centre, other depressions	organic	0.13	15-20	22-48	*Sphagnum,* lush green appearrance, often overgrowing TW1.
TW4		Wet	Low polygon centre, other depressions	organic	0.05	0-5	22-38	*Potentilla palustris, Sphagnum* and sedges.

Table 3. Soil characteristics.

component	fraction	density kg/m3	specific heat kJ/kg/K
mineral	4%	2650.00	0.90
organic	36%	1300.00	1.92
water	30–60%	1000.00	4.18
air	remaining	1.20	1.01

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Table 4. Parameters of the ecosystem respiration model (Eq. 2).

p_3	p_2	p_1	p_0	Q_{10}
–8.37E–06	5.15E–03	–1.04E+00	6.94E+01	1.80

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Table 5. Comparison of units of methane fluxes.

carbon flux units $gC\ m^{-2}\ day^{-1}$	methane mass flux units $gCH_4\ m^{-2}\ day^{-1}$	GWP units (*) $gCO_2e\ m^{-2}\ day^{-1}$	GWP in carbon flux units $gC\text{-}CO_2e\ m^{-2}\ day^{-1}$
12/16 = 0.75	1	25	25×(12/44) = 6.8
0.11	0.15	4.0	1

Table 6. Up scaled methane fluxes, based on estimated areal fraction of terrain types on the Terrace/tundra near the eddy correlation tower and on the floodplain. Fluxes are given as mg CH_4 m^{-2} h^{-1}.

vegetation class	2004	2005	2006
FD1	−0.01±0.01	−±−	−0.10±0.11
FD2	0.46±0.60	−0.04±0.18	−0.14±0.23
TD1	−0.10±0.16	−0.08±0.14	−0.15±0.18
TD2	−0.02±0.01	0.01±0.00	−0.18±0.06
TD3	0.06±0.09	0.001±0.001	0.02±0.003
TD4	−0.06±0.02	0.01±0.23	−0.04±0.10
FW1	7.38±0.77	1.09±1.99	3.47±1.82
FW2	6.79±3.83	3.61±4.26	2.91±1.67
TW1	1.60±0.60	0.69±0.47	0.82±0.40
TW2	0.03±0.02	0.02±0.02	−±−
TW2	0.41±0.35	0.19±0.25	0.03±0.20
TW3	0.33±0.15	0.05±0.08	0.22±0.20
Floodplain	13.66±6.63	4.76±6.51	6.41±4.10
Terrace/Tundra	2.46±1.73	1.08±1.29	1.11±1.28

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Table 7. Comparison of daily and annual carbon dioxide and methane fluxes measured at arctic tundra. In some occasions, hourly fluxes were integrated to daily fluxes by multiplying with 24 h per day.

