

Supplement of Biogeosciences Discuss., 11, 9215–9247, 2014
<http://www.biogeosciences-discuss.net/11/9215/2014/>
doi:10.5194/bgd-11-9215-2014-supplement
© Author(s) 2014. CC Attribution 3.0 License.



Supplement of

Using atmospheric observations to evaluate the spatiotemporal variability of CO₂ fluxes simulated by terrestrial biospheric models

Y. Fang et al.

Correspondence to: Y. Fang (yyfang@stanford.edu)

10 **S1 Terrestrial biospheric models (TBMs)**

11 Four TBMs participating in the North American Carbon Program Regional Interim Synthesis
12 (NACP RIS) project (Huntzinger et al., 2012) are used in the analyses presented in this work.
13 These TBMs were selected because their NEE fluxes are available at 3-hourly and 1°×1°
14 resolution. The four models are the CASA coupled with the Global Fire Emissions Database
15 (CASA-GFED, Van Der Werf et al., 2006), Simple Biosphere (SiB3, Baker et al., 2008),
16 Organising Carbon and Hydrology In Dynamic Ecosystems (ORCHIDEE, Krinner et al., 2005)
17 and Vegetation Global Atmosphere and Soil (VEGAS2, Zeng et al., 2005). The runs used here
18 represent “off the shelf” simulations, and are therefore not based on a standardized protocol.
19 Any differences in their performance can therefore be driven not only by structural differences,
20 but also by differences in initial conditions, spin up, driver data etc. Table S1 provides a brief
21 summary of the key features of each model, with more detail available in Huntzinger et al.
22 (2012).

23 Table S1. Terrestrial Biospheric Models (TBMs) evaluated and their phenology, resolution,
24 photosynthetic and soil carbon decomposition formulations

Model	Phenology	Native temporal resolution	Native spatial resolution	Photosynthetic formulation	# Plant functional types	# Soil pools
CASA-GFED	Diagnostic	Monthly	1°	Light Use Efficiency	3	5
ORCHIDEE	Prognostic	30 min	0.5°	Enzyme Kinetic	12	8
SiB3	Diagnostic	Hourly	1°	Enzyme Kinetic	14	0
VEGAS2	Prognostic	Daily	1°	Light Use Efficiency	4	6

25

26 **S2 Covariance matrices (R and Q) used in the synthetic and real data experiments**

27 The model-data mismatch covariance matrix \mathbf{R} describe the expected magnitude of discrepancies
28 between the observed and modeled CO₂ concentrations. These errors are assumed here to be
29 uncorrelated in space and time, and \mathbf{R} is therefore a diagonal matrix with individual elements
30 representing variances (σ_R^2) that vary across measurement towers and months. The prior flux
31 covariance matrix \mathbf{Q} characterizes the spatially- and temporally-correlated flux deviations from
32 the model of the trend, and is modeled using a covariance function that varies as a function of the
33 separation distance between flux times and location, as in Gourdji et al. (2012):

$$\mathbf{Q} = \sigma_Q^2 \underbrace{\left[\exp\left(-\frac{\mathbf{h}_t}{l_t}\right) \right]}_{\text{temporal covariance}} \otimes \underbrace{\left[\exp\left(-\frac{\mathbf{h}_s}{l_s}\right) \right]}_{\text{spatial covariance}} \quad (\text{S1})$$

34 where σ_Q^2 is the asymptotic variance of flux deviation in space and time, \mathbf{h}_t and \mathbf{h}_s represent the
35 separation lags between estimation locations in space and time, respectively, and l_s and l_t are the
36 spatial and temporal correlation length parameters. The variance and correlation parameters vary
37 across months. Temporal correlations are only assumed across days for the same times of the
38 day, and not within days, so as not to risk smoothing out the diurnal variability.

39 **S2.1 Covariance matrices (R and Q) used in the real data experiments**

40 The covariance parameters for \mathbf{R} and \mathbf{Q} for the RD-one- $\xi\epsilon$ and RD-all- $\xi\epsilon$ experiments (see
41 Section 4 and Figure 2 in the main text) are estimated using Restricted Maximum Likelihood
42 approach (e.g., Gourdji et al., 2010; Gourdji et al., 2012; Michalak et al., 2004), which
43 minimizes the negative log-likelihood of the available atmospheric measurements with respect to
44 the covariance parameters in \mathbf{R} and \mathbf{Q} . The corresponding objective function for a given
45 candidate model \mathbf{X}_c is (Kitanidis, 1995):

$$L = \ln|\Sigma| + \ln|(\mathbf{H}\mathbf{X}_c)^T \Sigma^{-1} \mathbf{H}\mathbf{X}_c| + [\mathbf{z}^T (\Sigma^{-1} - \Sigma^{-1} \mathbf{H}\mathbf{X}_c ((\mathbf{H}\mathbf{X}_c)^T \Sigma^{-1} \mathbf{H}\mathbf{X}_c)^{-1} (\mathbf{H}\mathbf{X}_c)^T \Sigma^{-1}) \mathbf{z}] \quad (\text{S1})$$

46 where all variables are as defined in Section 3 of the main document.