region	site	vegetation type	remarks	NEE daily gC $m^{-2} d^{-1}$	NEE annual gC $m^{-2} yr^{-1}$	CH$_4$ flux daily gC-CO$_2$e $m^{-2} d^{-1}$	CH$_4$ flux annual gC-CO$_2$e $m^{-2} yr^{-1}$	reference
Alaska	Barrow	wet sedge tundra		-	-55.5	-	-	Kwon et al., 2006
Alaska	Barrow	moist tussock tundra		-	+18.3	-	-	Kwon et al., 2006
Alaska	Yuken-K. Delta	wet meadow		-	-	0.90	-	Bartlett et al., 1992
Alaska	Yuken-K. Delta	dry upland tundra		-	-	0.01	-	Bartlett et al., 1992
Alaska	Yuken-K. Delta	large tundra lakes		-	-	0.02	-	Bartlett et al., 1992
Alaska	Yuken-K. Delta	small tundra lakes		-	-	0.48	-	Bartlett et al., 1992
Alaska	Toolik Lake	tussock tundra		0.32	-	0.01-0.06	-	Oberbauer et al., 1998
Alaska	Toolik Lake	wet sedge meadow	control	-	-	0.50	30.7	King et al., 1998
Alaska	Toolik Lake	wet sedge meadow	moss removal	-	-	0.53	36.1	King et al., 1998
Alaska	Toolik Lake	wet sedge meadow	sedge removal	-	-	0.04	3.8	King et al., 1998
Greenland	Zackenberg	Cassiope heath		-0.60	-	0.00	-	Christensen et al., 2000
Greenland	Zackenberg	hummocky fen		-4.80	-	0.90	-	Christensen et al., 2000
Greenland	Zackenberg	continous fen		-2.88	-	1.35	-	Christensen et al., 2000
Greenland	Zackenberg	grasland		-4.80	-	0.41	-	Christensen et al., 2000
Greenland	Zackenberg	salix arctica		-0.60	-	0.00	-	Christensen et al., 2000
Greenland	Zackenberg	overall		-2.30	-	0.29	-	Christensen et al., 2000
Greenland	Zackenberg	fen area	July	-	-	0.75	23.2	Friborg et al., 2000
Greenland	Zackenberg	fen area	Aug	-	-	0.06-0.09	-	Friborg et al., 2000
Greenland	Zackenberg	integrated	pre-season	-	8.4	-	-	Soegaard et al., 2000
Greenland	Zackenberg	integrated	growing season	-	-18.8	-	-	Soegaard et al., 2000
Greenland	Zackenberg	integrated	winter	-	8.1	-	-	Soegaard et al., 2000
Greenland	Zackenberg	integrated	all year	-	-2.3	-	-	Soegaard et al., 2000
Greenland	Zackenberg	fen + grass	all year	-	-18.8	-	-	Soegaard et al., 2000
Greenland	Zackenberg	heath	all year	-	5.2	-	-	Soegaard et al., 2000
Greenland	Zackenberg	willow	all year	-	0.6	-	-	Soegaard et al., 2000
Sweden	Stordalen	bogs+mires	eddy covariance	-	-	0.46	-	Christensen et al., 2004
Sweden	Stordalen	bogs+mires	chambers	-	-	0.40-0.45	-	Christensen et al., 2004
Eur. Russia	Vorkuta	tundra wetland		-	-29.0	-	4.9	Heikinnen et al., 2002
Eur. Russia	Vorkuta	tundra wetland		-	-34.6	0.28	-	Heikinnen et al., 2004
Siberia	transsect	various	mean	-	-	0.29	-	Christensen et al., 1995
Siberia	Lena delta	Carex aquatilis	depression	-	-	0.33	-	Wagner et al., 2003
Siberia	Lena delta	mosses	rim	-	-	0.03	-	Wagner et al., 2004
Siberia	Lena delta	polygon depression	45% of area	-	-	0.18	-	Kutzbach et al., 2004
Siberia	Lena delta	polygon rim	55% of area	-	-	0.03	-	Kutzbach et al., 2004
Siberia	Chokurdakh	mixed moist tundra		-1.53	-92	0.47	28	*this study*
Siberia	Cherskii	horsetail grassland		-	-	1.44	-	Tsuyuzaki et al., 2001
Siberia	Cherskii	Carex grassland		-	-	0.46	-	Tsuyuzaki et al., 2001
Siberia	Cherskii	Eriophorum grassland		-	-	-0.01	-	Tsuyuzaki et al., 2001
Siberia	Cherskii	tussock tundra		-	-38.0	-	100	Corradi et al., 2005

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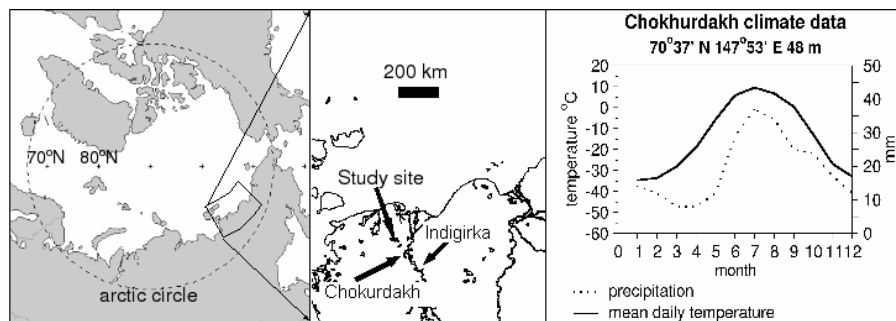
Fig. 1. Site location and climate of the Kytalyk site.

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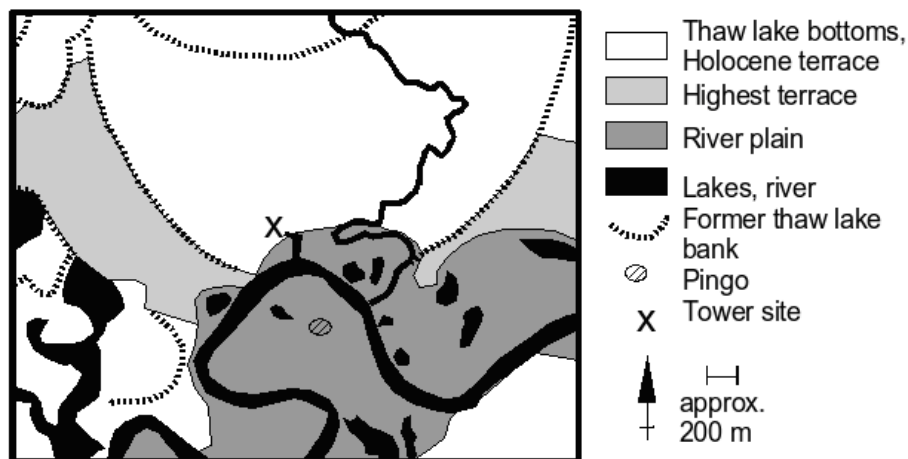
Fig. 2. Geomorphological map of the site, derived from field mapping and a satellite picture of the Kytalyk reserve on http://maps.google.com (http://maps.google.com/?ll=70.8275, 147.4790&spn=0.1578,0.5905&t=k).
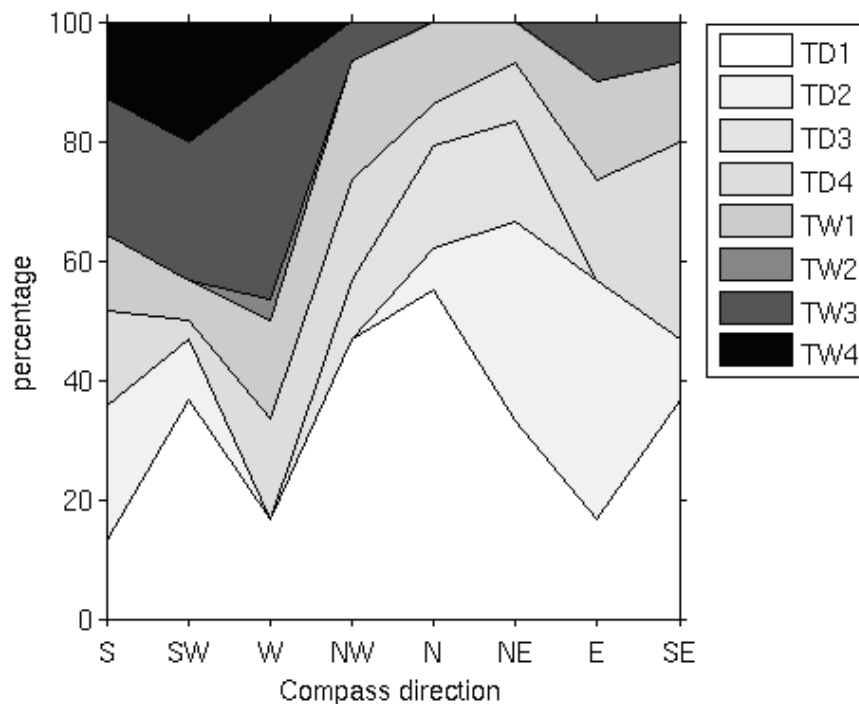
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Fig. 3. Areal proportion of vegetation/terrain classes along 150 m transects in different directions from the tower, TD classes: dry sites with a grass-moss-*Betula* nana vegetation (TD1,TD3) or *Eriophorum* hummocks (TD2). TD4 is a dried *Sphagnum* vegetation with very thin active layer. TW classes: wet sites with a *Carex-Eriophorum* vegetation (TW1, "sedge meadows"), open water (TW2) or *Sphagnum*-dominated vegetation (TW3,4). See Table 2 for further description of terrain classes.