47 Because the \mathbf{R} and \mathbf{Q} parameters depend on the candidate model of the trend \mathbf{X}_c , and the
 48 selection of the model of the trend is affected by \mathbf{R} and \mathbf{Q} (Eq. 6-7), the model selection and
 49 parameter optimization proceed iteratively. The final optimized \mathbf{R} and \mathbf{Q} for each experiment are
 50 henceforth denoted as \mathbf{R}_{RML} and \mathbf{Q}_{RML} . Note that for the RD-one- $\xi\epsilon$ experiments, different \mathbf{R}_{RML}
 51 and \mathbf{Q}_{RML} are obtained based the \mathbf{X}_c that includes biome-month combinations for each individual
 52 TBM, and these are themselves different from the single \mathbf{R}_{RML} and \mathbf{Q}_{RML} obtained for the RD-
 53 all- $\xi\epsilon$ experiment based on the \mathbf{X}_c that include biome-month combinations from all four TBMs.

54 **S2.2 Covariance matrices (\mathbf{R} and \mathbf{Q}) used in the synthetic data experiments**

55 For the SD-one- $\emptyset\emptyset$ experiments that do not consider model-data mismatch, all variances in \mathbf{R} are
 56 set to a nominal value of $\sigma_R^2 = 0.01 \text{ ppm}^2$ for all towers and all months. The remaining
 57 synthetic data experiments (SD-one- $\emptyset\epsilon$, SD-one- $\xi\epsilon$ and SD-all- $\xi\epsilon$) all include realistic model-
 58 data mismatch errors, and the variances in \mathbf{R} are set to be equal to those used in the analogous
 59 real data experiments (Section S2.1).

60 For synthetic data experiments with no additional spatiotemporal variability added to the
 61 underlying flux field (SD-one- $\emptyset\emptyset$ and SD-one- $\emptyset\epsilon$), the variance of flux deviations from a trend
 62 including all TBM biome-month combinations is technically zero, whereas the
 63 variance/covariance of flux deviations from a trend that includes none of the TBM biome-month
 64 combinations would be equal to that of the full underlying flux field. This second setup

65 represents a more conservative assumption, *i.e.*, does not prescribe a priori that the variability in
 66 the candidate TBM represents the true underlying variability. Consistent with this setup, the
 67 parameters of the matrix \mathbf{Q} are set to those representing the full variability of the underlying
 68 fluxes, where these parameters are obtained by minimizing the negative log likelihood of the
 69 fluxes (Gourdji et al., 2010; Gourdji et al., 2008; Mueller et al., 2008):

$$L_{\mathbf{Q}} = \ln|\mathbf{Q}| + \ln|\mathbf{X}^T\mathbf{Q}^{-1}\mathbf{X}| + \frac{1}{2}[\mathbf{s}^T(\mathbf{Q}^{-1} - \mathbf{Q}^{-1}\mathbf{X}(\mathbf{X}^T\mathbf{Q}^{-1}\mathbf{X})^{-1}\mathbf{X}^T\mathbf{Q}^{-1})\mathbf{s}] \quad (\text{S2})$$

70 Here, \mathbf{X} is a simply column of ones, such that the covariance parameters represent the correlation
 71 structure of the full flux field. \mathbf{Q} estimated using this method is referred as \mathbf{Q}_{krig} , and is different
 72 for each TBM.

73 For synthetic data experiments with the presence of spatially-correlated flux residuals (SD-one-
 74 $\xi\epsilon$ and SD-all- $\xi\epsilon$), the \mathbf{Q} applied is \mathbf{Q}_{RML} derived from the RD-all- $\xi\epsilon$ experiment.

75 **S3 Flux residuals (ξ) and model-data mismatch errors (ϵ) in the synthetic data** 76 **experiments**

77 In all synthetic data experiments, measurements are generated as $\mathbf{z} = \mathbf{H}(\mathbf{s}_{\text{TBM}} + \xi) + \epsilon$, in which,
 78 \mathbf{s}_{TBM} is simulated NEE from a TBM, ξ represents any spatiotemporally-correlated flux residuals
 79 beyond the variability represented by the TBM, and ϵ represents the model-data mismatch errors.