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$$y = (0.998 \pm 0.005) \cdot x - (14.0 \pm 0.8)$$
$$\sigma_y = 60.6$$
$$N = 9835$$
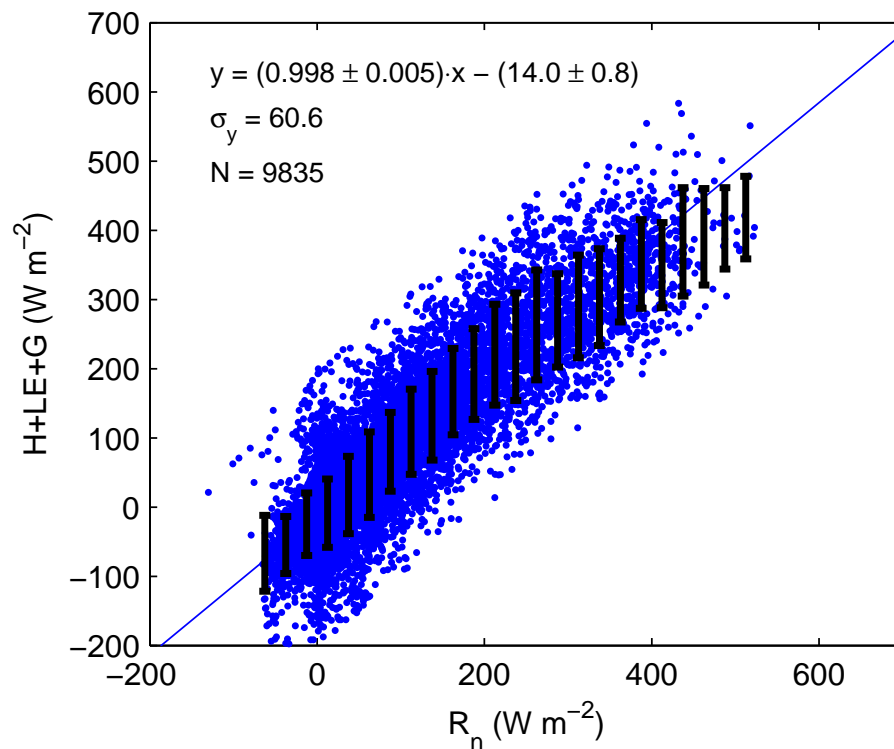
Fig. 4. Energy balance closure based on half hourly fluxes. The x-axis indicates the amount of net radiation (Rn) received and the y-axis the amount of energy spent on the sensible (H), latent (LE) and soil heat fluxes (G).

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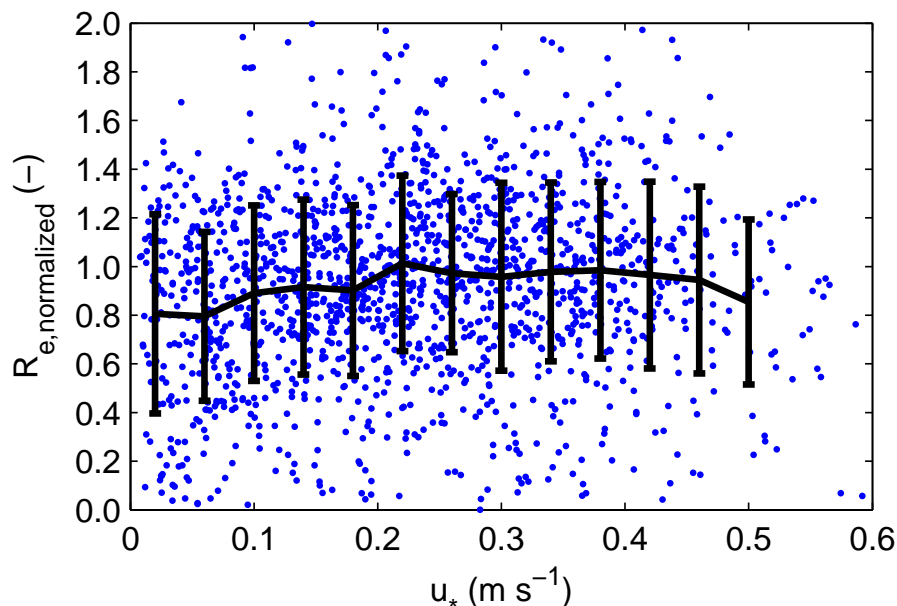
Fig. 5. The ecosystem respiration rate, determined from nighttime ($R_g < 20\,\mathrm{W\,m^{-2}}$) eddy co-variance measurements, normalized for temperature and time influences, as a function of u_*, showing a slight underestimation for $u_* < 0.2\,\mathrm{m\,s^{-1}}$.
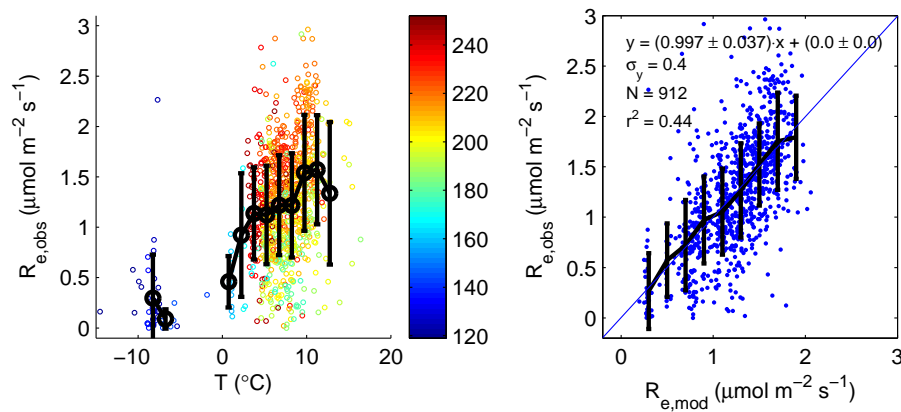
Fig. 6. (a) the variation of observed soil respiration rate as a function of soil temperature and time. **(b)** performance of the soil respiration model.

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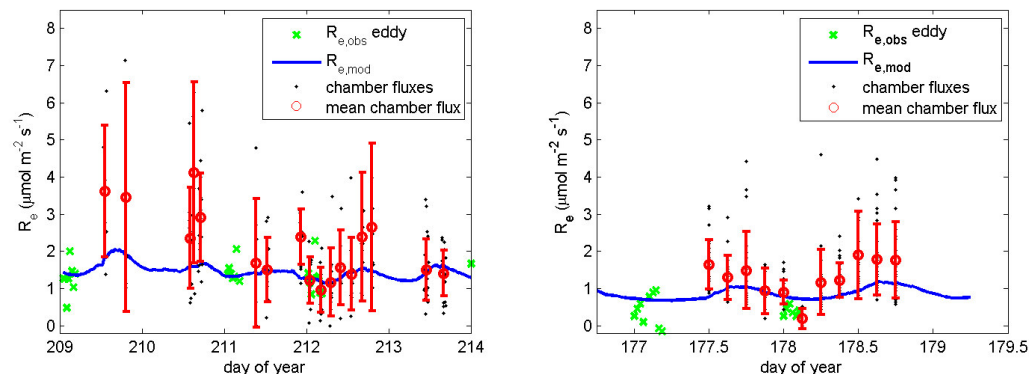
Fig. 7. Comparison of $R_{eco,obs}$: observed respiration rate (nighttime eddy fluxes when $u_* > 0.2\,\mathrm{m\,s}^{-1}$), $R_{eco,mod}$ (Eq. 1), chamber fluxes at individual locations and the mean chamber fluxes. The left figure shows data for 27 July to 1 August 2004 and the right figure for 26 to 29 June 2005.
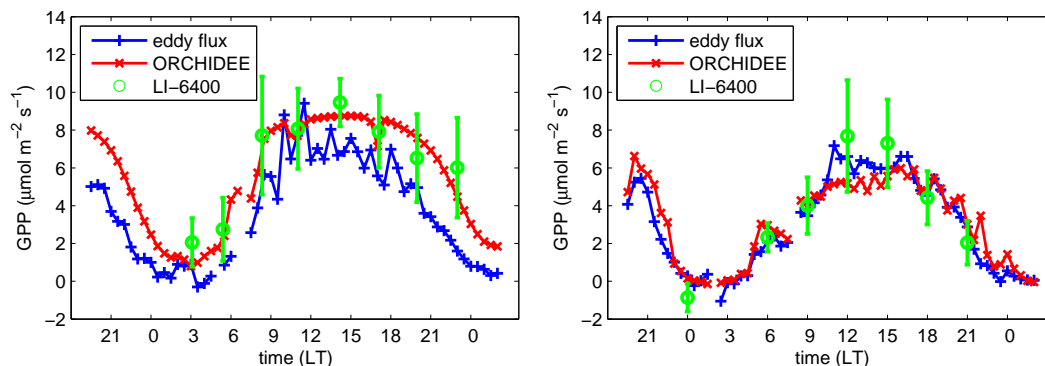
Fig. 8. Diurnal cycles of *GPP* determined by 1) partition of eddy covariance measurements of *NEE*, 2) model simulations with ORCHIDEE, using local meteorology and 3) leaf level measurements. The left figure shows the diurnal cycles for 14 July 2003, and the right figure for 30 July 2004.