80 For the SD-one- $\emptyset\emptyset$ experiments, ϵ is a zero vector. In the SD-one- $\emptyset\epsilon$, SD-one- $\xi\epsilon$, and SD-all- $\xi\epsilon$
 81 experiments, ϵ is a randomly-generated vector of independent normally-distributed values with
 82 variances corresponding to the diagonal elements (σ_R^2) of \mathbf{R}_{RML} for the analogous real data
 83 experiment and a mean of 0.

84 When no additional spatiotemporally correlated flux residuals are included (SD-one- $\emptyset\emptyset$ and SD-
 85 one- $\emptyset\epsilon$), ξ is a zero vector. In all SD cases that include realistic flux deviations (SD-one- $\xi\epsilon$, and

86 SD-all- $\xi\epsilon$), ξ is a randomly-generated vector of normally-distributed values with a covariance
87 structure equal to \mathbf{Q}_{RML} from the RD-all- $\xi\epsilon$ experiment.

88 **References**

- 89 Baker, I. T., Prihodko, L., Denning, A. S., Goulden, M., Miller, S., and Da Rocha, H. R.: Seasonal
90 drought stress in the Amazon: Reconciling models and observations, *Journal of Geophysical*
91 *Research: Biogeosciences*, 113, G00B01, 2008.
- 92 Gourdjji, S. M., Hirsch, A. I., Mueller, K. L., Yadav, V., Andrews, A. E., and Michalak, A. M.: Regional-
93 scale geostatistical inverse modeling of North American CO₂ fluxes: a synthetic data study,
94 *Atmos. Chem. Phys.*, 10, 6151-6167, 2010.
- 95 Gourdjji, S. M., Mueller, K. L., Schaefer, K., and Michalak, A. M.: Global monthly averaged CO₂ fluxes
96 recovered using a geostatistical inverse modeling approach: 2. Results including auxiliary
97 environmental data, *Journal of Geophysical Research: Atmospheres*, 113, D21115, 2008.
- 98 Gourdjji, S. M., Mueller, K. L., Yadav, V., Huntzinger, D. N., Andrews, A. E., Trudeau, M., Petron, G.,
99 Nehrkorn, T., Eluszkiewicz, J., Henderson, J., Wen, D., Lin, J., Fischer, M., Sweeney, C., and
100 Michalak, A. M.: North American CO₂ exchange: inter-comparison of modeled estimates with
101 results from a fine-scale atmospheric inversion, *Biogeosciences*, 9, 457-475, 2012.
- 102 Huntzinger, D. N., Post, W. M., Wei, Y., Michalak, A. M., West, T. O., Jacobson, A. R., Baker, I. T.,
103 Chen, J. M., Davis, K. J., Hayes, D. J., Hoffman, F. M., Jain, A. K., Liu, S., Mcguire, A. D.,
104 Neilson, R. P., Potter, C., Poulter, B., Price, D., Raczka, B. M., Tian, H. Q., Thornton, P.,
105 Tomelleri, E., Viovy, N., Xiao, J., Yuan, W., Zeng, N., Zhao, M., and Cook, R.: North American
106 Carbon Program (NACP) regional interim synthesis: Terrestrial biospheric model
107 intercomparison, *Ecological Modelling*, 232, 144-157, 2012.
- 108 Kitanidis, P. K.: Quasi-Linear Geostatistical Theory for Inversing, *Water Resources Research*, 31, 2411-
109 2419, 1995.
- 110 Krinner, G., Viovy, N., De Noblet-Ducoudré, N., Ogée, J., Polcher, J., Friedlingstein, P., Ciais, P., Sitch,
111 S., and Prentice, I. C.: A dynamic global vegetation model for studies of the coupled atmosphere-
112 biosphere system, *Global Biogeochemical Cycles*, 19, GB1015, 2005.
- 113 Mueller, K. L., Gourdjji, S. M., and Michalak, A. M.: Global monthly averaged CO₂ fluxes recovered
114 using a geostatistical inverse modeling approach: 1. Results using atmospheric measurements,
115 *Journal of Geophysical Research: Atmospheres*, 113, D21114, 2008.
- 116 Van Der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Kasibhatla, P. S., and Arellano Jr, A. F.:
117 Interannual variability in global biomass burning emissions from 1997 to 2004, *Atmos. Chem.*
118 *Phys.*, 6, 3423-3441, 2006.
- 119 Zeng, N., Mariotti, A., and Wetzal, P.: Terrestrial mechanisms of interannual CO₂ variability, *Global*
120 *Biogeochemical Cycles*, 19, 2005.

121

122