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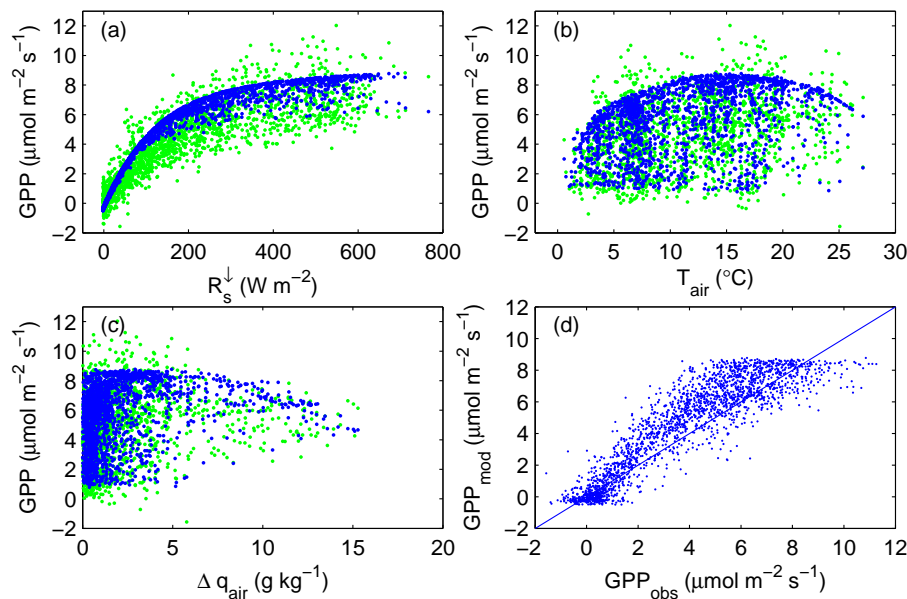
Fig. 9. Performance of the ORCHIDEE model to simulate *GPP* as a function of **(a)** global radiation, **(b)** air temperature **(c)** vapour pressure deficit. The blue dots represent the modelled *GPP* and the green ones the *GPP* obtained by partitioning the observed *NEE*. **(d)** shows modelled versus observed *GPP*. In this comparison data only mid summer data have been used (9 July to 18 August).

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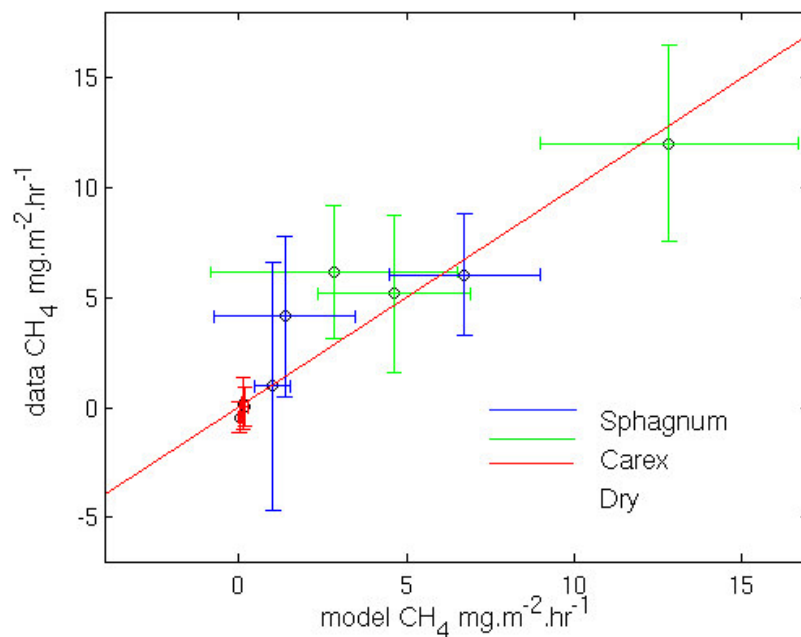
Fig. 10. Validation of the methane fluxes simulated by the PEATLAND-VU model versus observations. Each data point represents the weekly average of methane measurements for the indicated vegetation type in a specific year.

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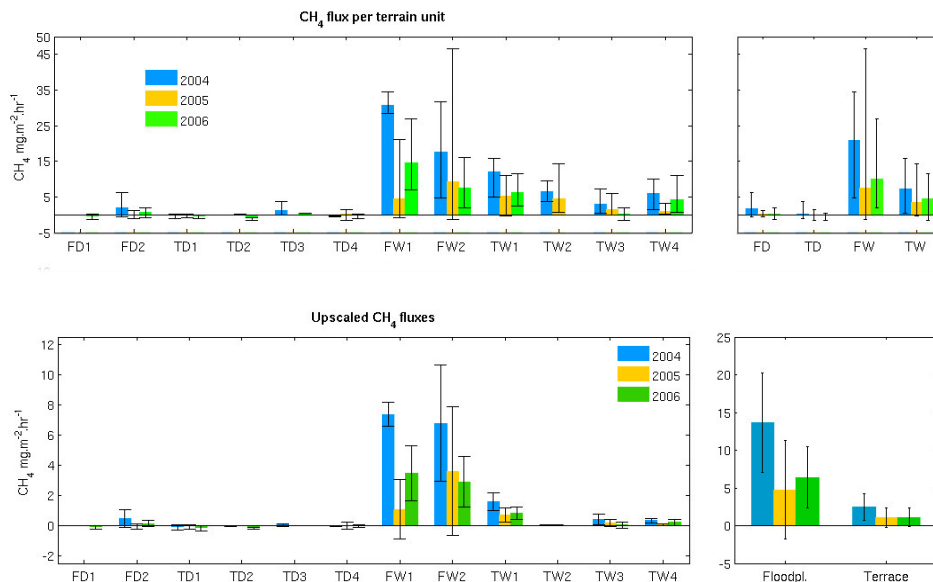
Fig. 11. Top: Average CH_4 fluxes per terrain/vegetation unit. The error bars indicate the standard deviation of the average. Top right: averages for all wet and dry groups of terrace and floodplain. Bottom: upscaled fluxes by weighting of the measured fluxed by areal proportion in the terrain.
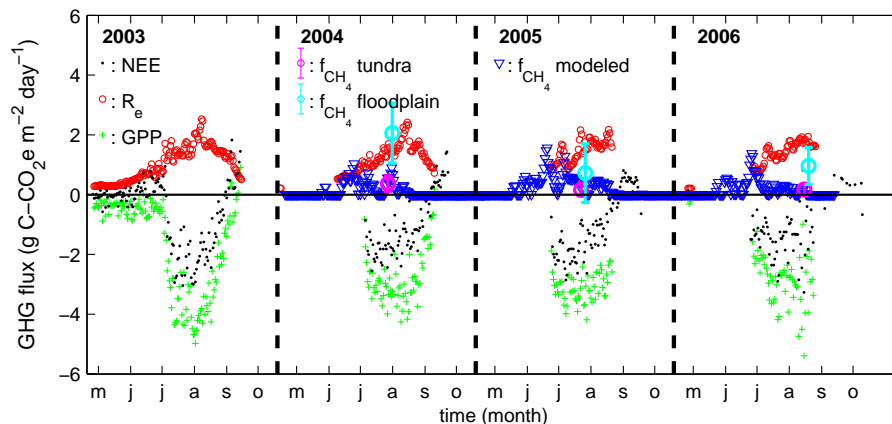
Fig. 12. The seasonal cycles of *GPP*, R_{eco}, *NEE* and f_{CH4}. The individual data points represent daily total fluxes, except for the observed methane fluxes (data points with errorbars) which represent weekly averages. The time axis is compressed in the winter months.

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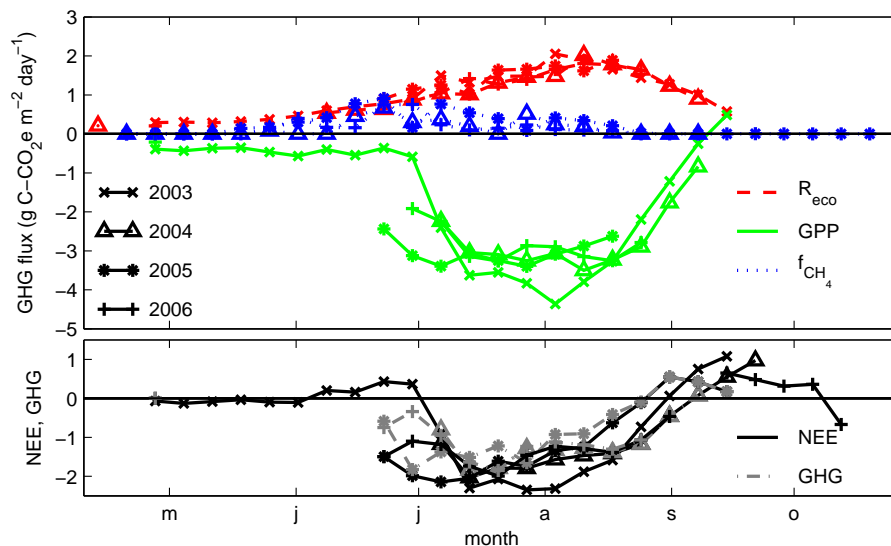
Fig. 13. The interannual variability in the seasonal cycles of R_{eco}, *GPP*, f_{CH4} (upper panel) and *NEE* and the net GHG balance (lower panel). The data points represent weekly mean daily fluxes. The lower panel has the same scale as the upper panel.

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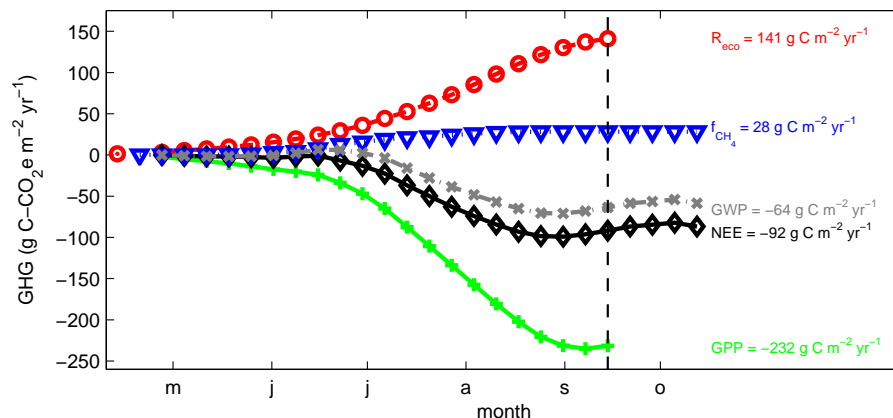
Fig. 14. Annually cumulative *GPP*, R_{eco}, *NEE* and methane fluxes and the resulting GHG balance. The numbers at the right indicate the total flux in mid September, indicated by the vertical dashed line, when eddy covariance measurements were no longer available.
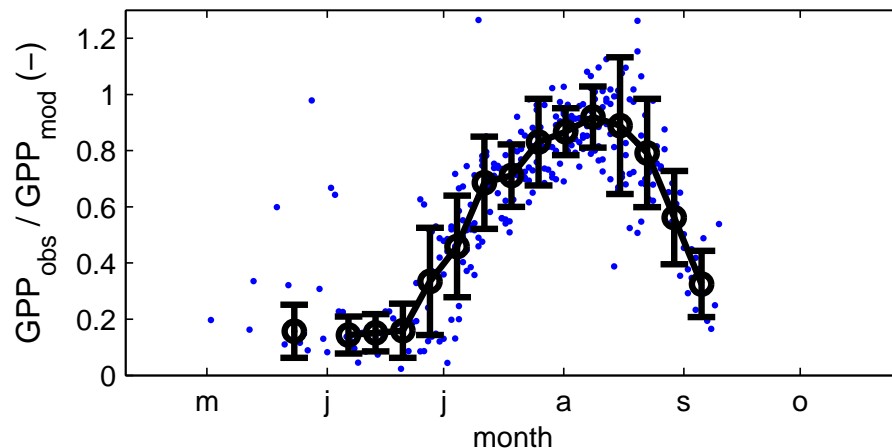
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Fig. 15. The ratio of observed and simulated *GPP*, which represents the percentage of the maximum *GPP* that is achieved, as a function of the day of year. *GPP*mod was calibrated for the days in the middle of the growing season and phenology was not simulated.